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**DETAILED STRATIGRAPHIC SECTIONS OF THE MISSISSIPPIAN ENDICOTT
GROUP IN THE CENTRAL FRANKLIN AND EASTERN ROMANZOF MOUNTAINS,
ARCTIC NATIONAL WILDLIFE REFUGE, ALASKA**

by

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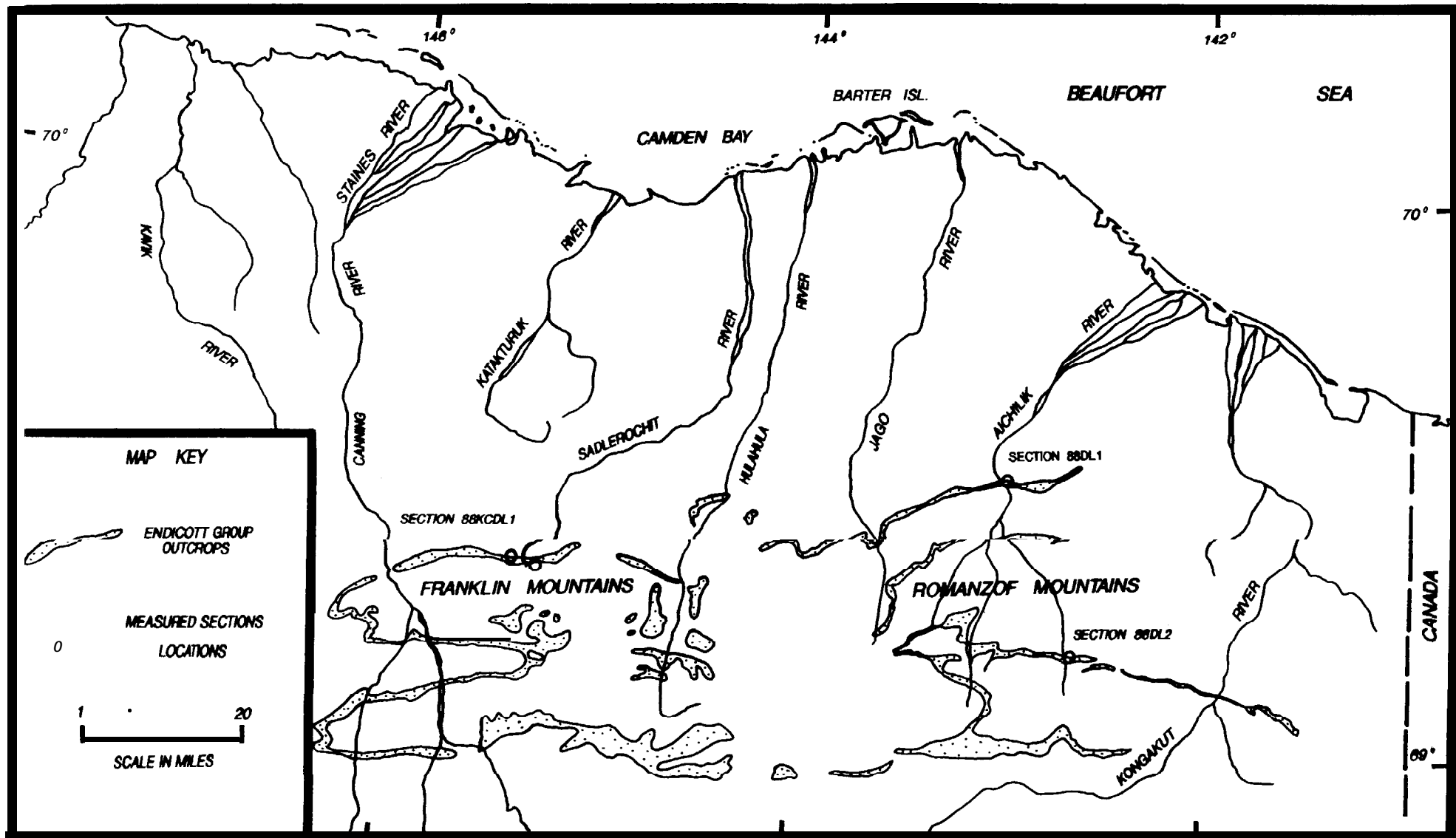
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INTRODUCTION

This report contains measured sections and preliminary environmental interpretations on the sedimentologic and stratigraphic evolution of the Mississippian Endicott Group in the Franklin and Romanzof Mountains, Arctic National Wildlife Refuge (Fig. 1). All data were collected during the 1988 field season. Measured sections and locations maps are included in the appendix at the end of this report. This project is part of an ongoing research effort in the Brooks Range by the Tectonics and Sedimentation Research Group at the University of Alaska-Fairbanks.

