# U.S. GEOLOGICAL SURVEY

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Earthquake Locations Determined by the Southern Alaska Seismograph Network for October 1971 through May 1989

by

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## CONTENTS

	Page
Introduction	. 3
Instrumentation	. 3
Data Processing	. 10
Velocity Models	
Traveltime Delay Models and Trial Focal Depths	13
Magnitude	
Analysis of Hypocentral Quality	. 17
Hypocenter Precision	18
Focal Depths	
Completeness of Catalog	19
Seismicity of Southern Alaska	
Availability of Data	
Acknowledgements	
References	
ILLUSTRATIONS	Dago
	Page
Figure 1 Map showing principal seismograph stations used in locating earthquakes	
2 Block diagram of the USGS telemetered seismograph system	
3 System response curves of typical USGS telemetered seismograph stations	
4 Histograph showing the number of earthquakes located per month	
5 Map showing earthquake epicenters with magnitudes greater than 4.0	
6 Map showing earthquake epicenters with depths equal to and below 30 km	
7 Map showing earthquakes epicenters with depths shallower than 30 km	
8 Map showing location of cross sections	
9 Cross sections showing depth distribution of earthquake hypocenters	
10 Error ellipsoid relationships	. 54
TABLES	
	Page
Table 1 Station parameters	4
2 Alaska velocity models	
3 Geographical boundaries, starting depths, velocity models, and delay models	
4 P-phase and S-phase traveltime delays	15
APPENDICES	
	Page
Appendix A Magnitude 4 and larger earthquakes	_
B List of previously published catalogs	
C Criteria for processing earthquakes	37

#### INTRODUCTION

The U.S. Geological Survey (USGS) has operated a regional network of seismographs in southern Alaska since 1971. The principal purpose of this network has been to record seismic data to be used to precisely locate earthquakes in the seismic zones of southern Alaska, delineate seismically active faults, assess seismic risks, document potential premonitory earthquake phenomena, investigate current tectonic deformation, and study the structure and physical properties of the crust and upper mantle. A task fundamental to all of these goals is the routine cataloging of parameters for earthquakes located within and adjacent to the seismograph network.

The initial network of 10 stations, 7 around Cook Inlet and 3 near Valdez, was installed in 1971. In subsequent summers additions or modifications to the network were made. By the fall of 1973, 26 stations extended from western Cook Inlet to eastern Prince William Sound, and 4 stations were located to the east between Cordova and Yakutat. A year later 20 additional stations were installed. Thirteen of these were placed along the eastern Gulf of Alaska with support from the National Oceanic and Atmospheric Administration (NOAA) under the Outer Continental Shelf Environmental Assessment Program to investigate the seismicity of the outer continental shelf, a region of interest for oil exploration. Since then the region covered by the network has remained relatively fixed while efforts have been made to make the stations more reliable through improved electronic instrumentation and strengthened antenna systems. The majority of the stations installed since 1980 have been operated only temporarily (from one to several years) for special studies in various areas within the network. Due to reduced funding the network was trimmed substantially in the summer of 1985 with the closure of 15 stations, 13 of which were located in and around the Yakataga seismic gap. To further reduce costs, two telephone circuits were dropped and multiple radio relays were installed in their place. This economy reduced the reliability of these telemetry links. In addition, data collection from the areas around Cordova and Yakutat was compromised by the necessity of relying on triggered event recording using PC-based systems (Rogers, 1993) that were not fully developed and which proved to be less reliable than anticipated.

The principal means of recording throughout the time period of this catalog was 20-channel oscillographs on 16-mm film (Teledyne Geotech Develocorder, Model RF400 and 4000D). Initially one Develocorder was operated at the USGS headquarters in Anchorage, but recording was shifted to the National Oceanic and Atmospheric Administration (NOAA) Palmer Observatory (currently the Alaska Tsunami Warning Center) in 1972. The Develocorders were turned off at the end of May 1989, and since that time recording has been done in digital format at the Geophysical Institute of the University of Alaska in Fairbanks (GIUA). Thus, this catalog covers the entire period of film recording.

#### INSTRUMENTATION

The locations of seismograph stations of the USGS network that contributed to this catalog are shown in Figure 1 and listed in Table 1, along with stations operated by other institutions from which reading were obtained. Most of the USGS stations have only single, vertical-component sensors, but horizontal components seismometers are also operated at a few selected sites (see Table 1).

TABLE 1. List of seismograph stations used to determine hypocenter parameters reported in this catalog. INST - principal operating institution (USGS - U.S. Geological Survey; GIUA - Geophysical Institute, University of Alaska; NOAA - Alaska Tsumani Warning Center, Palmer; EMRC - Energy Mines and Resources, Canada [now Pacific Geoscience Center]). Open date left blank (or = 0) if unknown; closed date left blank if operation continued after May 1989.

