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# GEOLOGIC REPORT NO. 34

Geology and Geochemistry Diana Lakes Area Western Talkeetna Mountains, Alaska

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# GEOLOGY AND GEOCHEMISTRY OF

# THE DIANA LAKES AREA

# WESTERN TALKEETNA MOUNTAINS, ALASKA

By

Robert E. Anderson

# ABSTRACT

Highly fractured granodiorite and greenstone with some later volcanic and related rocks lie along the western edge of the Talkeetna Mountains between Iron Creek and Sheep River, west of Rainbow Lake.

The granodiorite forms a discordant northwest-trending stock which interrupts the general northeast structural trend of the Talkeetna Mountains.

Results from geochemical sampling of stream sediments and rock exposures were put into the University of Alaska's IBM 360 computer for statistical calculations. Trend surface analyses through the fifth degree from the rock sample data were drawn. Comparison of the stream sediment samples with the rock sample trend surfaces and their residuals indicate possible hydrothermal mineralization at depth in several locations. These are viewed as possible target areas for future ore mineral exploration.

## INTRODUCTION

An area located on the western edge of the Talkeetna Mountains between Sheep River on the southwest and Iron Creek on the northeast, called the Diana Lakes area, was selected for geologic and geochemical investigation as a continuation of a project to examine the iron stained areas in the Talkeetna Mountains reported by Rose (1965). Iron staining or capping has been indicative of a number of hydrothermal and porphyry mineral deposits in the western United States. Another encouraging feature for possible mineralization came from a study of aerial photos of the area which indicated several major faults near the iron stained sections. The southeast side of the mapped area (fig 1) lies about seven miles north-northwest of a known area of copper mineralization along Iron Creek.

Before the work of Rose (in 1965 and 1967), very little geologic work had been done in the western Talkeetna Mountains. The earliest geologic map of the area was done by S. R. Capps in 1919 ( p 187-205). That map is very generalized and on a scale of 1 to 500,000. In 1925 some of the prospects in the Talkeetna Mountains were visited and a brief report accompanied by sketches of the properties were submitted to the Alaska Railroad (Townsend, 1925). A more complete report but still quite general was made by Capps in 1940. The geologic map in that report is on a 1:250,000 scale. There has been no other geologic mapping on a broad scale since Capps' 1940 report. Grantz (1960 A & B) investigated the southeastern corner of the Talkeetna Mountains quadrangle putting out two geologic maps on a 1:48,000 scale.

The Diana Lakes area has been highly fractured and faulted and there are numerous 25 to 75 foot scarps making good rock exposures above 2500 feet elevation. Stream sediment sampling was difficult below the 2500 foot level because of heavy growth of alder and willow brush.

The Division of Mines and Geology field party was in the Diana Lakes area for a total of 26 days from July 11 to August 26, 1968.

The writer was assisted by Gardner F. Gillespie, III, who collected the stream sediment samples and assisted in collecting rock samples. Access to the area was by float plane. Don Sheldon of Talkeetna Air Service worked out the transportation and support problems.

# GENERAL GEOLOGY

A highly fractured area of principally greenstone and intrusive rock lies between Iron Creek and Sheep River (fig 1). Earlier geologic maps (Capps, 1940 and Dutro and Payne, 1957), show all but the southeast side of this area as part of a batholith trending NE through the Talkeetna Mountains. The present work shows that part of what was previously mapped as a granitic batholith is in fact greenstone with minor thin interbedded meta-argillite and pyroclastic rocks. The intrusive forms a discordant stock trending NW. It is probably separate from the concordant NE trending granitic batholith shown by Capp (1940) for the Talkeetna area. Emplacement of the stock resulted in doming of the overlying lavas and associated sills which nearly erased the general NE structural trend of the Talkeetna Mountains. During or following emplacement of the stock, faulting occurred which placed the country rock into isolated contact with the intrusive rocks. Magmatic stoping may have been the mechanism for the fracturing and isolation of large blocks of country rock in the intrusive. Metamorphism of the country rock to greenstone probably occurred during emplacement of the stock. The greenstone is generally blocky, showing little shistosity or bedding, i.e., a hornfels, and in places grading into an andesite granofel or a gabbro granofel. The most recent magmatic rocks in the area are rhyolite volcanic flows and several rhyolite dikes. Associated with the volcanic activity is a tuffaceous sandstone.

Glacial advances covered the area during middle to late Pleistocene (Coulter and others, 1965). Much of the southeast portion of the mapped area (fig 1) is rounded and smoothed with numerous glacially-transported boulders appearing in the tundra. In much of the area north of Rusty Creek (fig 1) numerous scarps and sharp ridges indicate post-glacial movement along the northwest and east-west trending fault zones.