Regional Perspective

In the northeastern Brooks Range, as in the subsurface at Endicott Field (and possibly in Umiat and Meade basins located to the west), the Endicott Group is **autochthonous-to-parautochthonous** and unconformably overlies pre-Upper Devonian basement rocks of the Franklinian Sequence (**Nilsen, 1981; Moore and Nilsen, 1984; and Woidneck, et al, 1987**). In the central and western Brooks Range the Endicott Group is present in three north-vergent thrust sheets, within which it rests conformably above carbonates of the Baird Group (**Nilsen, 1981; Moore and Nilsen, 1984**). The stratigraphy and **sedimentology** of this dominantly siliciclastic succession changes markedly between these three areas (Fig. 2). Relative to its occurrence in the



Adapted from Donovan and Tailleux (1975)

Figure 1 - Exposures of the Mississippian Endicott Group north of the continental divide in the Arctic National Wildlife Refuge, Alaska (measured sections shown with open circles).

central Brooks Range and the north slope subsurface, the Endicott Group in the northeastern Brooks Range is poorly understood.

Information on the relationships between the various Endicott Group sequences is critical to our understanding of the Devonian and Mississippian depositional and tectonic evolution of northern Alaska, and the circum-Arctic region. A detailed analysis of the Endicott Group in northeastern Alaska will aid in determining these relationships.

PURPOSE AND OBJECTIVES

The purpose of this study is to gain a better understanding of the stratigraphy, sedimentology, paleogeography, and tectonic setting of the Endicott Group in northeastern Alaska. The objectives are:

- 1.) To determine the sedimentologic and structural nature of the pre-Mississippian unconformity, its relationship to the Kekiktuk Conglomerate and the structural evolution of the Franklin and Romanzof Mountains.
- 2.) To perform a petrofacies analysis of the framework constituents to determine their provenance.
- 3.) To carry out a detailed facies analysis of the Endicott Group, with emphasis on the internal geometry and lateral variations within the Kekiktuk Conglomerate, and the sandstones and **bioclastic/biohermal** limestones of the Kayak Shale. The facies analysis will also include characterizing the nature of the transition between the Kekiktuk

Conglomerate and the Kayak Shale, and the Kayak Shale and the Alapah Limestone (Lisburne Group).

4.) To determine the depositional environments recorded in the Endicott Group.

5.) To reconstruct the paleogeography of northeastern Alaska during Endicott time.

REGIONAL GEOLOGIC SETTING

The Endicott Group in the northeastern Brooks Range crops out along the northern and southern flanks; of several regional anticlinoria, which extend approximately from the Canning River eastward to the Canadian border and beyond (Bader and Bird, 1986). Extensive exposures of the complexly deformed and metamorphosed sediments and igneous rocks of the Franklinian Sequence occur within the core zones of these anticlinoria (Mull, 1982). The Endicott Group forms the base of the Ellesmerian Sequence in this region, and rests with distinct angular discordance above the metasediments and igneous rocks of the Franklinian Sequence (pre-Upper Devonian basement). Regional studies of the northeastern Brooks Range indicate that the pre-Upper Devonian basement rocks were deformed by thrust faulting during a Middle Paleozoic orogeny (Mull, 1982). Pre-Upper Devonian rocks were subsequently thrust northward during the **Cretaceous Brookian** Orogeny to form large duplex structures (W. K. Wallace, oral **commun.**, October and November 1988).

A limited body of literature exists on the geology of the northeastern Brooks Range, however, several informative papers are available for more detailed **discussions** of regional geology: Bird and Bader (1987); Bird and **Molenaar**; **Bowsher** and Dutro (1957); Brosge, et al. (1962); Reed (1968); Sable (1977). **Nilsen** (1981) and Moore and **Nilsen** (1984) provide an excellent general regional perspective of the Endicott Group in northern Alaska. Melvin (1987a and **1987b**) and Woidneck and others (1987) provide useful discussions of the depositional environments, petrography, and reservoir characteristics of the Kekiktuk Formation in the subsurface at Endicott Field.