CODE	SITE NAME	N LATITUDE DEG MIN	W LONGITUDE DEG MIN	ELEV M	INST	OPEN	CLOSED
ARF	AUKE BAY ADAK AUGUSTINE ISLAND AGASSIZ LAKES AUGUSTINE ISLAND ALCAN ANVIL MOUNTAIN ANVIL MOUNTAIN AUGUSTINE ISLAND AUGUSTINE ISLAND AUGUSTINE ISLAND AUGUSTINE ISLAND AUGUSTINE EAST AUGUSTINE FLOW AUGUSTINE DOME H AUGUSTINE ISLAND AUGUSTINE HOUND AUGUSTINE PINICLE	58 22.88	134 38.60	3	USGS	80/ 7/21	
ADK	ADAK	51 53.02	176 41.07	116	NOAA	66/ 1/ 0	
AG1	AUGUSTINE ISLAND	59 22.80	153 25.20	580	GIUA	71/8/0	76/ 1/ 0
AGA	AGASSIZ LAKES	60 9.25	141 2.00	1024	USGS	83/10/ 5	88/ 8/26
AGI	AUGUSTINE ISLAND	59 22.80	153 25.20	580	GIUA	71/ 8/ 0	76/ 1/ 0
ALC	ALCAN	62 37.35	141 0.50	582	USGS	79/10/25	82/ 9/21
ANM	ANVIL MOUNTAIN	64 34.60	165 22.28	323	GIUA	76/ 9/ 0	
VVA	ANVIL MOUNTAIN	64 33.60	165 22.28	323	GIUA	76/ 9/ 0	
AU1	AUGUSTINE ISLAND	59 22.39	153 25.23	494	GIUA	71/ 8/ 0	76/ 1/ 0
AU2	AUGUSTINE ISLAND	59 22.21	153 22.69	195	GIUA		76/ 1/ 0
AU4	AUGUSTINE ISLAND	59 20.16	153 30.77	18	GIUA	76/ 1/ 0	
AU5	AUGUSTINE ISLAND	59 23.19	153 27.35	152	GIUA	00/10/ 1	76/ 1/ 0
AUE	AUGUSTINE EAST	59 21.54	153 22.33	1/2	GIUA	88/10/ 1	00/0/06
AUF	AUGUSTINE FLOW	59 23.27	153 27,45	100	GIUA	77/ 7/28 92/ 8/26	80/ 8/25
HUM	AUGUSTINE DOME H	59 21.63	153 20.01	202	CIUN	78/ 1/ 1	
AUT	AUGUSTINE ISLAND	59 20.11	153 25.00	203	GIVA	76/10/17	78/4/6
אטא	MICHEMENT INVA TION	50 20.03	153 25.02	360	GIUA	80/10/29	70/ 4/ 0
MIIA	AUGUSTINE MAININ	59 22.95	153 20.07	106	GIUA	75/ 9/ 0	80/10/29
AUL	AUGUSTINE PINICLE	59 21 74	153 25.23	1033	GIUA	88/10/ 1	00/20/25
AUW	AUGUSTINE WEST	59 22.02	153 28.25	320	GIUA	91/ 7/27	
BAL	AUGUSTINE PINICLE AUGUSTINE WEST BALDY BEAVER CREEK U3 BANCAS POINT BANCAS POINT BELUGA BIG MOUNTAIN BIG MOUNTAIN BLACK RAPIDS BURWASH LANDING BREMNER RIVER	61 2.12	142 20.83	1265	USGS	73/ 8/24	
BC3	BEAVER CREEK U3	63 3.95	141 46.96	848			
BCP	BANCAS POINT	59 57.20	139 38.10		USGS	79/ 9/ 4	
BCS	BANCAS POINT	59 56.90	139 38.10 139 37.00	10	USGS	76/ 6/25	79/ 9/ 4
BGA	BELUGA	61 13.34	150 57.95	20	USGS	75/10/ 8	76/ 3/27
BGM	BIG MOUNTAIN	59 23.56	155 13.76		USGS	78/ 9/ 8	
BIG	BIG MOUNTAIN	59 23.34	155 13.02	567	USGS	72/ 7/31	78/ 9/ 8
BLR	BLACK RAPIDS	63 30.10	145 50.70	810	NOAA	65/ 3/ 0	85/8/8
BLY	BURWASH LANDING	61 22.35	139 1.56	799	USGS	74/ 7/22	78/ 6/ 1
BMR	BREMNER RIVER	60 58.09	144 36.18	823	USGS	79/ 8/19	85/ 7/16
BRW	BARROW CAETANI RIVER CLEAR CREEK BUTTE CAPE DOUGLAS CAPE DOUGLAS	71 18.20	156 44.90	0	GIUA	88/10/ 1	n. / n /nn
CAE	CAETANI RIVER	60 5.05	140 59.33	716	USGS	83/ 7/25	84/ 8/27
CCB	CLEAR CREEK BUTTE	64 38.80	147 48.33	219	GIUA	83/ 9/16	01/0/16
CDA	CAPE DOUGLAS	58 57.32	153 31.//	500	GIUA	78/ 1/ 1 81/ 8/17	81/ 8/16
CDD	CANDLE	66 6.84	161 39.40	312	GIUA	78/12/16	
	COLLEGE FIORD	61 10.96		312	USGS	74/ 7/31	
	CONGABUNA	61 4.14		160	USGS	75/10/ 8	76/ 3/27
	CAPPS GLACIER	61 18.46	152 0.40	1082	USGS	81/ 9/22	10/ 3/2/
	CHAIX HILLS	60 3.78	141 7.00	1067	USGS	74/ 9/ 4	85/ 8/25
	CHEKOK LAKE	59 57.58		732	USGS	72/ 7/29	78/ 9/ 9
	CHINA POOT	59 31.55	151 14.16	564	USGS	83/ 7/ 1	
	COLLEGE OUTPOST	64 54.00		320	USGS	64/ 1/ 0	
	CRATER PEAK	61 16.02	152 9.33	1622	USGS	81/ 8/26	
	CIRQUE	60 45.40	143 8.35	1853	USGS	88/ 7/ 1	
	CHILDS GLACIER	60 39.66	144 51.30	678	USGS	84/ 7/22	85/ 8/10
		60 57.96	141 20.30	1490	USGS	79/ 8/28	
CUT	CHULITNA	62 24.28	150 16.17	168	GIUA	86/ 7/18	
		-					

<sup>\*</sup> Three-component site during at least part of the operation period

TABLE 1 (cont)

