

SURFACE WATER

Runoff - Runoff is that portion of precipitation that leaves an area as streamflow. Average annual runoff in the Copper River basin ranges from less than 5 in. in the lowland areas to more than 54 in. in the high mountain areas (fig. 9). Annual runoff for the entire basin above Chitina averages 24.8 in., or 27,680 ft³/in. The runoff values and patterns shown on the map are based on measured and estimated streamflow and on other factors such as estimated evapotranspiration and area-averaged runoff calculations. The map is intended only to delineate the general area distribution of runoff from the basin and should not be used to estimate the flow of any specific stream.

Stream Types - The streams draining the Copper River basin can be classified into two general types: nonglacial streams, which drain lowland and low-altitude mountain areas, and glacial streams, which drain the high-altitude mountains.

Streams in the Copper River Lowland areas (fig. 2), have a relatively low gradient and derive their flow from snowmelt and rain. Streamflow is low from September through March, but with the increased solar radiation and warmer temperatures of April and May, flow reaches a peak. This peak results mainly from melting of snow and channel ice. A general recession in flow then takes place during June, July, and August. About 70 percent of the total annual flow occurs during the open-water period, May through September. Squirrel Creek (fig. 9) northwest of Tonsina is an example of a lowland stream.

Low-altitude mountain streams are nonglacial, as their drainage basins lie along the mountain flanks at altitudes too low for glaciers to exist. The Little Tonsina River (fig. 9) typifies such streams. Flow of the Little Tonsina increases due to snowmelt from late August through June and then declines during July, August, and September. About 80 percent of the total annual flow takes place between May and September, the open-water period.

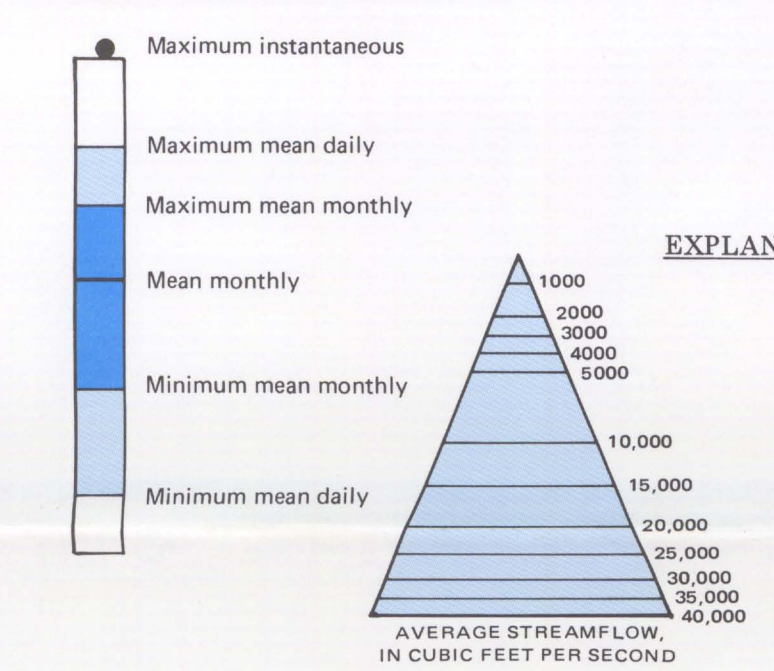
High-altitude mountain streams (glacial) exhibit the greatest seasonal variability of flow. About 88 percent of their total annual flow takes place during the approximately 5-month open-water period (May through September). High flows usually occur in July (fig. 9) when the highest seasonal temperatures cause maximum melting of glacier ice and snow.

The production of meltwater from a glaciated basin appears to be closely related to annual precipitation. This relation can be seen by comparing the cumulative departure of average annual flow of glaciated Tonsina River (fig. 10) to the departure curve of precipitation for the same period (fig. 4).

In the use and management of water, seasonal streamflow variation is of more concern than long-term variation. Streams throughout the basin experience high flows during the spring and summer and low flows during fall and winter. The seasonal and annual variability of streamflow is illustrated in figures 8a-8i. In general, streamflow in the basin shows the greatest variability in mean annual discharge for nonglacial streams and the least in streams affected by glacial meltwater. Glaciers tend to reduce the variation in annual discharge because they release water from ice and firm storage in dry, warm years and store water as snow and firm during cool, wet years. However, during the relatively dry years of 1969-70 even glacial streams experienced the lowest annual average discharge in the period of record (figs. 8c, 8e, 8h, and 8i).

A typical hydrograph of nonglacial streams, such as the Gulikana River (fig. 11), shows sharp May rises during the spring snowmelt, a general recession during the summer months, and a slight increase in streamflow during the early fall rainy period. In contrast, high flow on glacial streams, such as the Tonsina River, coincides with the peak melting of snow and ice in June, July, and August (fig. 11). Rainfall during these same months may produce even higher discharge when the rivers are already high from glacial runoff.