### STRUCTURAL GEOLOGY

### Faults

Two major fault directions are apparent in the mapped area. Steeply-dipping, nearly east-west trending faults and fractures appear from aerial photographic study to be restricted to the area around the intrusive stock and appear to have been caused by the stock emplacement. Dips range from vertical to 55°S; no north dips were found. The other prominent faults of the area trend northwest with nearly vertical dips ranging from 80°SW to 75°NE. At many places dips on the northwest trending faults were impossible to measure.

Structures trending northwest are of a more regional nature than the east-west faults. Northwest trending structures are apparent on aerial photographs and on topographic maps in many places throughout the central to southwestern part of the Talkeetna Mountains quadrangle. The northeast trending regional structure so apparent along the upper Talkeetna River (Anderson, 1969) was almost obliterated by the stock emplacement in all but the southeast portion of the mapped area (fig 1). The greenstone-granite contact in the southeast side of the map (fig 1) follows the regional northeast trend. One fault and an associated rhyolite dike southeast of Diana Lakes may also reflect the northeast regional trend. Intense faulting and fracturing parallels both the northwest and the east-west structural trends. Only the major faults are plotted on the map (fig 1).

#### Joints

A study of jointing in the area shows that most of the joints measured are steeply dipping with no particular directional trend indicated. Rock exposures along the perimeter of the mapped area (fig 1) are infrequent, but jointing in the exposures observed tend to outline the doming effect of the stock emplacement.

#### PETROLOGY

No age dating on any of the rocks in this area was done. Ages tentatively assigned to the various formations are arrived at by the inference that similar rock types correlate with the dated rocks in the southeast part of the Talkeetna Mountains quadrangle (Grantz, 1960).

To aid the petrologic discussion, various rock samples were studied in thin section by polarizing microscope and selected samples were analyzed by x-ray diffraction for major minerals by the Division of Mines and Geology Laboratory, College, Alaska.

# Metamorphic Rocks

Greenstone (g) -- Greenstone is the principal metamorphic rock in the mapped area. The greenstone in this area is tentatively placed in the Talkeetna formation of lower Jurassic age. The greenstone varies from an orthoandesite granofel which shows good flow structure in a few places and is porphyritic or porphyroblastic in many places, to an orthodiorite granofel. In most places metamorphism of the volcanic and nearsurface intrusive andesite and diorite is low order. In a few places, interbedded with the granofels, are thin outcrops of rock showing phyllitic texture.

In thin section the orthodiorite shows abundant plagioclase feldspar, moderate to abundant green hornblende (tremolite?), moderate to abundant epidote, moderate chlorite, and moderate to minor quartz. Accessory pyrite was noted in several places in the field and accessory sphene was observed in thin section. In thin section the orthoandesite shows abundant feldspar, probably plagioclase but difficult to determine due to metamorphism, moderate to abundant green hornblende (tremolite?), moderate chlorite, and moderate epidote. Quartz is a minor constituent of the orthoandesite. Accessory pyrite was noted. In places pyrite is surrounded by accessory hematite. In the northwest section of the mapped area hematite veinlets were also observed.

In a number of places interbedded with the greenstone, but not shown on the map (fig 1), are thin beds of slightly metamorphosed arkose and pyroclastic rocks. Neither the arkose or the pyroclastic rocks contain any optically apparent pyrite or other sulfide minerals.

# Intrusive Rocks

Granodiorite (gd) -- The intrusive stock ranges from quartz monzonite, through granodiorite to quartz diorite. No contacts or distinct changes were observed, so the rocks are considered all part of a single intrusive stock and are referred to as granodiorite in the text and on the map (fig 1). In age, the stock may be related to the Talkeetna Mountains batholith, placing it between lower and middle Jurassic in age. However, the author feels that the stock is slightly younger than the batholith since the stock breaks up the general NE structural and stratigraphic trend probably caused by emplacement of the batholith. In thin section a sample of quartz monzonite from the vicinity of the Diana Lakes contains in decreasing abundance; plagioclase showing albite and Carlsbad twinning; abundant to moderate amounts of quartz; moderate green hornblende showing good cleavage and strong pleochroism; moderate epidote and chlorite; moderate to minor orthoclase(?); accessory apatite; sphene; and pyrite,

Quartz diorite in thin section from samples taken near the center of the mapped area, (fig 1) contains in decreasing abundance; plagioclase showing excellent albite and Carlsbad twinning in addition to normal zoning; anhedral quartz; strongly pleochroic green hornblende, in places showing brown alteration rims; augite(?) showing moderate relief, and faint green color, strongly pleochroic, brown biotite; and abundant accessory pyrite, altered in places to hematite.