		N LA	TTUDE	W LO	NGITUDE	ELEV				
CODE	SITE NAME	DEG	MIN	DEG	MIN	M	INST	OPEN	CLC	SED
_ ~~ -							~			
CVA	CORDOVA CAPE YAKATAGA DONNELY DOME DRIFT RIVER DEZADEASH LAKE	60	32.84	145	44.73	120	USGS	71/ 8/31		
CYT	CAPE YAKATAGA	60	4.47	142	24.68	323	USGS	78/ 8/ 8	80/	9/22
DDM	DONNELY DOME	63	47.14	145	51.45	900	GIUA	90/ 8/10		
DFR	DRIFT RIVER	60	35.51	152	41.16	1090	USGS	88/ 8/15		
DLY	DEZADEASH LAKE	60	22,20	137	3.90	738	EMRC			
DMA	DEVIL MOUNTAIN	66	17.80	164	31.35	238	GIUA	77/ 8/ 0	82/	0/0
MMG	DELTA MICROWAVE	64	3.23	145	43.52	346	GIUA			8/19
DOT	DOT LAKE	63	38.92	144	3.75		GIUA		- ,	
חכם	DISCHANTMENT BAY	60	4 60	139	32.70	640	USGS		76/	6/ B
DOR	DISK ISLAND	60	30.12	147	38.81	1.5	USGS		76/	6/7
DWY	DAWSON CTTY	64	3 20	139	25.90	346	EMRC		76/	6/7
FRN	FRNESTINE	61	26.65	145	6.74		USGS	71/ 9/16	73/	8/29
493	COLLEGE OUTPOST	64	54 00	147	47.60			91/8/1		• /
FID	FIDALCO	60	43 73	146	35.79			74/10/ 7		
ETC	FIDE TOLAND	63	8 65	150	13 11	76		74/ 9/24	76/	5/4
FVII	EUDA AIRUM	66	23 96	145	13.11 13.90 58.27 58.70	137		88/10/1	. 07	0, 1
CVD	CARNER	63	50 22	149	58 27	590		87/10/20		
Chv	CDANITE DAY	60	25 03	147	58.27	495	USGS			
CUO	CLODYHOLE	61	46 33	140	55,45	1021	USGS			
GIO	GLUKINOLE CREEK	6.4	40.33	147		350	AAON			
GIT	GILMORE CREEK	64	30.30	147	56.08		GIUA		06/	3/10
GKC	GOLD KING CREEK	64	10.72	14/			USGS		867	3/10
*GLB	DEVIL MOUNTAIN DELTA MICROWAVE DOT LAKE DISCHANTMENT BAY DISK ISLAND DAWSON CITY ERNESTINE COLLEGE OUTPOST FIDALGO FIRE ISLAND FORT YUKON GARNER GRANITE BAY GLORYHOLE GILMORE CREEK GOLD KING CREEK GILAHINA BUTTE GLACIER ISLAND	9.7	70.31	143	48.63				011	9/17
GLC	GLACIER ISLAND	60	53.44	14/	4.38 5.65	420	USGS		84/	9/1/
GL1	GLACIER ISLAND GILMORE DOME GLENNALLEN GRANITE MOUNTAIN GUYOT HARDING LAKE	60	52.78	14/	5.65	429	USGS			
GLM	GILMORE DOME	64	59.24	14/	23.34 32.93	820	GIUA			
GLN	GLENNALLEN	62	6,65	145	32.93	45/	GIUA			
GMA	GRANITE MOUNTAIN	65	25.72	161	13.92		GIUA			
GYO	GUYOT	60	8.78	141	28.29		USGS			
HDA	HARDING LAKE	64	24.34	146	56.45		GIUA			
					30.10	611	USGS	7.4/10/ 3		
HMT	HAMILTON HOMER HARLEQUIN HURRICANE HAINES JUNCTION ILIAMNA ILIAMNA	60	20.19	144	15.64	620	USGS	77/ 8/16		
MOH	HOMER	59	39.50	151	38.60	198	GIUA	81/ 1/ 1		
НQИ	HARLEQUIN	59	27.10	138	52.62	372	USGS	74/10/ 1		
HUR	HURRICANE -	62	58.72	149	38.60	496	GIUA	77/ 2/14		
$\mathtt{HYT}$	HAINES JUNCTION	60	49.50	137	30.24	1416	EMRC	81/ 7/27		
ILI	ILIAMNA	60	4.81	152	57.57	823	USGS	87/ 9/15		
ILM	ILIAMNA	60	10.92	152	48.97	550	USGS	71/ 8/ 7	87/	9/16
ТТИ	ILIAMNA	60	10.92	132	40.91	330	0363	/1/ 6/ /	72/	9/30
	NIATHUOM MAIDHI		4.11		40.72	1380	NOAA			
INK	INUVIK		17.50	133	30.00	40	EMRC			
KAI	KAYAK ISLAND		55.61		24.98	311	USGS	82/ 8/ 3		
KDC	KODIAK	57	44.87	152	29.50	13	NOAA	78/ 9/20		
KEY	KLUANE LAKE	61	3.00	138	30.10	785	<b>EMRC</b>	78/ 8/26	81/	7/25
KLU	KLUTINA	61	29.57	145	55.21	1021	USGS	72/ 7/22		
KMP	KIMBALL PASS	61	30.78	145	1.09	1143	USGS	76/ 8/ 1	85/	7/19
	KNIGHT ISLAND		20.92	147	44.16	434	USGS	85/ 7/21		
	KNIK GLACIER		24.75			595	USGS	73/ 8/11		
	KOIDERN RIVER		58.20		24.50	686	EMRC	78/ 8/29	81/	4/1
	KOTZEBU		50.48		35.22	24	GIUA	76/ 9/ 8		0/0
	KANTISHNA HILLS		33.19		55.26	1172	GIUA	88/10/ 1	•	
	KATMAI		19.48		22.59	945	USGS	73/ 7/27	75/	8/24
	KAYAK ISLAND		52.10		31.39	375	USGS			8/2
	LOUIS BAY		27.93		38.66	490	USGS		•	
	LATOUCHE ISLAND	60			51.25	302	USGS	88/ 7/ 1		
				'						