EXPLANATION FOR FIGURES 8a-8i



EXPLANATION FOR FIGURE 8

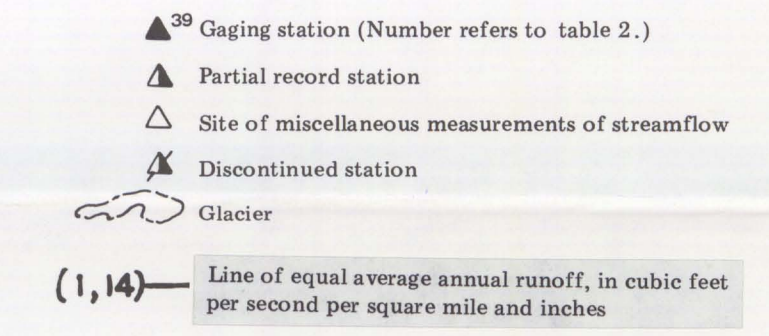
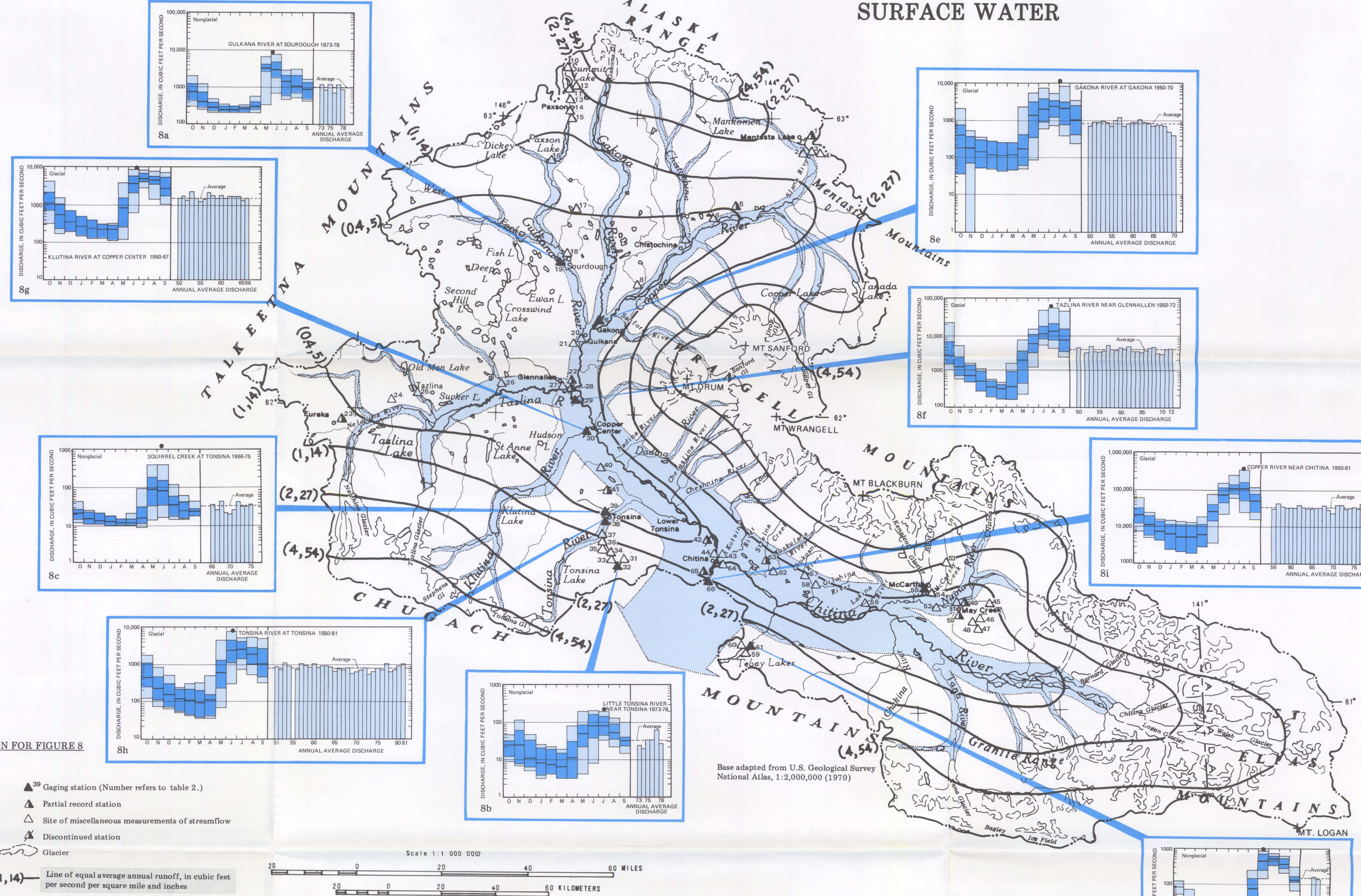


FIGURE 8. - SURFACE-WATER DATA-COLLECTION SITES, SUMMARY OF STREAMGAGING STATION RECORDS, AND BASIN AVERAGE ANNUAL RUNOFF.

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Streamflow - Total basin outflow has been measured continuously at a streamgaging station on the Copper River near Chitina since 1905. The average flow of the river at this point is 37,680 ft³/s, or 27,680,000 acre-ft/yr. It is estimated (fig. 8) that about three-fourths of this discharge is derived from the high mountain snowfields and glaciers. The average flow of the Chitina River, the Copper River's largest tributary, is estimated to be 20,000 ft³/s. Thus the Chitina provides slightly more than half the total basin runoff even though it drains only 38 percent (7,970 mi²) of the area of the Copper River basin. This high runoff is due to the Chitina River basin's relatively high altitude and consequent high precipitation.

Streamflow data have been collected at 10 continuous-recording gaging stations, 10 partial-record stations and 46 miscellaneous measurement sites in the Copper River basin (fig. 8). Information regarding these stations and sites - such as period of record, drainage area, and streamflow characteristics - is given on table 2.

The mean annual discharge of ungaged streams in the basin can be estimated using a regression analysis equation developed by Scully and Freestrey (1972). The equation is based on data from nearby gaged streams and values for basin characteristics such as drainage area and precipitation. Scully and Freestrey concluded that estimates of mean annual discharge of glacial streams and large nonglacial streams (drainage area greater than 100 mi²) are reliable. However, similarly derived estimates for small nonglacial streams are not reliable.