In several places greenstone inclusions were noted in the granodiorite rocks. At most places they are only a few inches across, and in a few places they are a foot or over in longest dimension. One very large dike-shaped inclusion of greenstone is shown in the northwest portion of the map (fig 1).

# Volcanic and Related Rocks

Rhyolite(?) (rh) -- The volcanic rocks and dike rocks observed are porphyritic rhyolite, possibly middle to upper Tertiary in age. The dikes are generally a drab, medium dark, grayish brown, while the lavas are tan and frequently iron stained, particularly along Rusty Creek (fig 1). Due to the fine grained nature of these rocks, thin section mineral identification was difficult. Many of the phenocrysts were identified as plagioclase crystals. One point that was noted in all of the samples collected is the conspicuous shortage of mafic minerals. Examination of the rhyolite by x-ray diffraction showed the following minerals listed in decreasing order of abundance; quartz; feldspar (in some places feldspar was slightly more abundant than quartz); muscovite; and chlorite. In the field, accessory pyrite up to 10% was noted in places along Rusty Creek which accounts for the prominent iron staining. Pyrite in such abundance is not normal in rhyolite and may be of epigenetic origin.

Tuffaceous Sandstone (ts) -- Tuffaceous sandstone is an irregular bed confined to the central eastern part of the mapped area (fig 1). The tuffaceous sandstone shows mostly quartz, minor feldspar, and sercite in a quartz matrix giving the appearance of a sub-graywacke. In some places toward the greenstone contact (fig 1) limonite staining which results from weathering of amphibole and biotite is apparent. In other places toward the granite contacts, the amphiboles and biotite are not weathered and form as much as 1% of the rock. Some quartz has been recrystallized.

# GEOCHEMICAL INVESTIGATIONS

The results of the stream sediment sample analyses were put into a computer to tabulate the data and to calculate the statistical characteristics. The program for this was written by L. E. Heiner, Mining Engineer, Mineral Industry Research Laboratory, University of Alaska. Rock sample values were put into the computer on a program to calculate trend surface analyses. The IBM 360 Model 40 computer at the University of Alaska performed the computations. To compare stream sediment geochemical results with rock sample geochemical results, threshold and anomalous stream sediment values were plotted on the fifth degree trend surface analysis maps computed from the rock sample values. All samples were given numbers in the field and no changes were made for the maps or the computer programs.

## STREAM SEDIMENT SAMPLES

Stream sediment samples were collected from all flowing creeks in the mapped area (fig 1). The samples were dried in the College Laboratory of the Division of Mines and Geology, then forwarded to the U. S. Geological Survey Field Laboratory in Anchorage to be analyzed for thirty elements by semiquantitative spectrographic methods. Only twenty-three elements were detected and the results are shown in Appendix 1. Intervals of estimate and detection limit for this method of analysis are shown in Appendix 2. Eliminated from Appendix 1, which shows the computer calculation results of threshold and anomalous values, are silver (Ag), titanium (Ti), and beryllium (Be), because not enough of the samples showed these elements to be of statistical significance. For the remaining twenty elements the mean and the standard deviation were calculated. From these measures of central tendency, the threshold value, or upper limit of normal background fluctuation, and anomalous values were determined for each element. The computer also plotted a histogram of frequency distribution for copper (Appendix 3).

The threshold and anomalous values for each element were computed by methods described in Hawkes and Webb (1962, p 30). The threshold value is taken as the mean plus twice the standard deviation; anomalous values are taken as the mean plus three standard deviations. These values are meaningful for a normal distribution, the further the data departs from normalcy the less reliable are the computed threshold and anomalous value. Appendix 1, for stream sediment samples, shows the calculated average, standard deviation, threshold, and anomalous values for each element detected.

The concentration of an element in a given sample is either in the background range, between the threshold value and anomalous value, or greater than the anomalous value. Samples are considered possibly anomalous if the concentration is between the threshold value and the anomalous value and probably anomalous if the concentration of an element is above the anomalous value. A listing of all possible and probable anomalous values is shown on Table 1. Locations of all samples taken are shown on the map (fig 1). Stream sediment samples containing probable anomalous amounts of copper, lead, zinc, nickel, chromium, cobalt, and molybdenum are indicated, and samples which have two or more of the above listed elements in possible anomalous concentrations are also indicated.