CODE	SITE NAME	N LA'	TITUDE MIN	W LO	NGITUDE MIN	ELEV M	INST	OPEN	CLOSED
LVY	SITE NAME  LEVY  MOULD BAY  MCKINLEY PARK  MCNEIL RIVER  MIDDLETON ISLAND  MALASPINA GLACIER  MCNEIL RIVER  MARTIN RIVER  MOOSE CREEK  MOOSE PASS  MONTAGUE ISLAND  MONTAGUE ISLAND  NELCHINA  NORTH CRESENT  NENANA  NORTH GASLINE  NIKOLSKI  NINILCHIK  NIKISHKA  NIKOLSKI  NINILCHIK  NORTH RIVER  NUNITAK  OCEAN CAPE  OIL POINT  PAXSON  [see PIN]  POINT CAMPBELL	64	13.00	149	15.20	230	GIUA	72/ 7/ 0	89/ 9/19
MBC	MOULD BAY	76	14.50	119	21.60	15	EMRC	61/10/18	, -,
MCK	MCKINLEY PARK	63	43.94	148	56.10	618	GIUA	64/ 0/ 0	
MCN	MCNEIL RIVER	59	6.06	154	11.99	273	GIUA	76/10/15	81/ 8/22
MID	MIDDLETON ISLAND	59	25.67	146	20.34	37	NOAA	76/ 5/21	
MLS	MALASPINA GLACIER	59	46.00	140	9.00	30	USGS	74/ 9/ 4	81/ 9/ 5
MMN	MCNEIL RIVER	59	11.11	154	20.20	442	GIUA	81/ 8/22	
MRN	MARTIN RIVER	60	32.10	144	0.50	957	USGS	75/ 8/22	76/ 3/31
MSE	MOOSE CREEK	61	50.30	148	58.03	1318	USGS	84/ 9/11	85/ 9/30
MSP	MOUSE PASS	6U 6.0	29.35	149	21.03	7 Ø Å	0262	73/8/3	05/ 7/35
MTG	MONTAGUE ISLAND	59	59.71	147	30 02	737	11565	14/10/ 3 85/ 7/21	85/ //15
NC A	MELCHINA	61	59.27	146	19.02	741	CTITA	86/ 7/17	90/6/0
NCT	NORTH CRESENT	60	33.79	152	55.57	1079	USGS	88/ B/14	30, 0, 0
NEA	NENANA	64	34.63	149	4.63	364	GIUA	81/3/5	
NGL	NORTH GASLINE	60	49.25	149	59.89	122	USGS	74/ 9/26	76/ 3/27
NIK	NIKOLSKI	52	58,46	168	51.11	207	NOAA	71/ 5/17	76/ 3/27
NIN	NINILCHIK	60	0.67	151	32.13	110	USGS	71/ 8/28	72/ 8/24
*NKA	NIKISHKA	60	44.58	151	14.28	100	USGS	71/ 9/14	
NKI	NIKOLSKI	52	56.56	168	51,45	8	NOAA	76/ 3/27	85/12/26
NNL	NINILCHIK	60	2.66	151	17.36	381	USGS	72/ 8/24	
NRA	NORTH RIVER	63	53.51	160	30.86	107	GIUA	76/ 8/ 0	78/ 6/ 8
NTK	NUNITAK	59	52.66	139	2.11	1050	USGS	74/ 9/ 9	76/ 5/31
occ	OCEAN CAPE	59	32.55	139	51.61	22	USGS	76/ 1/ 1	76/ 4/24
OPT	OIL POINT	59	39.16	153	13.78	450	GIUA	88/10/ 1	
PAX	PAXSON	62	58.25	145	28.12	1130	GIUA	88/10/ 1	
PCA	PAXSON [see PIN] POINT CAMPBELL PEDRO BAY PINNACLE PALMER (USGS) PORT MOLLER PALMER EAST PALMER OBSERVATORY ARCTIC VALLEY PENINSULA PORTAGE POTTER	61	0 57	1 5 0	0 02	0.0	11000	72/0/4	74/ 9/30
PCT	PEDDO BAY	67	47 27	120	11 55	305	0565	70/0/0	14/ 9/30
PUB	PEDRO BAI	59	5 90	134	15.00	975	11565	78/ 9/ 9	
D1.D	DAIMED (HEGG)	61	35 53	140	7 95	100	USGS	84/9/20	
AMG	PORT MOLLER	55	58.72	160	29.83	315	GIUA	91/8/1	
PME	PALMER EAST	61	37.90	149	1.70	232	NOAA	80/ 8/19	90/ 3/ 1
PMR	PALMER OBSERVATORY	61	35.53	149	7.85	100	NOAA	67/ 9/ 1	
PMS	ARCTIC VALLEY	61	14.68	149	33.63	716	NOAA	67/ 5/25	
*PNL	PENINSULA	59	40.06	139	23.82	585	USGS	74/ 9/ 2	
*PRG	PORTAGE	60	51.87	149	1.21	55	USGS	72/ 8/29	
PTR	POTTER	+ -	0					/	74/ 9/15
PWA	HOUSTON		39.05		52.72	137	NOAA	77/ 7/ 1	
PWL	PORT WELLS		51.56		20.09	549	USGS	74/8/3	
	RAG		23.22		40.51	739	USGS	B4/ 7/22	
	RASPBERRY ISLAND		3.63		9.55	520	GIUA	75/10/ 0	01/7/10
	RICHARD D. SIEGRIST		49.59		8.68	510	GIUA	77/ 6/12	91/ 7/18
	REDOUBT		34.39		24.32 46.31	930	USGS	71/ 8/ 9 81/11/10	90/ 7/19
	REDOUBT VOLCANO RAGGED MOUNTAIN		25.19 13.15		32.74	1064 610	GIUA USGS	74/ 9/30	77/ 8/16
	RIOU		52.65		13.80	15	USGS	76/8/3	81/ 9/28
KIU	REINDEER		24.37		51.17	991	GIUA	88/10/ 1	01/ 3/20
BUN	REMOTE		41.47		12.21	470	GIUA	71/ 9/22	74/10/31
	SAWMILL		48.49		19.98	740	USGS	73/ 8/31	-,, - ,
	SHEEP CREEK FACILITY				2.35	67	GIUA	71/ 9/21	75/10/ 7
	SHEEP MOUNTAIN		50.00		19.66	1020	GIUA	91/ 8/ 1	
	SCOTTY LAKE		19.15	150	17.83	140	GIUA	71/ 9/21	75/ 6/ 0
SDE	SADIE COVE	59	26.60	151	16.92	770	USGS	83/ 7/ 4	84/ 6/28

TABLE 1 (cont)