Flow Duration - Flow-duration curves are commonly used as a means of expressing the variability of streamflow. The curves show frequency distribution of average daily flows for the period of record and the percent of time any flow is equaled or exceeded. The variability of daily flows is shown for glacial (fig. 12) and nonglacial (fig. 13) streams. Various factors affect flow duration, including melting of snow and glacier ice, glacial outburst flooding, climate, geology, and ground-water inflow.

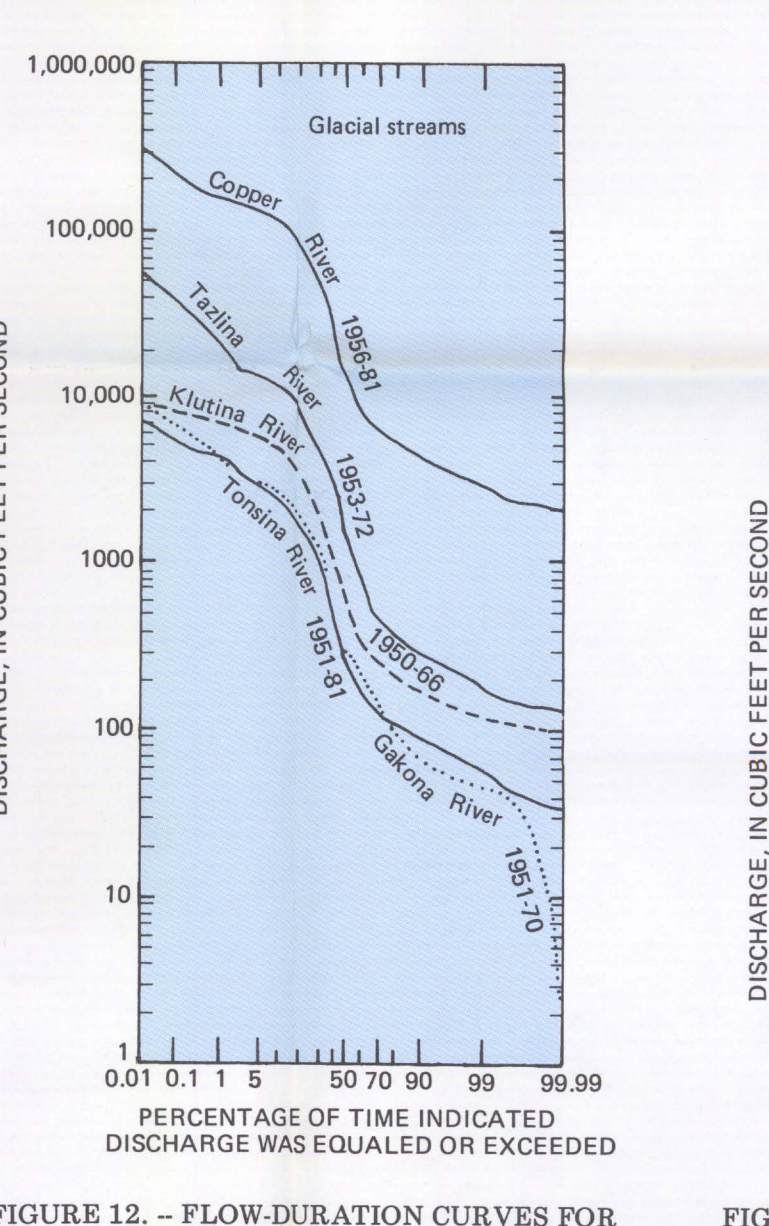


FIGURE 12. - FLOW-DURATION CURVES FOR FIVE GLACIAL STREAMS.