STRATIGRAPHY

Excellent exposures of the Endicott Group occur along Straight Creek (Lat. 69 deg. 18 min. N and Long. 145 deg. **31** min. W) and on the east side of a large cirque-basin, located at the headwaters of the third major drainage west of Whistler Creek (Lat. 69 deg. 17 min. N and Long. 145 deg. 15 min. W), both in the Franklin Mountains (figure 1). Several partial sections of the Endicott Group containing portions of the Kekiktuk Conglomerate and Kayak Shale occur along a tributary to **Leffingwell** Fork in the **Romanzof** Mountains (Lat. 69 deg. 11 min. **N** and Long. 142 deg. 39.5 **W**).

Kekiktuk Conglomerate

The Kekiktuk Conglomerate west of Lake Peters consists of a **59-to-79** meter thick, fining upward succession of polymictic, chert and quartz-pebble conglomerate and medium- to very coarse-grained sandstone (measured sections **88KCDL1** and **88KCDL2**). Minor amounts of platy, black slate and **gray-black carbonate clasts** occur mixed in with the chert and quartz. The lower 40-to-45 meters is distinctly **fluvial** and contains multiple fining and coarsening upward cycles from 7- to-15 meters thick. Each cycle consists of an erosional lower contact which is overlain by a massive-to **CROSS-**stratified, multiply-erosional channel-fill succession. Individual channel-fills consist of imbricated, granule-to pebble conglomerate that generally fines upward to medium-to very coarse-grained sandstone. Near the tops of fining upward cycles, individual channel-fills grade upward from coarse-to very coarse-grained to medium-grained sandstone. Individual channel-fills range from < 0.5 meter to approximately 1.0 meter in thickness.

The top-most portions of many sandstone beds show evidence of extensive bioturbation and, near the top of the **fluvial** sequence, megaripples are preserved on some exhumed bedding surfaces. The bases of many individual channel-fill sequences cut into gray-black, organic-rich, argillaceous silty-to fine-grained sandstones 10-to-60 cm thick. These finer-grained lithologies contain abundant broken and carbonized plant fossils, show extensive bioturbation, and often exhibit flaser bedding. Similar fine-grained deposits

are observed as mud drapes overlying the tops of some channel-fill sequences. Cycle thicknesses range from five to 17 meters. These are interpreted to be either abandoned channel fill, slack water, or flood-plain deposits.

The upper-most 19-to-34 meters of the Kekiktuk Conglomerate consists of multiply-erosional channel-fill successions, as in the lower Kekiktuk (sections **88KCDL1** and **2**). However, they are restricted to the lower-most beds, and are composed of medium-to very coarse-grained sandstone, and grade upward near the top of the Kekiktuk into extensively bioturbated, medium-bedded, fine-to **coarse-grained** chert-quartzose sandstone. Many of the **finer-grained** uppermost beds are trough cross-stratified with internal ripple-drift cross-laminae. Gray-black **mudstone** interstratified with the finer-grained sandstones become more abundant near the top of this interval. Plant fossils progressively decrease in abundance upward in this interval.

East of Lake Peters, in the eastern Romanzof Mountains north of Bathtub Ridge, the basal portion of the Kekiktuk Conglomerate exhibits greater grain-size variability. In a disrupted section located **immediatetly** east of Leffingwell Fork, on the west bank of an unnamed tributary (section **88DL2** at Lat. 69 deg. 11 min. N and Long. 142 deg. 39.5 min. W), the basal beds of the Kekiktuk consist of polymictic, granule-to pebble-conglomerate, which grade upward into medium-to coarse-grained sandstone; whereas exposures of the lower-most beds of the Kekiktuk immediately west of the Aichilik River consist of polymictic, cobble conglomerate,

containing **clasts** as large as 30 centimeters.

The contact between the Kekiktuk Conglomerate and the Kayak Shale has **not** been precisely defined as yet, due to the interfingering nature of the two formations. We propose herein that the contact be placed at the base of the first prominent occurrence of shale, siltstone, or **mudstone** occurring above the uppermost interval of the Kekiktuk Conglomerate that is laterally continuous at outcrop scale.