		N LA	TITUDE	W LON	NGITUDE	ELEV			~	
CODE	SITE NAME	DEG	MIN	DEG	MIN	M 	INST	OPEN	CLO	OSED
SDG	SOURDOUGH SAND POINT SHERMAN GLACIER SHUYAH ISLAND SITKINAK ISLAND SITKANAK ISLAND SITKALIDAK ISLAND SITKALIDAK ISLAND SKILAK SKWENTNA SKILAK SELDOVIA SHEMYA SPURR SUSITNA SUNSHINE POINT SUSITNA STEPHENS GLACIER STONY RIVER SUCKLING HILLS SAVOONGA SPARREVOHN SEWARD TANA GLACIER TALKEETNA MOUNTAINS	62	31.62	145	32.60	625	GIUA	86/ 1/ 0		
SDN	SAND POINT	55	20.48	160	29.83	23	AAON	78/10/11		
SGA	SHERMAN GLACIER	60	32.04	145	12.42	424	USGS	76/ 8/16		
SHU	SHUYAH ISLAND	58	37.68	152	20.93	34	GIUA	74/ 0/ 0	90/	8/27
511	SITKINAK ISLAND	56	33.60	154	10.92	500	GIUA	75/ 8/ 9 40/ 0/ 0		
SIT	SITKA	5 / 61	1 00	130	24 20	705	NUAA	79/12/ 5	80/	3/27
SKU	SIEVER CITI	57	9.85	153	4.82	135	GTIIA	70/ 1/ 1	807	3/2/
*SKL	SKILAK	60	30.86	150	12.96	640	USGS	71/ 9/ 9	84/	7/28
*SKN	SKWENTNA	61	58.82	151	31.78	564	USGS	72/8/8	,	, _ ,
SLK	SKILAK	60	30.74	150	13.26	655	USGS	84/ 7/29		
SLV	SELDOVIA	59	28.28	151	34.83	91	USGS	72/ 9/30	85/	6/29
SMY	SHEMYA	52	43.85	174E	6.18	58	NOAA	70/11/17		
SPU	SPURR	61	10.90	152	3.26	800	USGS	71/ 8/10		
รรห	SUSITNA	61	27.83	150	44.60	1297	USGS	72/ 8/15		
SSP	SUNSHINE POINT	60	12.30	142	49.80	305	USGS	74/ 9/10		<b>~</b> / <b>~</b>
SST	SUSITNA	61	26.05	150	46.82	780	USGS	71/ 8/24		
STG	STEPHENS GLACIER	61	25,24	146	23.69	1326	USGS	74/ 7/11 72/ 7/24		
SII	STONI KIVER	2.0 2.1	4 42	104	46 62	454	USGS USGS	74/10/ 2		
SUC	SAVOONGA	63	41 70	170	28 80	15	CTIIA	77/ 8/17	63/	8/10
SVW	SPARREVOHN	61	6.49	155	37.30	762	NOAA	67/8/0		
SWD	SEWARD	60	6.22	149	26.96	91	USGS	72/ 8/22		
TGL	TANA GLACIER	60	45.35	142	49.78	1234	USGS	88/ 7/ 1		
TLK	TALKEETNA MOUNTAINS	62	29.63	147	52.68	1719	USGS	74/ 7/10	76/	7/ 1
TMW	TANA GLACIER TALKEETNA MOUNTAINS TOK MICROWAVE TANANA TOLSONA TSINA TATALINA TERENTIEV LAKE TZERO VALDEZ VALDEZ SOUTH VALDEZ WEST	63	19.28	142	59.48	488	GIUA	90/ 8/20		
TNN	TANANA	65	15.40	151	54.70	504	GIUA	65/ 1/ 0	79/	1/3
TOA	TOLSONA	62	6.29	146	10.34	909	NOAA	71/ 9/15		
TSI	TSINA	61	13.57	145	20.24	1113	USGS	76/ 8/15	85/	7/17
ATT	TATALINA	62	55.80	156	1.32	914	NOAA	78/ 9/20		- 4- 4
TTV	TERENTIEV LAKE	61	3.29	147	7.29	533	USGS	84/ 9/19		
TZ0	TZERO VALDEZ	63	48.16	145	44.10	602	GIUA	87/10/ 0	89/	7/17
7 V L Z	VALDEZ COUTU	61	2 65	146	18 32	669	USGS USGS	71/ 9/ 2 72/ <b>7</b> /22	76/	8/15
V23	VALDEZ SCOTH	61	3 54	146	33 24	796	11565	72/ 7/17	707	8/13
WAX	VALDEZ WEST WAXELL RIDGE	60	26.89	142	51.06	991	USGS	75/ 8/22		
WHC	WHITEHORSE	60	44.20	135	5.90	732	EMRC	71/ 9/ 1		
	WILLOW MOUNTAIN		46.42		11.88	985	USGS	71/ 9/ 1	72/	8/ 0
WRG	WHITE RIVER GLACIER	60		142	1.97	550	USGS	74/ 9/10		
WRH	WOOD RIVER HILL		28.28	148	5.39	314	GIUA	83/ 9/16		
	SELDOVIA		27.28		40.30	320	GIUA	91/ 6/20		
	YAHTSE		21.51		44.70	2135	USGS	74/ 9/ 5		
	YELLOWKNIFE ARRAY		29.59		36.32	200	EMRC	62/ 0/ 0		
	YELLOWKNIFE		28.70		28.40	198	EMRC	64/ 7/15	0.5 /	
	YAKATAGA	60	4.20		25.33 52.62	46	USGS	72/10/ 8		9/3
	YAKUTAT YAKUTAT		27.10 33.23		43.50	372 40	USGS NOAA	72/10/ 6 78/ 9/20	14/	9/30
	BRADLEY LAKE		45.83		53.38	622	USGS	80/10/11		
	BRADLEY LAKE NE		54.65		39.13	1219	USGS	80/10/12	84/	6/28
	BRADLEY LAKE NW		50.25		10.15	582	USGS	80/10/12		6/28
	BRADLEY LAKE SE		42.33		40.25	975	USGS	80/10/10		7/ 2
	BRADLEY LAKE SW		38.46	151		951	USGS	80/10/12		6/28
	HEATHER LAKE	61	1.75		55.88	3	USGS	83/ 7/14		9/30
CBKT	KADIN LAKE	61	6.63	147	11.61	275	USGS	83/ 7/15	83/	9/30

TABLE 1 (cont)

		N LA	TITUDE	W LON	NGITUDE	ELEV			
CODE	SITE NAME				MIN		INST		CLOSED
	~				<del>-</del>			~	<b>-</b>
CBLB	LONG BAY	60	57.93		13.30	5	USGS	83/ 7/15	83/ 9/30
CBUI	UNAKWIK INLET	60	54.66	147	32.72	8	USGS	83/ 9/11	83/ 9/30
ECNI	CHENEGA ISLAND	60	16.94	148	2.11	4	USGS	87/ 7/ 9	87/ 8/ 7
EFLB	FOUL BAY	60	34.94	148	3.65	4	USGS	87/ 7/ 9	87/ 8/ 7
EGRI	GREEN ISLAND	60	17.53	147	23.37	4	USGS	87/ 7/ 9	87/ 8/ 7
ELAI	LATOUCHE ISLAND	60	3.68	147	49.10	4	USGS	87/ 7/ 9	87/ 8/ 7
<b>EMTC</b>	MONTANA CREEK	62	9.55	150	1.00	131	USGS	86/ 8/28	86/ 9/11
ENKI	NAKED ISLAND	60	37.84	147	27.55	4	USGS	87/ 7/ 9	87/ 8/ 7
<b>EPKS</b>	PARKS HIGHWAY	62	15.30	150	14.50	وو ·	USGS	86/ 8/28	86/ 9/11
ETAL	TALKEETNA	62	18.25	150	4.92	137	USGS	86/ 8/28	86/ 9/11
<b>EWTF</b>	WEST FORK	62	19,13	150	28.55	200	USGS	86/ 8/28	86/ 9/11
GALN	GALENA BAY	60	57.47	146	44.48	3	USGS	72/ 7/14	72/ 8/13
<b>JACK</b>	JACK BAY	61	0.26	146	31.03	30	USGS	72/ 7/14	72/ 8/13
LOWE	LOWE RIVER	61	4.17	146	2.57	58	USGS	72/ 7/14	72/ 8/13
MNRL	MINERAL CREEK	61	12,60	146	19.16	236	USGS	72/ 7/14	72/ 8/13
SENW	SAINT ELIAS NW	60	13.80	140	54.40	1250	USGS	79/ 7/20	79/ 8/17
SHOU	SHOUP BAY	61	7.84	146	34.26	15	USGS	72/ 7/14	72/ 8/13
SWML	SAWMILL BAY	61	3.94			15	USGS	72/ 7/14	72/ 8/13