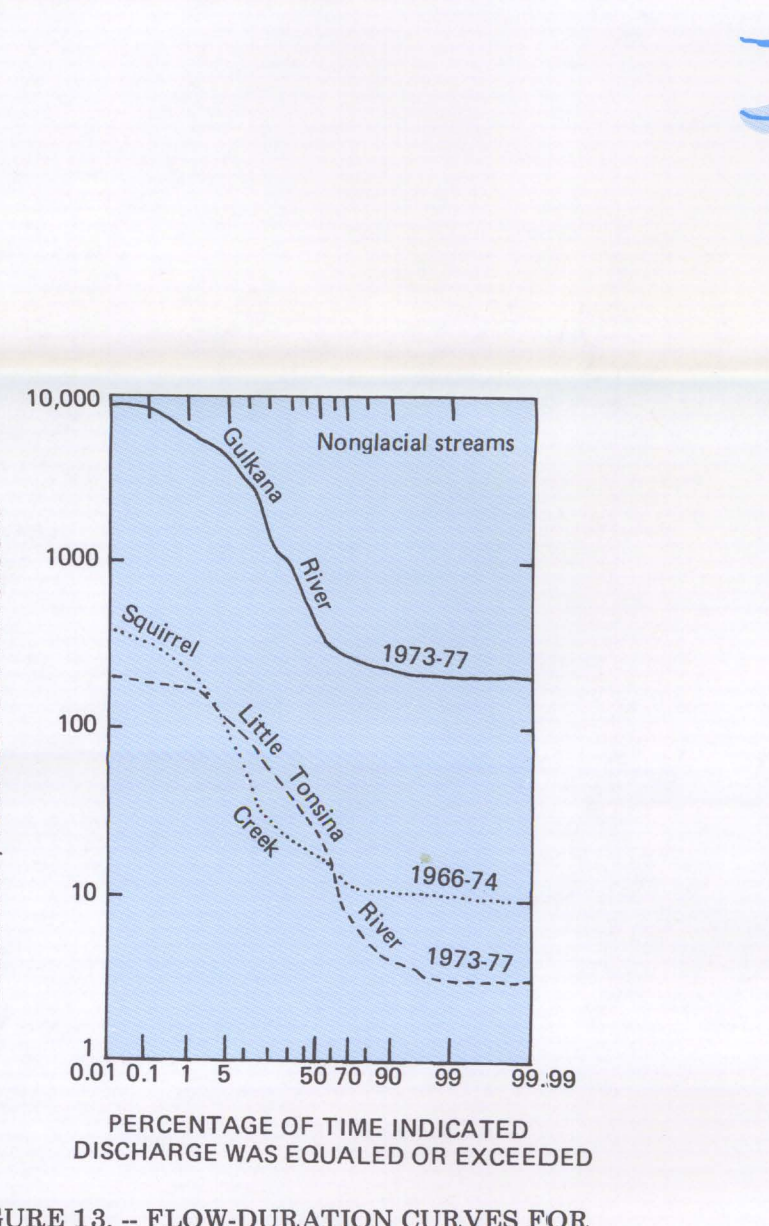


FIGURE 13. - FLOW-DURATION CURVES FOR THREE NONGLACIAL STREAMS.

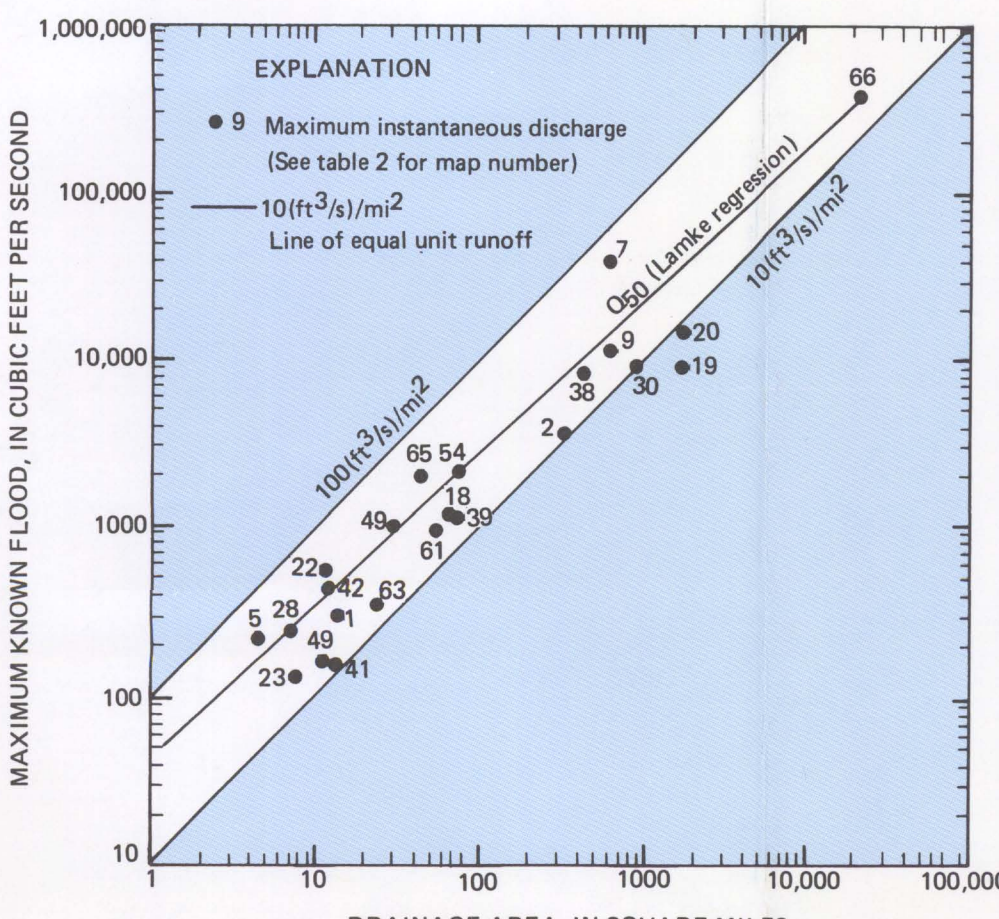


FIGURE 14. - RELATION BETWEEN MAXIMUM DISCHARGE AND DRAINAGE AREA.