Kayak Shale

The upper beds of the Kekiktuk Conglomerate interfinger with organic-rich shale, siltstone, and quartzarenite of the Kayak Shale (section **88KCDL1**). Measured thicknesses of the Kayak Shale west of Lake Peters range from **151-to-162** meters. These estimates may exaggerate the true stratigraphic thickness of the Kayak Shale, since in this region (as elsewhere) it serves as a detachment horizon along which there may have been structural thickening (Wallace, personal **commun.**, October 1988). Additional work is necessary to address the degree of structural thickening in this area.

The lower third (40-to-70 meters) of the Kayak Shale consists of siltstone and silty-shale, interrupted by several laterally continuous, coarse-grained intervals (sections **88KCDL1** and 2). The coarse-grained intervals consist of medium-to very coarse-grained, quartzose sandstone and chert-quartz granule-conglomerate that range from 1-to-16 meters in thickness. Internally, the coarser

grained beds are trough cross-stratified, contain **ripple-**drift lamination, are extensively bioturbated near their **tops**, and contain **strongly** undulatory bedding surfaces. Some of the thicker coarse-grained units contain large-scale lateral accretion surfaces (epsilon cross bedding).

The upper two-thirds (80-to-125 meters) of the Kayak Shale consists predominantly of gray-black, organic-rich, siltstone and shale, with **occasional** interbeds of bioclastic limestone (sections **88KCDL1** and 2). The bioclastic beds are composed of normally graded, broken and abraded crinoid and brachiopod fragments (0.1-1.0 centimeter thick in section **88KCDL1**, Franklin Mountains; and 10-to-100 centimeters thick southeast of section **88DL2**, Romanzof Mountains) that appear to have been deposited from episodic, waning-flow events. The Kayak Shale becomes increasingly calcareous upsection and, in sections **88KCDL1** and 2, is abruptly overlain by the Alapah Limestone (Lisburne Group).

ENVIRONMENTAL INTERPRETATION

The Endicott Group records a major transition from terrigenous **clastic** to carbonate depositional environments. The Kekiktuk Conglomerate represents deposition in braided, **fluvial** and deltaic systems. Retrograding environments gradually drowned the Kekiktuk dispersal system and initiated deposition of the Kayak Shale. The Kayak Shale represents an extensive suite of lagoonal, tidal and **storm-**dominated inner shelf environments. The uppermost beds in

the Kayak Shale consist of bioclastic and biohermal silty limestone. These record the final transition from terrigenous **clastic** deposition to deposition of the major carbonate platform of the Lisburne Group.