The instrumentation used in the USGS seismograph network is illustrated in the block diagram in Figure 2. The standard equipment at each field site includes: a vertical seismometer with a natural frequency of 1.0 Hz (Mark Products, Model L-4C); an electronics package consisting of an amplifier, voltage-controlled oscillator (Develco Model 6203 or equivalent) until about 1980, and then an A1VCO with calibrator and gain-ranging (Rogers and others, 1980); and either "air-cell" storage batteries, or a solar panel with 80 amp-hr storage batteries.

Data are telemetered via a combination of VHF (162-174 MHz) radio links and leased telephone circuits, some of which use satellite links having a 0.27 s transmission delay per hop. The radio equipment consists of low-power (100 mW) transmitters and receivers, many adapted from HT-200 Motorola handie-talkie transceivers, and either Yagi antennae with 9 db directional gain (Scala, Model CAS-150) or log-periodic antennae (Scala, Model CL-150). At the receive sites, where the seismic signals enter the telephone circuits, base-station radio receivers (G.E. Model R46AP66B) with greater sensitivity are used. The central recording facility incorporates a bank of discriminators (USGS designed NCER J101 or Develco Model 6203), four 16 mm-film 20-channel oscillographs (Teledyne Geotech Develocorder, Model RF400 and 4000D), a 14-track FM magnetic tape recorder (Bell and Howell Model VR3700B), three 3-channel drum recorders (Teledyne Geotech Helicorder, Model RV301B), and a time-code generator (Datum, Model 9100).

The principle of operation is as follows: The seismometer translates ground velocity into an electrical voltage that is fed into the amplifier/VCO unit. There the amplified voltage causes the frequency of VCO to fluctuate about its center frequency. The frequency-modulated (FM) tone from the amplifier/VCO unit is carried directly to the recording site by VHF radio links and/or voice-grade telephone circuits. Signals from eight seismograph stations can be transmitted on a single telemetry circuit using standard frequency division multiplexing techniques with a 340 Hz separation between carriers and a constant bandwidth of 250 Hz per channel. The channel center frequencies range from 680 to 3,060 Hz. A ninth channel with center frequency of 340 Hz and 125 Hz bandwidth is also commonly used. At the recording site the FM seismic signal is demodulated by a discriminator. The demodulated signal, which is simply an amplified and filtered form of the initial signal from the seismometer, is recorded on the oscillograph and tape recorder together with time signals from the time-code generator. Twenty-four hours of data from 18 stations can be recorded on a single 43 m-long roll of 16-mm film, while data from nine stations can be recorded on a single track of a 2,195 m-long, 14-track tape. Several stations are also recorded on Helicorder records for monitoring purposes.

Figure 3 illustrates the response characteristics of the entire seismic system from seismometer to film viewer. The response level at each station is adjusted in steps of 6 decibels so that the ambient seismic noise produces a small deflection of the trace on the film. As a result, the actual response for an individual station may differ from that of the typical station by a factor of 2, 4, 8, etc. The magnification of the typical station is about 6 x 10<sup>4</sup> at 1 Hz and 1 x 10<sup>6</sup> at 10 Hz.

Digital seismic-event recorders were developed internally and deployed to temporarily augment network recording in areas of interest. A description of these instruments can be found in Rogers and Lahr (1986) and Fogleman and Rogers (1987).

#### DATA PROCESSING

The 16-mm films, magnetic tapes, and Helicorder records are mailed weekly from the Alaska Tsunami Warning Center in Palmer, Alaska to the USGS in Menlo Park, California where the seismic data are processed by the following multi-step routine:

- 1. <u>Scanning</u>. The scan film, which records data from 18 stations distributed throughout the network, is scanned to identify all seismic events, including those of local, regional, and teleseismic origin, and to note the earliest P-arrival time, the time interval between the P and S phases (S-P time), and the duration of the signal (see section on Magnitude) for the first 3 stations. Stations not recorded on the scan film are arranged on the remaining films by geographic region.
- 2. <u>Timing</u>. For the "well-recorded" local earthquakes identified in the scanning process, the following data are read from each station: P- and S-wave arrival times, direction of first motion, duration of signal in excess of a given threshold amplitude, and period and peak-to-trough amplitude of maximum recorded signal. The P and S times are assigned weights according to the reader's confidence of the precision of the picks. The precision is influenced by the impulsiveness of the phases and the recording quality. Weights range from a full weight (coded 0) for the highest quality readings to no weight (coded 4) for times too poor to be used for hypocenter determination.

From October 1971 through the end of September 1973, the criteria for choosing earthquakes to be timed were based on S-P times and the number of stations which clearly record the earthquake. Beginning with October 1973, an additional criterion based on the signal durations is used. The area within which earthquake locations are routinely determined has changed a number of times since 1971. The various criteria used to determine which earthquakes to time from October 1971 to May 1989 are summarized in Appendix C. From September 1, 1985 to May 31, 1989, for example, the area is bounded approximately by longitudes 156° and 138°W and by latitudes 58° and 62.5°N, and is subdivided into western and eastern regions at longitude 145°W. In the western region, only events with average signal durations longer than 30s are routinely timed. In the eastern region, all earthquakes that are recorded by at least three stations and that produce at least four clear arrivals are timed. These criteria were established to select from the large number of earthquakes recorded by the network those shocks that are of greatest interest to current research objectives. In areas where special studies are being conducted, exceptions to the standard criteria may be made in order to locate more events.

Until October 1982 all of the available data recorded on the Develocorder films were read for each timeable earthquake. Since then, in order to keep the data processing current, only the scan film and the film(s) which contain the stations in the region the earthquake occurred were timed. A record of which films are timed for each earthquake is kept with the daily scan sheet.