# Discussion of Major Elements

Copper -- The average copper content of 50 ppm (Appendix 1) is in the range of what would normally be expected from soil samples (Hawkes & Webb, p 365). The average of 50 ppm is slightly below another section of the Talkeetna Mountains, (Rose, 1967), where, by discounting two very high copper anomalies, the average copper content is 80 ppm. In the Diana Lakes area the highest concentration of copper lies in the northwest part (fig 1). The samples were taken near the point of convergence of three faults. Bedrock in the vicinity of the samples is greenstone, which in this vicinity is higher in quartz content than most of the other greenstone observed.

Lead -- The lead average of 22 ppm (Appendix 1) is about what would be expected from soil samples (Hawkes & Webb, p 367) and is well above the lead average of 7.57 ppm in another section of the Talkeetna Mountains (Rose, 1967). There are no probable lead anomalies, but nine samples are above threshold and are considered possibly anomalous. Most of the possibly anomalous samples come from the Rusty Creek drainage (fig 1). Bedrock crossed by this drainage is granodiorite and rhyolite. The rhyolite in places shows abundant accessory pyrite. Manganese -- Two manganese anomalies occur with high lead, zinc, molybdenum, and other elements in sample numbers 81 and 82 (table 1).

Zinc -- The zinc average of 52 ppm (Appendix 1) is below the average of 93 ppm obtained in another section of the Talkeetna Mountains (Rose, 1967). It is about average for soil samples as defined by Hawkes and Webb (p 376). Threshold samples 76, 78, 82, and anomalous sample 81, are from the Rusty Creek drainage and are apparently associated with lead occurrences.

Molybdenum -- The molybdenum average of 6 ppm (Appendix 1) is well above the average of 2 ppm from another section of the Talkeetna Mountains (Rose, 1967). The average of 2 ppm is also what would be expected from normal soils (Hawkes & Webb, p 369). The single probable molybdenum anomaly, sample 70, occurs on Rusty Creek, upstream from the lead and zinc concentrations (fig 1) and is not associated with any other threshold or anomalous metal concentration (fig 2). Of the threshold molybdenum samples, numbers 38 and 73 are not associated with other metals, number 50 is associated with threshold copper and threshold lead; numbers 81 and 82 are associated with lead and zinc along Rusty Creek.

Nickel -- The nickel average of 25 ppm (Appendix 1) is just slightly lower than the average of 33 ppm obtained in the other section of the Talkeetna Mountains (Rose, 1967) and well below what would be expected from streams cutting greenstone (Hawkes & Webb, p 371). Only two samples lie above threshold value (Appendix 1) and both of these are probable anomalous samples 31 and 61. Both of these samples are associated with anomalous chromium. Sample 31 is also associated with threshold copper in the northwest section of the mapped area (fig 1). Sample 61 is on Gard Creek which drains the lowest of Diana Lakes.

Chromium -- The chromium analysis average of 56 ppm (Appendix 1) is far below what would normally be expected. Soil sample average is given at 200 ppm (Hawkes & Webb, p 363). Three anomalous chromium samples are recorded (Appendix 1). Sample 31 is associated with threshold nickel, cobalt, and copper; sample 32 is associated with above threshold copper and cobalt; and sample 61 is associated with above threshold nickel.

Cobalt -- The cobalt average of 15 ppm (Appendix 1) is slightly above the average given for cobalt soil samples (Hawkes & Webb, p 363). Samples 31 and 32 are anomalous (Appendix 1) and are associated with above threshold values of copper and chromium (table 2). Sample 82 is a threshold reading and is associated with threshold lead, molybdenum, and zinc.

Discussion of Trace Element Associations

Geochemical mineral associations are specific for some rock types, both sedimentary and igneous, and for most hydrothermal sulfide ores. At the surface, supergene mobility of the existing elements is dependent on the minerals in which they occur and the conditions of weathering (Andrews-Jones, 1968, p 5). Consequently, stream sediment samples are a result of both weathering and the original rock element associations.

Table 2 shows the possible and probable anomalous occurrence of element associations from the stream sediment samples. A few of the associations may reflect their origin. The base metals in sediment samples 31, 32, and 61 are associations which are found in either a hydrothermal sulfide ore or in some ultramafic rocks. The associations of molybdenum with lead, copper and/or zinc in samples 50, 81, and 82 reflect hydrothermal sulfide mineralization (Andrews-Jones, 1968, p 7).