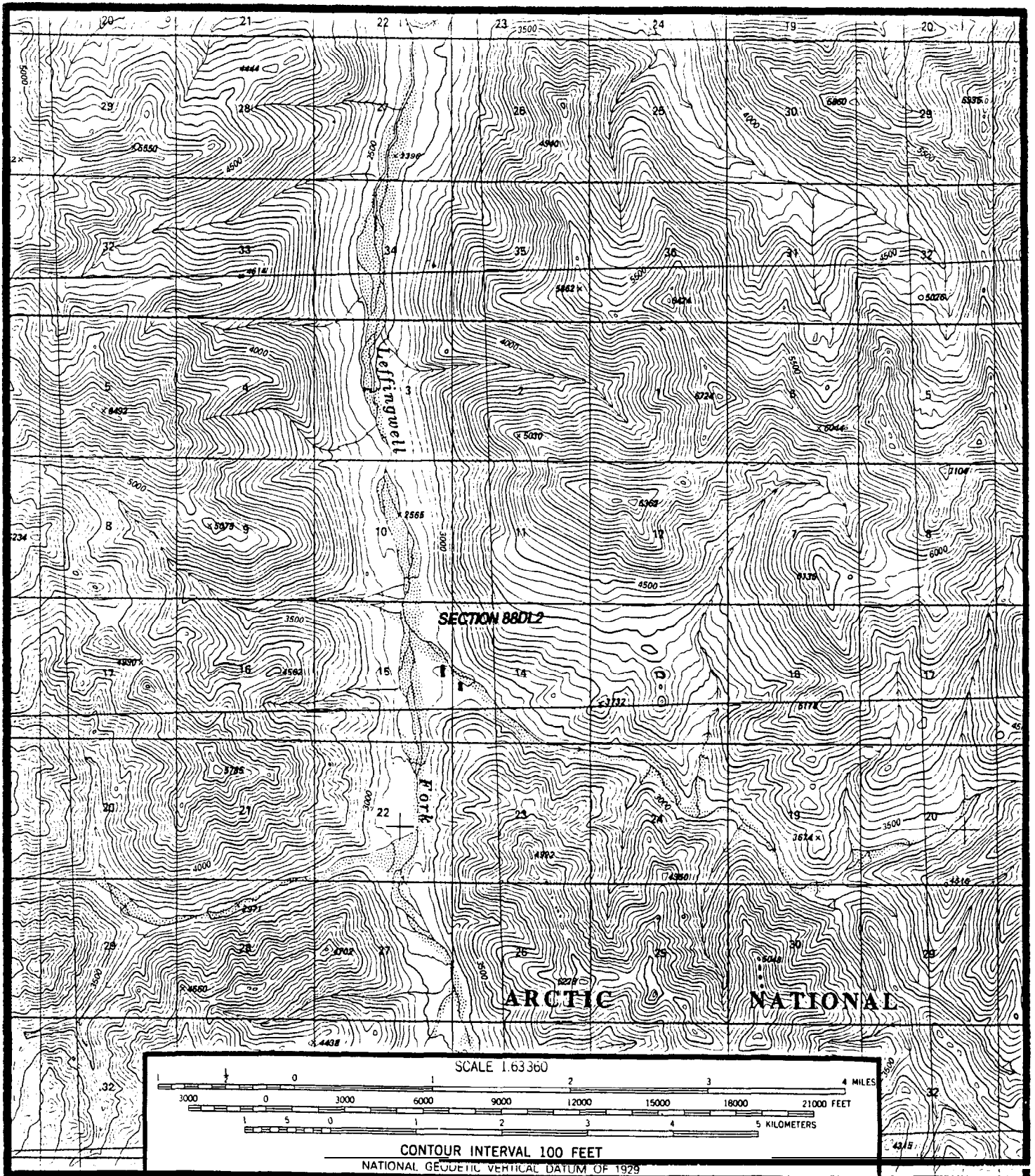
ACKNOWLEDGMENTS

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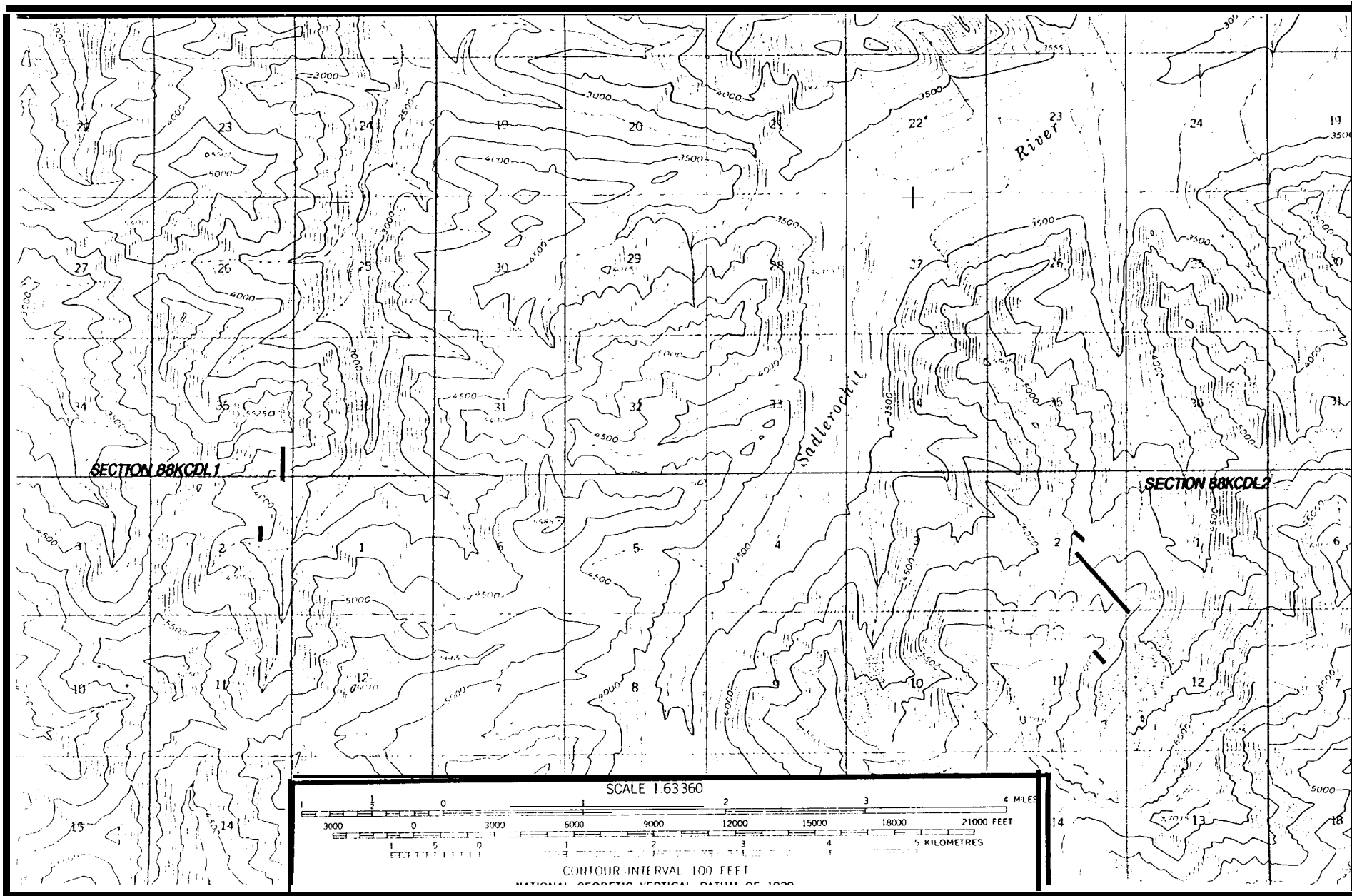
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Topographic map showing the location of section 88DL2 in the Romanzof Mountains, north of Bathtub Ridge (Demarcation Point Quadrangle).



Topographic map showing the location of sections 88KCDL1 and 88KCDL2 in the central Franklin Mountains, west of Lake Peters (Mt. Michelson Quadrangle).

SECTION KEY

Ripple-drift
cross lamination



Planar cross
stratification



Trough cross
stratification



Mutually erosive
channels



Lateral accretion
surfaces



Vertical burrows



Plant fragments



Log imprints



Rootlets



Echinoderm
fragments



Brachiopod
fragments



Key to symbols used in measured sections.

GROUP

FORMATION

THICKNESS (Meters)

GRAPHIC COLUMN

LITHOLOGIC

DESCRIPTION

CLAY SAND CONG
SILT
Fn Md Cs Gf Ppb Cob

GROUP

KAYAK SHALE

70

Covered interval.

60

Salt and pepper colored, thin-to medium-bedded, litharenite, dark gray siltstone, and minor gray-black shale. Sandstones are locally channelized and some contain lateral accretion surfaces.