Due to changes in the distribution of stations, to variations in seismicity rates, and to modifications in the criteria used for selecting which events to locate, the number of shocks located each month varied with time (Figure 4). The large increase seen beginning with March 1979 is due mainly to the aftershock sequence of the February 1979 M<sub>s</sub> 7.1 St. Elias earthquake which elevated the level of activity until the mid 1980's. The gradual decay in the number of events per month from 1979 to 1988 is due partially to decay in the rate of St. Elias aftershocks and to the closure of 15 stations, 13 of which were located in and around the Yakataga seismic gap, in the summer of 1985. Large spikes in the monthly number between 1982 and 1988 are due mainly to aftershock sequences. The large increase beginning in 1989 is due predominantly to the integration of data processed by the Geophysical Institute of the University of Alaska in Fairbanks with the USGS southern Alaska data.

The bulk of the timing is done by projecting the seismic traces from the film onto a one-film wire-grid or four-film sonic (Astrue and others, 1983) computer-based digitizing table, where the P-and S-phases, maximum amplitude, and coda duration are input as x-y coordinates into a computer and reformatted for input into a hypocentral location program. Since the fall of 1983, some of the timing has utilized digital waveform data obtained by digitizing the daily FM magnetic tapes at 100 samples per second. In the latter case an interactive, computer-based processing system (Stevenson, 1978) is used to display the waveforms and to pick the phase data.

3. <u>Initial computer processing</u>. The phase data for the timed events are batch processed by computer using the program HYPOELLIPSE (Lahr, 1989) to obtain origin times, hypocenters, magnitudes and, if desired, first-motion plots for fault-plane solutions. The HYPOELLIPSE computer program determines hypocenters by minimizing differences between observed and computed traveltimes through an iterative least-squares scheme. In many respects the program is similar to HYPO71 (Lee and Lahr, 1972), from which it was derived. Important features available in HYPOELLIPSE, but not in HYPO71, include multiple velocity and delay models, calculation of confidence ellipsoids, and incorporation of a station-history data base to keep the station gains and polarities updated.

The earthquake locations are based on P- and S-arrivals. S-arrivals provide important constraints on epicenters of shocks outside the network, and depths of events in the Wadati-Benioff zones beneath Cook Inlet and the Wrangell volcanoes (see section on Focal Depths). For some large events timed from the films S-arrivals cannot be read at any station because the traces on the film overlap each other or are too faint to read. However, S-arrivals not readable from the films can often be picked on paper records generated from playbacks of the magnetic tapes.

4. Analysis of initial computer results. Each hypocentral solution is checked for: large traveltime residuals (see section on Analysis of Hypocentral Quality), a root-mean-square (RMS) residual greater than 1 s, a focal depth greater than 35 km in an area where no Wadati-Benioff zone is known to exist, individual station magnitudes which differ from the average event magnitude by more than 0.5, and a poor spatial distribution of stations. Events with potential timing errors are re-read and additional readings, including those from souces other than the USGS network, are sought for shocks with a poor distribution of recording stations.

Initially, when the network spanned a relatively small area, traveltime residuals greater than 0.4 s were checked for data that was cataloged. As the network expanded eastward, it became necessary to increase the threshold for checking traveltime residuals due to the increase in the average epicentral distance to the stations used to locate the earthquakes. Beginning with 1978 data, the criterion for checking traveltime residuals was to check: the five closest stations for P-residuals greater than 0.6 s and S-residuals greater than 1.0 s; the stations after the first five and up to 150 km from the epicenter for P- and S-residuals greater than 0.9 s and 1.5 s, respectively; and stations with an epicentral distance between 150 to 350 km for P-residuals greater than 1.5 s and S-residuals greater than 2.0 s. Stations with epicentral distances greater than 350 km were generally weighted out by HYPOELLIPSE and checked for P- and S-residuals greater than 2.0 s and 3.0 s, respectively.

5. Final computer processing. Poor hypocentral solutions are rerun with corrected and/or additional data, and the new solutions are checked for large residuals that might indicate remaining errors. Corrections are made as required before the final computer run. Generally no more than three computer runs were allowed for any earthquake by an analyst. Additional runs might be made by a geophysicist to correct problem events. An empirical study of 100 earthquakes from October and November 1979 revealed that correcting all large traveltime residuals for each solution regardless of the overall quality of the solution resulted in little improvement in hypocenter precision and was time consuming. The high-quality solutions with only a few large traveltime residuals generally changed less than one minute in latitude and longitude and one km in depth with additional re-reads and computer runs. Consequently, since January 1980 earthquakes with only a few large traveltime residuals are checked only if the analyst feels the hypocentral solution may be affected.

#### VELOCITY MODELS

Our experience with locating earthquakes in southern Alaska suggests that significant lateral variations are present in the velocity structure across the network. Such variations might be expected from the complex geology and tectonics of the region (e.g., Plafker, 1967; Page and others, 1986). Over the years different methods have been used to account for variations in the crustal structure across southern Alaska. There have also been changes in the location program HYPOELLIPSE which would modify the computed locations (Lahr, 1989). For these reasons, all of the earthquake data from October 1971 through May 1989 were relocated using a single set of control parameters. Three velocity models were used in locating the earthquakes (Table 2). The Southern Alaska Model is based on a study of earthquakes below the Kenai Peninsula (Model A, Matumoto and Page, 1969); the Northern Alaska Model was developed by the Geophysical Institute of the University of Alaska, Fairbanks (N. Biswas, personal communication, 1988); the Gulf of Alaska Model is based on a study of the 1987 and 1988 M<sub>s</sub> 7.6 Gulf of Alaska earthquakes (Lahr and others, 1988a) and their aftershocks.

It is recognized that a model comprised of uniform horizontal layers is a poor representation of the actual velocity structure in the vicinity of a subduction zone (Mitronovas and Isacks, 1971; Jacob, 1972, McLaren and Frohlich, 1985); however, such a model does have the advantage of simplifying the computation of traveltimes. In order to determine any bias that might result from the approximation, a set of events in the Benioff zone below Cook Inlet was relocated using a ray-

tracing program of E. R. Engdahl and incorporating a more realistic, three-dimensional velocity model (Lahr and others, 1974; Lahr, 1975). Hypocenter shifts due to the oversimplified flat-layer model ranged from near zero at a depth of 60 km to as great as 25 km at 160 km depth. The offsets were oriented in such a way that the dip of the Wadati-Benioff zone would appear to be too great for locations based on a flat-layered model.

The choice of which velocity model to use in calculating the traveltime from an earthquake to a given station is based on the location of the earthquake. Table 3 summarizes the assignment of starting depth, velocity model, and station delay models. Work continues on improving our modeling of the first-order velocity features of southern Alaska.