# TABLE ]

# Threshold and Anomalous Stream Sediment Samples

Chemical Symbol not underlined = Possible Anomalous Value

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Chemical Symbol underlined = Probable Anomalous Value

Chemical Symbol		Element	Sample No.	Elements
Ag Ba Be Co Cr Lag Mo Nb I D S S Ti V Y Z r	<b>н н</b> аридин и и и и и и и и и и и и и и и и и и	Silver Boron Barium Beryllium Calcium Cobalt Chromium Copper Iron Lanthanum Magnesium Molybdenum Manganese Niobium Nickel Lead Scandium Strontium Titanium Vanadium Yttrium Zinc Zirconium	23 24 23 33 33 44 44 45 78 90 14 55 60 15 66 90 23 56 77 77 78 91 24 58 88 89	Mg, Sr B Fe, Sc, V Co, Cr, Cu, Fe, Mg, Ni, Sc, Sr, V Co, Cr, Cu, Mg Cu, Mg, Sr Sc B, Mo B Sr Fe, Mg, Mn, Sr, V Zn Fe, Zn Ba, Fe Nb Ba, Cu, La, Mo, Pb Ba Ba, Cu, Fe Nb Zn Mg, Sr Cr, Ni, Sc, V, Zr B B Nb Mo Nb Mo Pb Pb, Zn B, Pb Fe, La, Mn, Y, Zn Fe, Sc La, Mn, Mo, Pb, Y, Zn Ba, Co, Fe, Mn, Mo, Pb, Sc, Y, Zn Pb Pb

#### ROCK SAMPLES

Rock grab samples taken throughout the area are shown on figure 1. Samples were taken of every rock type and 75 samples were analyzed by quantitative atomic absorption by the Division of Mines and Geology laboratory for gold, copper, lead, and zinc (Appendix 4). The geochemical data were programmed through the University of Alaska IBM computer as described previously to calculate and plot trend surfaces through the fifth degree. The first four degrees of each element are shown in figures 3 through 6 in greatly reduced scale to allow the reader to see the development of the fifth degree trend surfaces in figures 7 through 10.

Table 2 shows the average gold, copper, lead, and zinc concentrations in each of the generalized rock types in the area and comparisons with the average concentrations that would be expected from these rock types as described by Hawkes & Webb (p 359-367).

#### Discussion of Metal Analyses

Copper -- The average copper found in both granodiorite and rhyolite in this area is about average for felsic rocks according to Hawkes and Webb (1965, p 364). The highest copper sample in the granodiorite (no. 119) shows 173 ppm. The highest copper sample in the rhyolite (G-3) shows 80 ppm. Most of the high copper values lie in the greenstone. The greenstone has a very wide range of copper occurrences from a low of 3 ppm to a high of 240 ppm. The average copper content in the greenstone is just a little over one-half of the average for mafic rocks, according to Hawkes and Webb (1965, p 364). Fifth degree trend surface analysis (fig 7) shows the highest background in the northwest portion of the mapped area.

Lead -- The average lead content for the area is slightly below average for both mafic and felsic rocks (Hawkes & Webb, 1965, p 367) with granodiorite (sample 138) carrying the highest content of 242 ppm. The lead content in the rhyolite rocks is only one-third of normal lead occurrence in felsic rocks. Fifth degree trend surface analysis (fig 8) shows the highest background on the west side of the mapped area trending in a northwestsoutheast direction.

Zinc -- Granodiorite and rhyolite are about average in zinc content for felsic rocks while the greenstone is less than one half the normal average for mafic rocks (Hawkes & Webb, 1965, p 376). Fifth degree trend surface analysis (fig 9) shows the highest background in the southern portion of the mapped area.

Gold -- The average for gold is very roughly calculated, using a figure of .004 ppm where gold was not detected since the detection limit is .008 ppm on the Division of Mines and Geology atomic absorption equipment. Granodiorite in the Diana Lakes area runs over twice as high in gold as would normally be expected (Hawkes & Webb, 1965, p 365). The average is high because of one sample, G-37, which contains .38 ppm gold. Without sample G-37 the average would be about .013 ppm which is only slightly above average. Greenstone in the area is generally metamorphosed mafic volcanic rocks and the figure from Hawkes and Webb (1965, p 365) that is used here (.035 ppm) comes from mafic igneous rocks. From this standard, the greenstone is below average in gold content. Rhyolite in both flows and dikes, is very slightly above average. Fifth degree trend surface analysis (fig 10) shows the highest background of gold in the central-aastern portion of the mapped area.

# TABLE 2

# Comparison of Diana Lakes Rock Type Mineral Averages With Rock Type Mineral Averages from Hawkes and Webb

#### Parts Per Million

	Go	1d	Cor	per	Le	ad	Zi	inc
	Diana Lakes	Hawkes & Webb						
Granodiorite 33 samples	.025	.01	27	30	26	48	65	60
Greenstone 30 samples	. 020	.035	76	140	7۲	12	52	130
Volcanic Flows & Associated Rocks 12 samples	.015	.01	31	30	14	48	52	60

# Trend Surface Analysis

In order to outline area trends and point to the best possible target areas for future mineral exploration the computer was called upon to make a statistical analysis of the geochemical data in terms of location and concentration of the following elements. For trend surface analysis, the method of least squares has the best general application for surface fitting (Crow and others, p 151). By using this technique a planar surface is established for geochemical data such that positive and negative measurement from the surface are at a minimum.