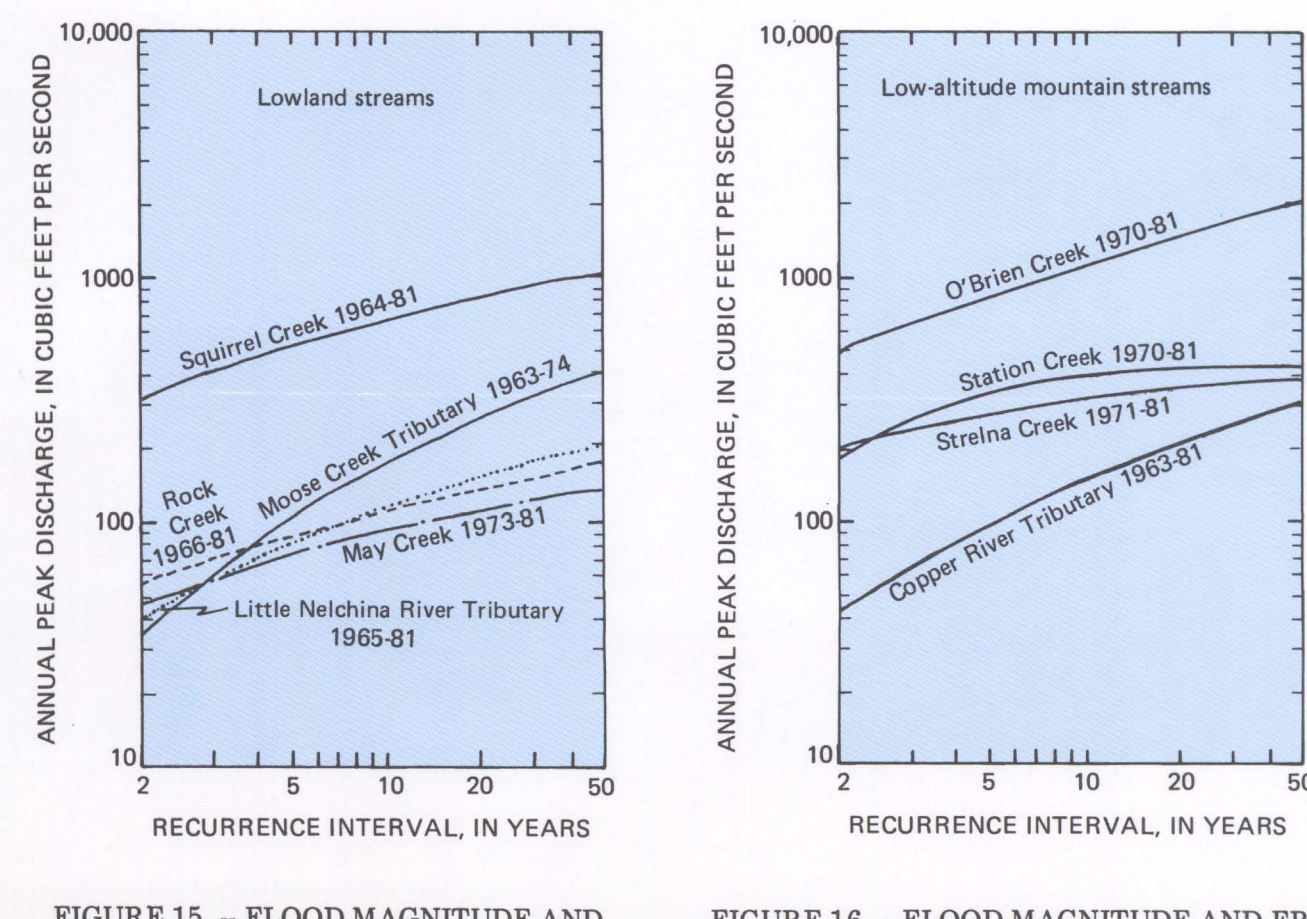


FIGURE 15. - FLOOD MAGNITUDE AND FREQUENCY FOR SELECTED LOWLAND STREAMS.

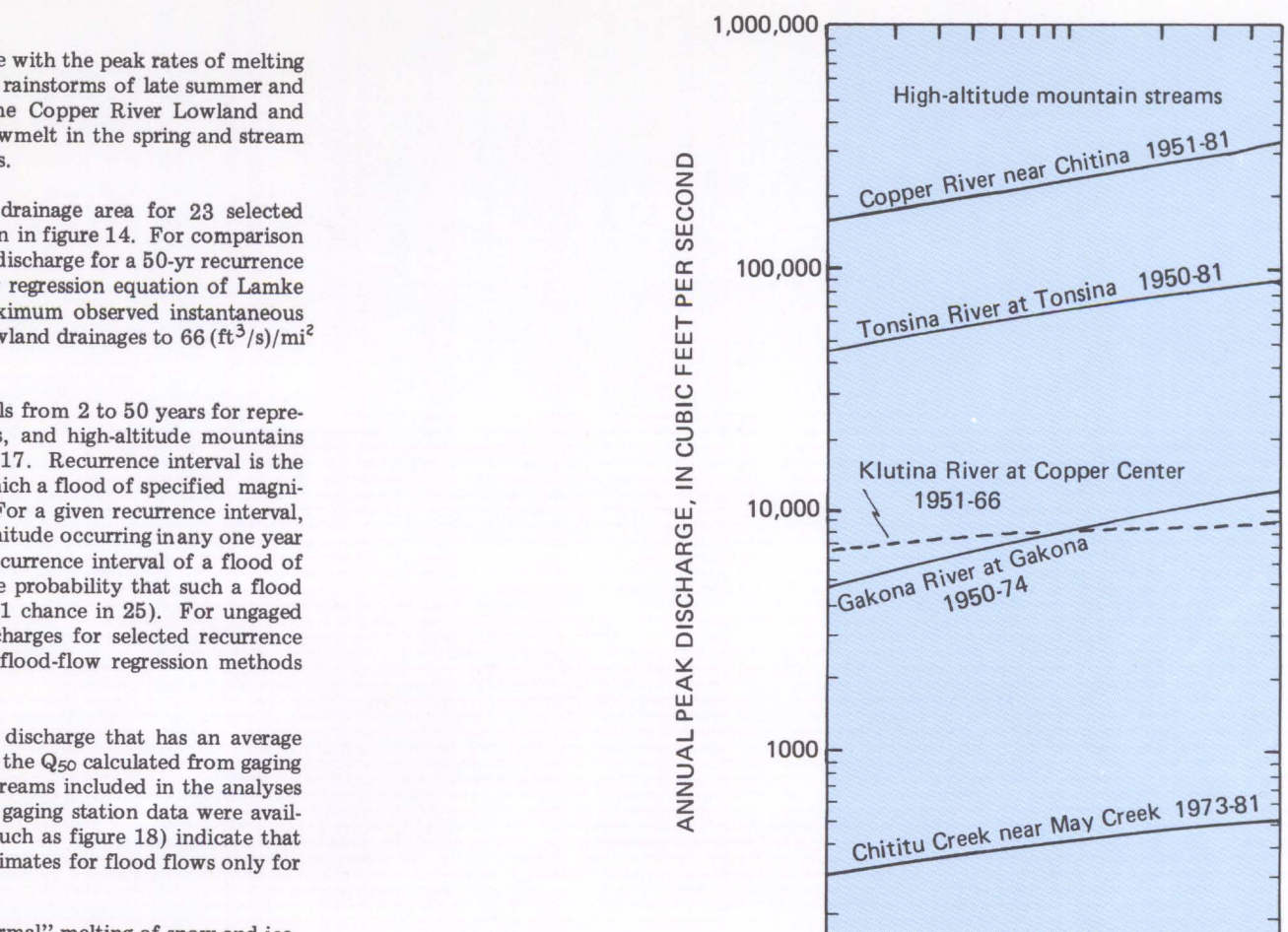


FIGURE 16. - FLOOD MAGNITUDE AND FREQUENCY FOR SELECTED LOW-ALTITUDE MOUNTAIN STREAMS.