50

ENDICOTT

CONGLOMERATE

40

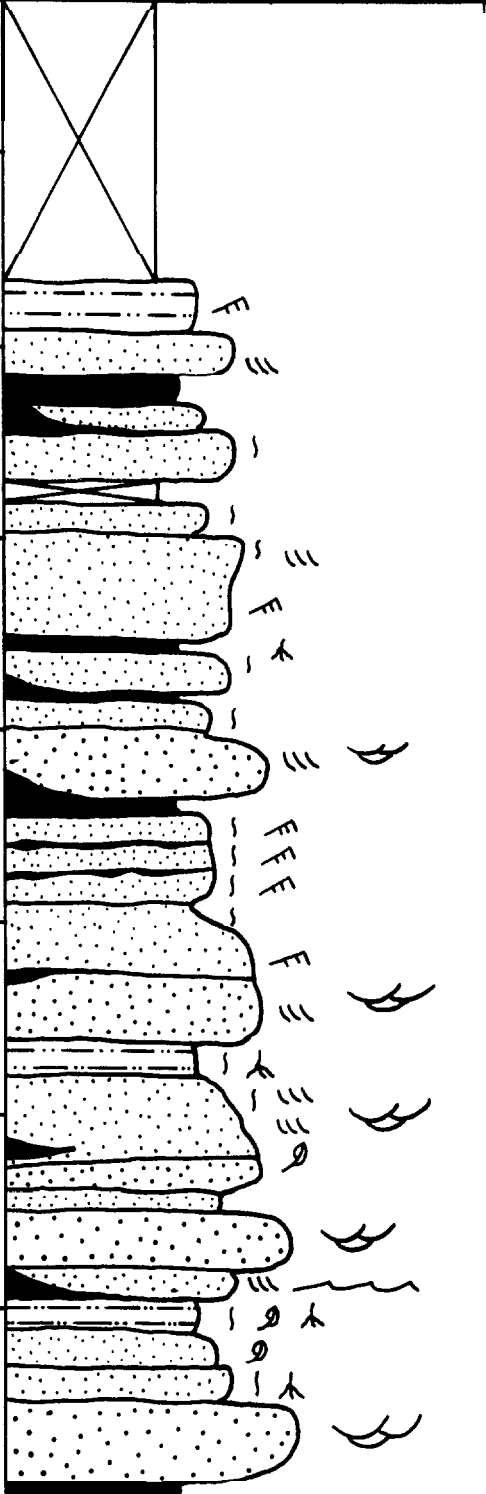
KEKIKTUK

30

Tan-brown, polymictic, grain-supported, conglomerate and litharenite, with minor gray-black siltstone and silty shale.

20

10



GROUP

FORMATION

THICKNESS (Meters)

GRAPHIC COLUMN

LITHOLOGIC

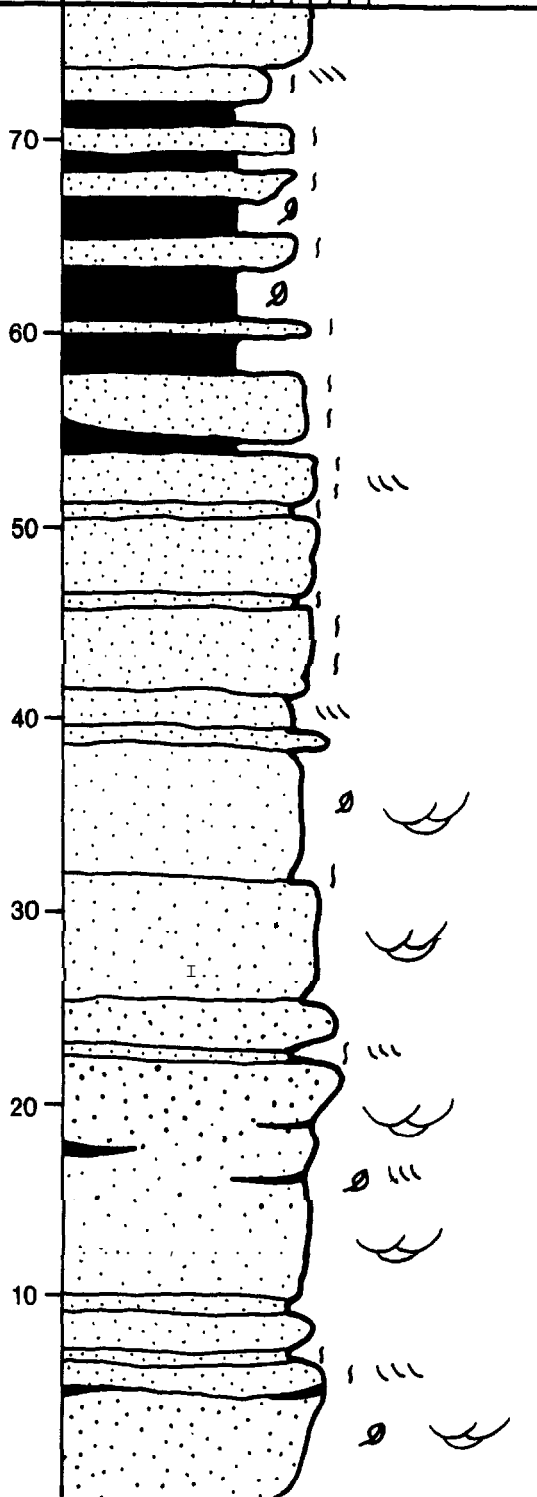
DESCRIPTION

CLAY SAND CONG
 SILT
 Fr Mn Co Gf Pcb Cob

GROUP

CONGLOMERATE

KEKIKTUK



Black, organic-rich sandy and silty shale, with irregularly interbedded, massive, bioturbated litharenite

Orange-brown, medium-bedded, massive, bioturbated, polymictic conglomerate and litharenite. Vertical burrows visible on most bedding surfaces and are surrounded by iron-oxide stains. Megaripples visible on many bedding surfaces.