Multiple regression trend surface analyses of geochemical data have been examined in a number of the known mineral deposits in the United States (M. P. Nackowski and others, p 1077) and frequently have been shown to outline the known mineralized area as well as point to petrologic contacts, faults, folds, or indicate possible topographic relationships.

"Trend surfaces applied to geological (and geochemical) data....are described by arithmetic equations in which the geologic data represent the dependent variable stated as a function of two independent variables which establish the areal sample locations. The mathematical operations of trend-surface fitting consist of finding the constants to the arithmetic equations such that the least-squares criterion is satisfied. In fitting a trend surface to geochemical data, the dependent variable z, represents quantity of indicator element in each sample and the independent variables, x and y, represent the planar sample location coordinates. These equations for linear, quadratic and cubic components are shown on table 3.

"Trend surfaces classified according to degree are more complex as the number of components in the equations describing them increases..... The first degree surface is a plane and contains only linear terms, whereas the generalized second degree surfaces, which contain both quadratic and linear terms, are either positive or negative bowl-shaped paraboloids. Third degree surfaces include an inflection, are more complex and contain cubic, quadratic, and linear terms". (Nackowski and others, 1967, p 1078)

# TABLE 3

# Classification of Trend Surface Equations Illustrating Three Components and Three Degrees (from Nackowski and others, 1967, p 1078, Table 2)

Trend Surface Classification	Dependent Varia <u>b</u> le	Linear Component	Quadratic Component	Cubic Component
First degree plane surface	2 Z=	A+B <sub>x</sub> +C <sub>y</sub>		
Second degree paraboloid	Ζ=	A+B <sub>x</sub> +C <sub>y</sub> +	D <sub>x</sub> <sup>2</sup> +E <sub>xy</sub> +F <sub>y</sub> <sup>2</sup>	
Three degree surface	z=	A+B <sub>x</sub> +C <sub>y</sub> +	Dx <sup>2+E</sup> xy+Fy <sup>2+</sup>	<sub>Gx</sub> <sup>3</sup> xH <sub>x</sub> <sup>2</sup> y+I <sub>xy</sub> <sup>2</sup> +J <sub>y</sub> 3

Letters A through J represent constants.

Each increasing degree adds another inflection in the trend surface and each equation representing this becomes more complex; the independent variables, x and y, are increased in power and new constants are fit into this part of the equation which is simply added on to the lower power equations, following the pattern shown in table 3.

Figure 2 shows the generalized trend surfaces for the first three degrees and variables.



Figure 2 shows generalized trend surfaces for the first three degrees and variables.

# Figure 2

Generalized trend surfaces for three surfaces for three degrees and other variables. (After Nackowski and others, 1967.)

"The trend surface itself represents the regional component or geochemical trend. This surface, or the value of any point on the surface represents a threshold value which is variable across the map area." (Heiner, written communication, 3-25-69)

Values which fall above or below the trend surface in any degree are termed residuals. Each advancing trend surface degree will generally cover some of the residuals from the degree below it.

Persistent residuals through increasing trend surface degrees are considered possible target areas for future prospecting and mineral exploration. Residuals standing well above the fourth and fifth degree trend surfaces for each of the four elements are shown in the following figures. On the fifth degree trend surfaces for copper, lead, and zinc, in addition to the residuals, probable and possible stream sediment anomalous values are plotted. Discussion and comparison of the results are included on the figure for each different element.