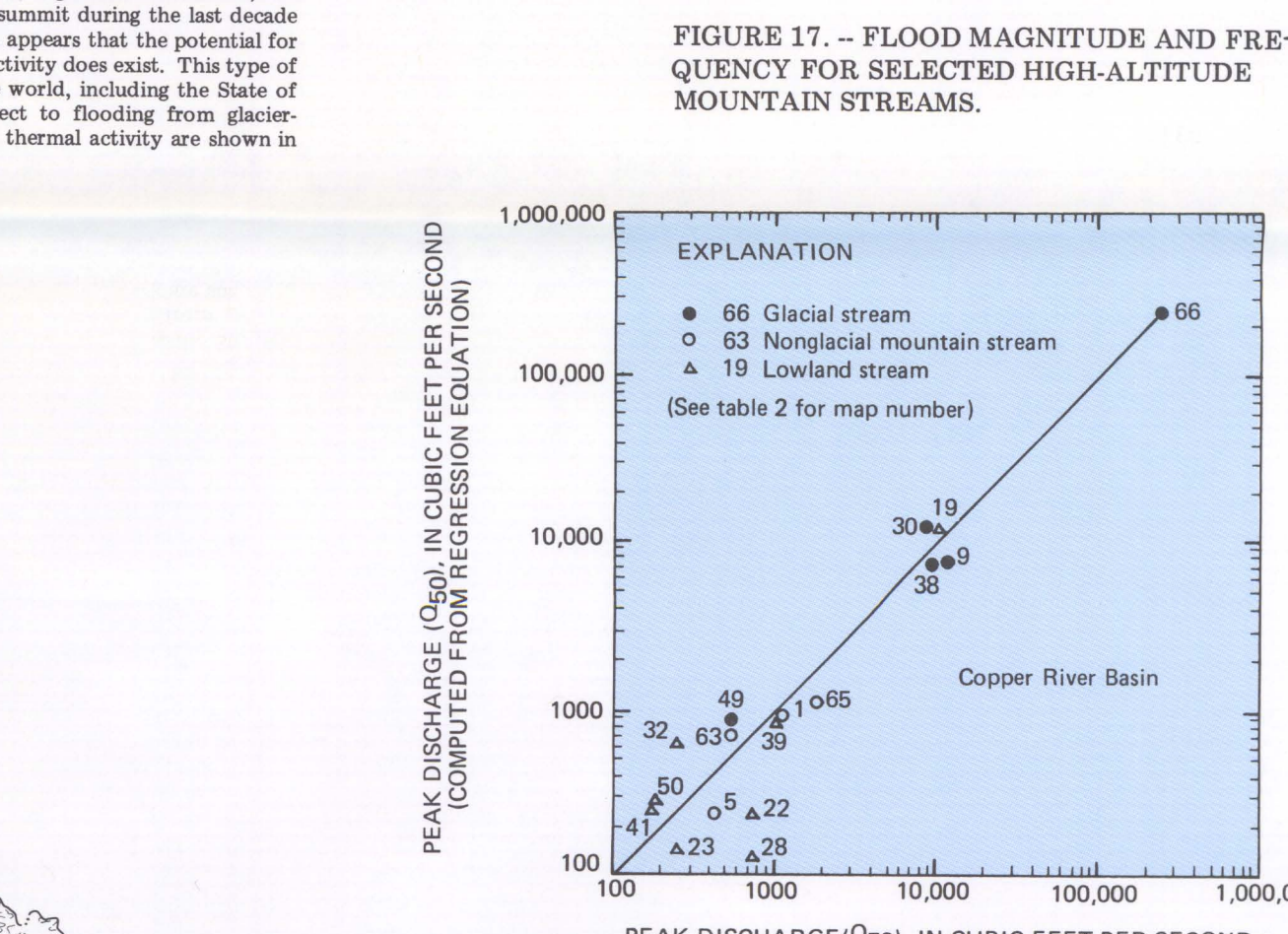


FIGURE 17. - FLOOD MAGNITUDE AND FREQUENCY FOR SELECTED HIGH-ALTITUDE MOUNTAIN STREAMS.