Tan-brown colored, polymictic, grain-supported conglomerate and litharenite, with minor gray-black siltstone and silty shale.

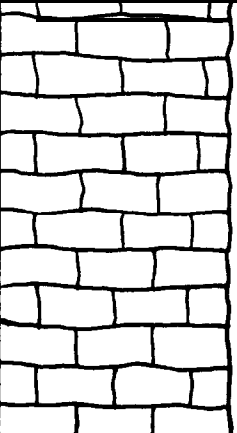
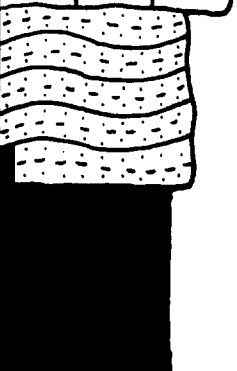
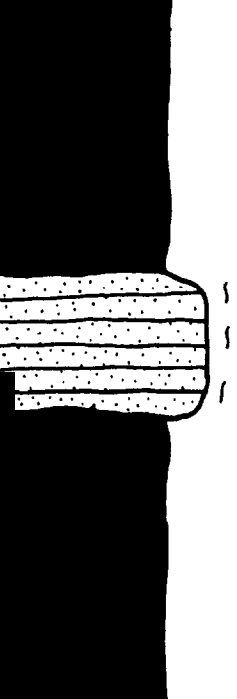
ENDICOTT

FRANKLIN **MOUNTAINS**

GROUP	FORMATION	THICKNESS (Meters)	GRAPHIC COLUMN	LITHOLOGIC DESCRIPTION
ENDICOTT	KAYAK	80 - 100	<p>CLAY SILT Fm Md Cs Gm Peb Cob</p> <p>SAND CONG</p>	<p>Black, silty shale</p> <p>Orange-brown, massive, bioturbated, litharenite.</p>
	SHALE	100 - 150		<p>Gray-black, siltstone, appears bioturbated (inferred from talus) .</p> <p>Tan-brown-to orange-brown, polymictic gonglomerate to litharenite, with minor silty, fine-grained sandstone.</p>

MISSISSIPPIAN ENDICOTT GROUP

FRANKLIN MOUNTAINS

GROUP	FORMATION	THICKNESS (Meters)	GRAPHIC COLUMN	LITHOLOGIC DESCRIPTION
LISBURNE GROUP	ALAPAH LIMESTONE	210		
GROUP	SHALE	200		<p>Calcareous siltstone</p> <p>Black, fissile, silty shale.</p>
ENDICOTT	KAYAK	170		<p>Dark red-brown, fine-grained sandstone. Unit contains small-scale overturned folds near top.</p>

GROUP	FORMATION	THICKNESS (Meters)	GRAPHIC COLUMN		LITHOLOGIC DESCRIPTION
			CLAY	SAND CONG	
ENDICOTT	KEKIKTUK	60			Gray-black, faintly laminated, arenaceous siltstone.
		50			Dark gray, thin-to medium-bedded, lithic wacke. In the middle of this interval beds are wavy and overlain by mud.
ENDICOTT	KEKIKTUK	40			Purple-gray, medium-bedded, sub-litharenite. Trough cross-stratification in sets 15-30 cm thick.
		30			Light gray, medium-bedded, litharenite. Bedding is disrupted and fractured, with some fracture surfaces containing slickenside striae. Numerous quartz veins cut through the zone oblique to bedding.
ENDICOTT	KEKIKTUK	20			Light gray, medium-bedded, litharenite. Trough cross-stratification in sets 15-20 cm thick.
		10			Light Gray, polymictic, matrix supported conglomerate and litharenite, arranged in 10-15 cm thick fining upward cycles; top of interval overlain by black, organic-rich, bioturbated silty shale.
	PRE-MISS. BASEMENT				Covered interval.