# Figure 3 COPPER TREND SURFACE ANALYSES

1



TREND SURFACE ANALYSES



ZINC TREND SURFACE ANALYSES



# Figure 6 GOLD TREND SURFACE ANALYSES







FIFTH DEGREE - the final degree in trend analysis indicates an increased trend toward the north and west and a dome on the eastern side. A decreased trend is indicated on the eastern and south-Granodiorite ern sides. Two strong positive residuals occur within high trend Granodiorite areas and may be of significance for possible future exploration. Gold was not detected in the stream sediment samples. How-Rhyolite ever, by semiquantitative spectrographic analysis the ronodiorite detection limit for gold is ten parts per million. In Rhyoli other sections of Alaska 0.04 parts per million is Greenstone the calculated average so a lack of gold detection 110 DIANA by semiquantitative spectrographic analysis is not Granodiorite AKES significant. It is noted that copper anomalies Rhyc show in both areas of strong residuals. The two possible copper anomalies present in ō the dome area may be good pathfinders 50 Granòdiorite for gold particularly since they are 1.stone 54 54 not associated with high copper sur-Tuttoccous Sandstone face trends or residuals. Granodiarite -N-LEGEND 5th degree rock sample Greenstone residuals ~ 65 ppm or more Possible stream sediment Rhyolite copper onomaly 2 Miles 0 Probable stream sediment Granodiorite copper anomaly APPROXIMATE SCALE Rhy<sup>olite</sup> REFERENCE CONTOUR = 50 PPM eenston RAINBOW LAKE Greènstone Figure 10 GOLD TREND SURFACE ANALYSIS

# CONCLUSIONS AND SUGGESTIONS

Geochemical analyses of stream sediment samples and rock samples indicate a few target areas for future metal exploration. To determine the significance and economic potential of the target areas which are discussed in figures 7 through 10, further geochemical or geophysical exploration is suggested. The trend surface analyses point to hydrothermal mineralization related to a general northwest structural trend.

No base metal minerals other than pyrite were observed in the field. If the hydrothermal mineralization developed an area of economic interest, it is likely to be found at depth.

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MENT SAMPLES	US VALUES	10.69 179.59 29.86 7.08 27.60 151.51	4.75 322.74 55.04 55.04 20.24 55.04	20.19 445.10 21.91 27.54 68.09 261.59 261.59	24.94 5/7.85 22.93 16.63 316.63 316.63	ds described	n plus twice the ndard deviations. ta departs from			
Appendix I A OF GEOCHEMICAL STREAM SEI	CO Cr Ni M Fet M91 Can Ba Sr	55.78         25.25         1200.29         8.85         1.76         2.19         362.1	30.76         1a.20         874.40         3.37         269.96         167.           7.20         874.40         3.37         0.59         0.72	9.02 53.64 2949.09 15.60 2.94 3.62	6.22 6.22 6.23 50 18.97 3.53 6.3.5	rmbols see Table I oid and anomalous values for each element were computed by m	and Webb (1962, p.30). The threshold value is taken as the r eviation: anomalous values are taken as the mean plus three es are meaningful for a normal distribution, the further the is less reliable is the computed threshold and anomalous value the less reliable is the computed threshold and anomalous value	LEGEMD	T - indicates threshold value.	<ul> <li>A - indicates anomalous value.</li> <li>E - indicates number eliminated due to error in assay report or in computer card punching.</li> </ul>
TABULATION		Average 50.2) 22.33 51.85 6.30	41.15         41.45         10.16         3.96           Standard         Deviation         3.96         3.96	Hawkes & 132.51 Webb Threshold Mean plus 2 standard deviations 2	Hawkes & 173.66 Webb Anomaly Aromaly Aran plus deviations deviations	For definition of chemical sy The thresh	in Hawkes a standard do These value normalcy th			
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Appendix 2

# Semiquantitative Emission Spectrographic Analyses Intervals of Estimate and Detection Limit

Parts Per Million

Be	1,000 500 100 70 70 70 70 70 70		Zr	1,000 500 700 50 10 L
Ba	5,000 2,000 1,000 200 200 200 200 200 200		٢	200 200 200 200 200 200
B	2,000 1,000 200 100 20 100 100		Λ	10,000 5,000 1,000 100 200 100 200 100 100 100
Cr	2,000 1,000 100 100 100 100 100 100 100 1		Ti	1 0.5 0.005 0.001 0.001 1 0.001
Co	2,000 1,000 500 100 100 100 100 100	4illion	Sr	5,000 2,000 1,000 200 50 100 50
ίN	5,000 500 500 1000 1000 1000 1000	Parts Per l	Sc	100 20 10 20 10 20
Мо	2,000 500 100 100 100 100 100		Mn	5,000 2,000 1,000 100 200 200 200 200 200 200 200 200
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Pŀ			La	, r r 200 200 200 200 200 200 200 200 200 200
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Си	20,000 5,000 10,000 1,000 100 100 100 100 100 10		Са	20 0.03 0.00 0.02 0.02 0.02 0.02 0.02 0.