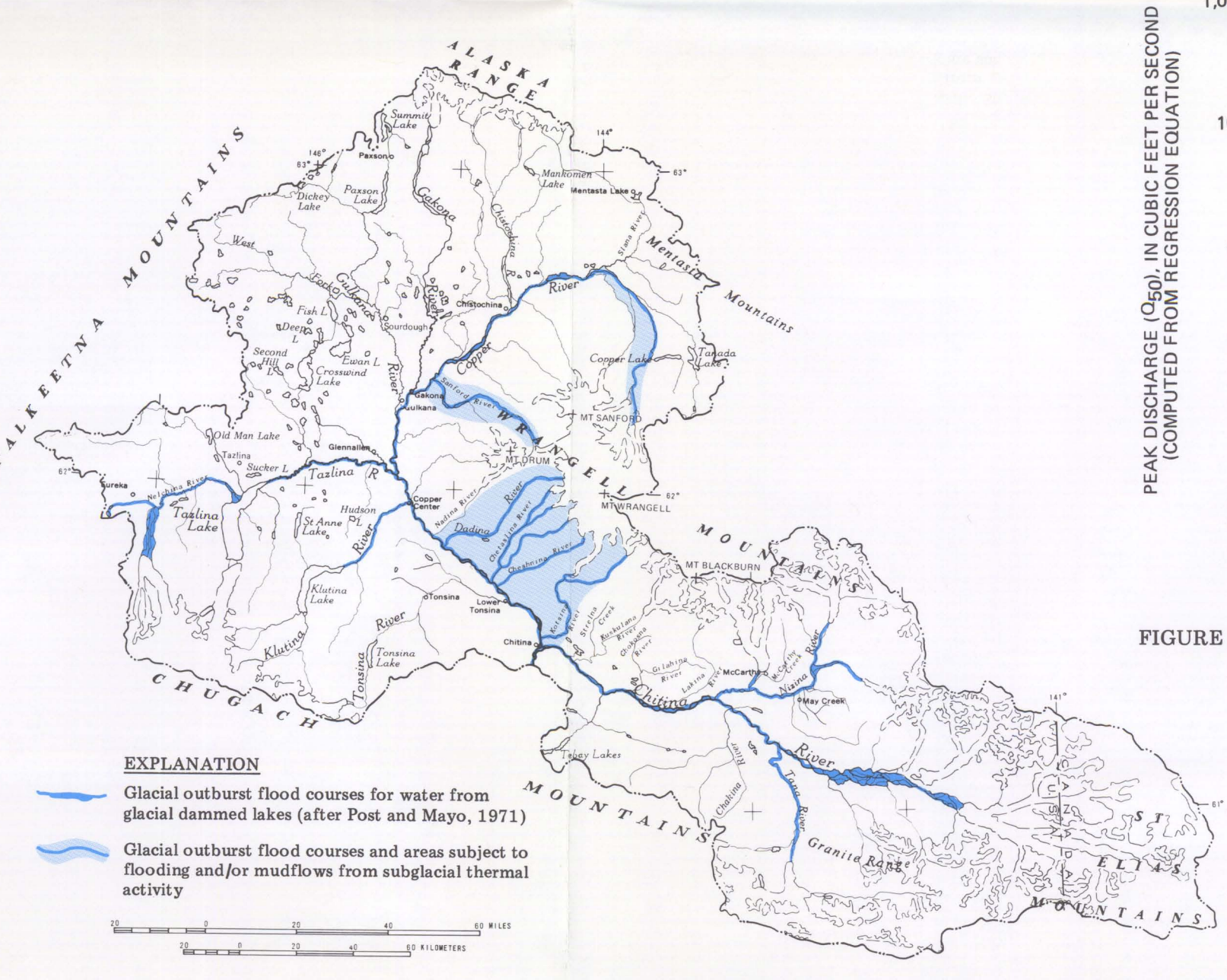


FIGURE 19. - AREAS SUBJECT TO FLOODING FROM GLACIAL OUTBURSTS.

EXPLANATION

Glacial outburst flood courses for water from glacial dammed lakes (after Post and Mayo, 1971)

Glacial outburst flood courses and areas subject to flooding and/or mudflows from subglacial thermal activity

Low Flow - Flow in lowland streams is at its minimum during periods of ice cover in late winter (March or April) or during periods of low precipitation in July, August, or September. However, minimum flow in both low-altitude and high-altitude mountain streams occurs during periods of late winter ice cover (March or April). Minimum measured flows of representative streams are listed in table 2. Methods given in Riggs (1972) were used to estimate low-flow characteristics at selected gaging stations and to calculate low-flow characteristics at partial-record stations and miscellaneous measurement sites (table 3). The low-flow analyses were made on a water year basis (October 1 through September 30). Low-flow characteristics for gaging stations with less than 10 years of record and for partial-record stations were estimated by correlation with continuous-recording gaging stations for which 10 or more years of record were available.

TABLE 3. - ESTIMATES OF MINIMUM DISCHARGE CHARACTERISTICS

Map No. (fig. 8)	Station No.	Station name	Drainage area (mi ²)	Annual flow (ft ³ /s) in cubic feet per second					
				7 consecutive days	10 consecutive days	30 consecutive days	60 consecutive days	90 consecutive days	
9	15202000	Gakona River	620	82	34	27	82	45	39
18	15202170	Squirrel Creek	68.0	0.2	0.2	0.2	0.2	0.2	0.2
19	15202200	Gulikana River	1,970	245	199	187	247	200	188
22	15201000	Dry Creek	11.4	NF	NF	NF	NF	NF	NF
23	15201100	Little Nelchina River Tributary	7.81	NF	NF	NF	NF	NF	NF
24	15202000	Chitina River	356	28	24	22	28	24	22
28	15201900	Moose Creek Tributary	7.11	NF	NF	NF	NF	NF	NF
29	15202000	Tazlina River	2,670	266	197	179	277	206	188
30	15202000	Chitina River	880	185	120	110	185	116	112
32	15207800	Little Tonsina River	22.7	4.4	2.8	2.6	4.7	2.9	2.7
38	15202000	Tonsina River	420	84	57	50	86	58	52
39	15202000	Squirrel Creek	70.5	11	9.3	9.0	11	9.7	9.5
41	15202000	Rock Creek	14.1	NF	NF	NF	NF	NF	NF
44	15201900	Copper River above Chitina River	12,630	2,800	1,600	1,600	2,800	1,600	1,600
50	15201900	May Creek	10.4	NF	NF	NF	NF	NF	NF
64	15201900	O'Brien Creek at mouth	7,970	1,800	1,200	1,200	1,800	1,200	1,200
65	15211900	O'Brien Creek	44.8	3.0	1.2	1.0	3.0	1.4	1.2
66	15212000	Copper River at Chitina	20,600	4,600	3,000	2,700	4,600	3,000	2,700