For definition of chemical symbols see table l

N = Not detected
L - Present, but below determination limit

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# Appendix 4

# Tabulation of Geochemical Rock Samples with their Threshold and Anomalous Values

Field and Map Sample No.	Rock Type	ppb <u>Gold</u>	ppm Copper	Lead	Zinc
Field and Map <u>Sample No.</u> G-3 G-4 G-12 G-22 G-31 G-37 G-50 G-74 G-75 G-89 G-94 G-100 G-111 G-113 G-127 a & b G-129 2 4 6 9 12 13 17 21 23 24 33 34 36 39 44 59 60 62 63 67 70 78 83 96 103* 106 119	Rock Type V Gr Gr Gs	ppb         Gold         25         ND         50         ND         25         380A         ND         ND         ND         ND         ND         16         12         27         16         ND         12         27         16         ND         12         20         20         30T         64         28         ND         37         28         40         30         37         28         40         30         32	ppm <u>Copper</u> 80 25 32 12 10 15 22 55 75 25 33 15 47 20 23 25 25 25 112 184A 121 20 43 240A 52 47 28 67 22 47 28 67 22 47 17 62 30 43 240A 52 47 17 62 30 43 240A 52 142T 105 11 30 16 18 13 19 173T	Lead 10 20 25 18 12 25 14 20 90T 44 17 8 17 70 60 14 44 12 11 14 18 21 14 14 14 14 14 14 14 14 14 1	Zinc 68 61 65 52 48 56 92 82 62 28 312A 70 50 83 55 47 50 82 20 22 22 312A 70 50 83 55 47 50 82 20 22 22 312A 70 50 83 55 47 50 82 20 22 22 312A 70 50 83 55 77 50 82 20 22 22 312A 70 50 83 55 77 50 82 20 22 22 53 22 53 22 53 22 53 55 77 50 82 20 20 20 22 22 53 22 53 22 55 77 50 82 20 20 22 22 53 22 53 55 77 50 82 20 55 77 50 82 20 55 77 50 82 20 55 77 50 82 20 55 77 50 82 20 57 73 46 10 9 57 73 46 114 27 57 73 46 114 27 57 73 46 114 27 57 73 46 114 27 57 57 57 57 57 57 57 57 57 5
136 138 144	Gr Gr Gs	10 21 13	16 16 105	5 242A 19	51 1721 32

\*Not included in computer calculations

Field and Map <u>Sample No.</u> 148 151 155 159 162 164 167 168 184 187 191 195 203 205 206 212 216 223 233 235 236 240 244 248 253 260 265 267	Rock Type Gr Gs Gr Gr V Ss Gr Gr V Ss Gs Gr Gs Gr Gs Gr Gs Gr Gs Gr Ss Gr Gs Gr Gs Gr V Ss Gs Gr Gr V V Ss Gr Gr Gr Gr Gr Gr Gr Gr Gr Gr Gr Gr Gr	ppb         Gold         28         21         ND         10         ND         15         19         13         ND         ND </th <th>ppm <u>Copper</u> 17 22 19 13 22 35 28 17 18 30 25 15 17 31 28 31 40 10 12 16 160T 55 18 13 3 25 11 11 13 13 12 16 160T 55 18 13 3 25 11 11 12 16 160T 55 18 13 3 25 18 11 10 12 16 16 16 16 16 16 16 16 16 16</th> <th>Lead 12 19 6 8 8 10 10 7 4 6 15 6 7 5 15 17 12 11 9 14 12 13 5 15 15 15 15 15 15 15 15 15</th> <th>Zinc 68 37 76 100 25 174T 57 26 47 65 132 8 71 94 92 112 90 9 28 43 32 36 19 76 36 31</th>	ppm <u>Copper</u> 17 22 19 13 22 35 28 17 18 30 25 15 17 31 28 31 40 10 12 16 160T 55 18 13 3 25 11 11 13 13 12 16 160T 55 18 13 3 25 11 11 12 16 160T 55 18 13 3 25 18 11 10 12 16 16 16 16 16 16 16 16 16 16	Lead 12 19 6 8 8 10 10 7 4 6 15 6 7 5 15 17 12 11 9 14 12 13 5 15 15 15 15 15 15 15 15 15	Zinc 68 37 76 100 25 174T 57 26 47 65 132 8 71 94 92 112 90 9 28 43 32 36 19 76 36 31
Gs - Greenstone Gr - Granodiorite V - Volcanic ND for gold means less than 8 PPB IS means insufficient sample		T - Threshold defined b A - Anomaly a Webb	or possible y Hawkes and s defined by	anomaly as Webb Hawkes and	
PPB means parts per PPM means parts per	billion million				
		ppb <u>Gold</u>	ppm Copper	ppm Lead	ppm <u>Zinc</u>
Average		22	43	21	58
Standard deviation		46	45	30	46
Threshold (as defined b and Webb) average plu standard deviations	y Hawkes is two	114	133	81	150
Anomaly (as defined by and Webb) average plu standard deviations	Hawkes s three	160	179	111	196