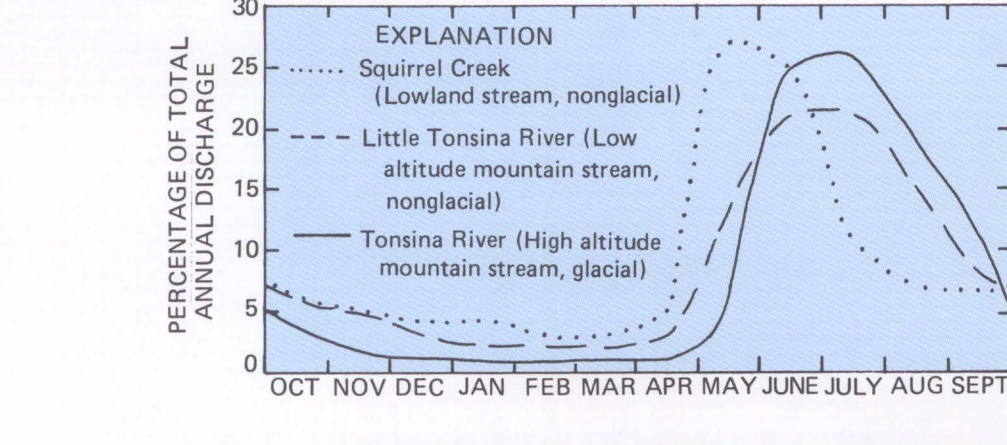


FIGURE 9. - MEAN MONTHLY DISTRIBUTION OF DISCHARGE OF THREE STREAMS.

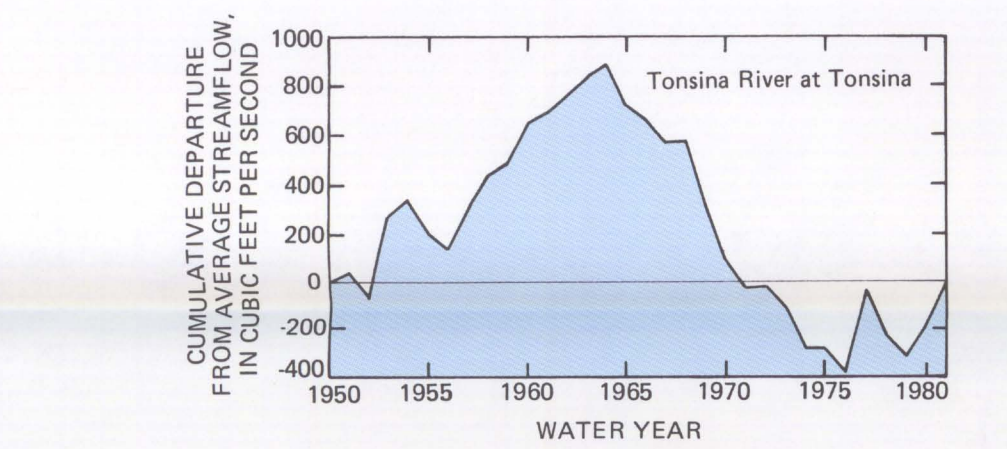


FIGURE 10. - CUMULATIVE DEPARTURE FROM AVERAGE STREAMFLOW, TONSINA RIVER AT TONSINA.

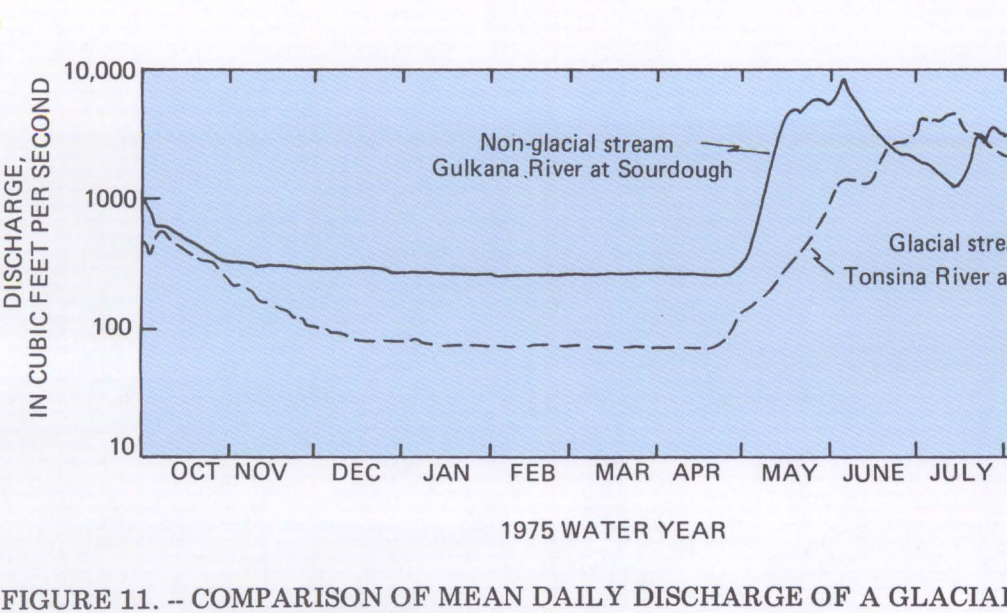


FIGURE 11. - COMPARISON OF MEAN DAILY DISCHARGE OF A GLACIAL VS. NON-GLACIAL STREAM.

WATER RESOURCES OF THE COPPER RIVER BASIN, ALASKA

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