#### TRAVELTIME DELAY MODELS AND TRIAL FOCAL DEPTHS

Corrections for P-phase and S-phase traveltime delays are applied at stations in the network that have consistently large residuals for most earthquakes (Table 4). The particular correction that is used to locate an earthquake is determined by the region in which the earthquake occurs (see Table 3).

Additional corrections are applied at several stations to correct for telemetry delays associated with one or more satellite links (0.27 s transmission delay per link) used in the telephone relay of the signal and are kept updated in the station-history data base.

Because the range of depths of earthquakes in southern Alaska vary from the surface to almost 200 km, hypocenters are computed starting at several trial focal depths to find the best solution. First, fixed-depth solutions are computed at depths z = 0 and 75 km, and a free-depth solution starting at z = 75 km is determined. If the latter solution has z < 20 km and a significantly lower RMS than at z = 0, then it is taken as the final solution. If, on the other hand, the z = 0 depth has a significantly lower RMS, then a free-depth solution starting at z = 0is used has the final solution. If neither the free-depth solution starting at 75 km nor the fixeddepth solution has significantly lower RMS, then the solution with the lower RMS is considered the best. Alternatively, if the free-depth solution starting at 75 km depth has z greater than or equal 20 km, then a solution with z fixed at 7.5 km is computed. Of the two fixed-depth solutions at 0 and 7.5 km, the one with lower RMS is used as a starting location for a free-depth solution. If the latter solution is within 7.5 km of the free solution that started at 75 km, then the solution with the lower RMS is reported as final. If the solutions are more than 5 km apart, then the one with lower RMS is reported as final. In the Gulf of Alaska where the velocity structure is poorly understood and hypocenters are not well constrained all of the depths are fixed at 10 km (see Table 3 for the border which is used to define the Gulf of Alaska).

Table 2. Alaska Velocity Models

Model 1. Gulf of Alaska

Layer	Depth (km)	P velocity (km/s)
1	0-7	5.0
2	7-12.5	6.8
3	below 12.5	<sup>*</sup> 8.1

Model 2. Southern Alaska

Layer	Depth (km)	P velocity (km/s)
1	0-4	5.3
2	4-10	5,6
3	10-15	6.2
4	15-20	6.9
5	20-25	7,4
6	25-33	7.7
7	33-47	7.9
8	47-65	8.1
9	below 65	8.3

Model 3. Northern Alaska

Layer	Depth (km)	P velocity (km/s)
1	0-24.4	5.90
2	24.4-40.2	7.40
3	40.2-76	7.90
4	76-301	8.29
5	301-545	10.40
6	below 545	12.60

Table 3. Geographical boundaries used to assign starting depth, velocity model, and delay models.

Earthquake Location	Trial Depth (km)	Velocity Model	Delay Model
Gulf of Alaska*	10.	1	2
Southern Alaska (South of 62.5°N)	0., 7.5, 75.	2	1
Northern Alaska (North of 62.5°N)	0., 7.5, 75.	3	1

<sup>\*(</sup>south of a line connecting 56°N, 154°W; 59.5°N, 146°W; 59.7°N, 143°W, 58.2°N, 139°W, 57.9°N, 137.3°W; 56°N, 135°W).

Table 4. P-phase and S-phase traveltime delays

Station	Mod	el 1	Mod	el 2
	P delay (sec)	S delay (sec)	P delay (sec)	S delay (sec)
HTP	01.10	00.00	00.00	00.00
ILI	00.44	00.78	01.11	01.98
ILM	00.44	00.78	00.44	0.783
ILN	00.44	00.78	00.44	0.783
NIN	01.47	02.59	00.67	01.19
NKA	02.16	03.82	01.45	02.58
NNL	01.47	02.59	00.75	01.34
PWA	00.70	01.25	00.41	0.730
RDT	00.36	00.64	-0.74	-1.32
SKL	00.10	00.18	00.10	0.178
SLK	00.10	00.18	-0.23	-0.41
SPU	00.39	00.69	00.39	0.694
SSN	00.67	01.19	00.30	0.534
SST	00.67	01.19	00.67	01.19
RDTE	00.36	00.64	-0.74	-1.32
RDTN	00.36	00.64	-0.74	-1.32
RDTZ	00.36	00.64	-0.74	-1.32

#### **MAGNITUDE**

Magnitudes are determined from either the coda duration or the maximum trace amplitude. Eaton and others (1970) approximated the local Richter magnitude, the definition of which is tried to maximum trace amplitudes recorded on standard Wood-Anderson horizontal torsion seismographs, by magnitude based on maximum trace amplitudes recorded on high-gain, high-frequency vertical seismographs, such as those operated in the Alaskan network. The amplitude magnitude, XMAG or  $M_X$ , used in this catalog is based on the work of Eaton and his co-workers and is given by the expression (Lee and Lahr, 1972):

$$XMAG = \log_{10} A - B_1 + B_2 \log_{10} D^2$$
 (1)

where A is the equivalent maximum trace amplitude in millimeters on a standard Wood-Anderson seismograph, D is the hypocentral distance in kilometers, and  $B_1$  and  $B_2$  are constants. Differences in the frequency response of the two seismograph systems are accounted for in A. It is assumed, however, that there is no systematic difference between the maximum horizontal ground motion and the maximum vertical motion. The term  $-B_1 + B_2 \log_{10} D^2$  approximates Richter's  $-\log_{10} A_0$  attenuation function (Richter, 1958, p. 342), where  $A_0$  is the trace amplitude for an earthquake of magnitude zero as a function of epicentral distance as observed for earthquakes in southern California. The constants used are  $B_1 = 0.15$  and  $B_2 = 0.08$  for D = 1-200 km, and  $B_1 = 3.38$  and  $B_2 = 1.50$  for D = 200-600 km. These constants have not been calibrated for southern coastal Alaska.

Coda durations are also used for determining magnitude because the maximum trace amplitude is often off scale due to the limited dynamic range of the film recording. For small, shallow earthquakes in central California, Lee and others (1972) express the duration magnitude, FMAG or  $M_D$  at a given station by the relation:

$$FMAG = -0.87 + 2.00 \log_{10} T + 0.0035 d$$
 (2)

where T is the signal duration in seconds from the P-wave onset to the point on the Develocorder film where the peak-to-peak trace amplitude of the coda envelope measured on a film viewer with 20X magnification falls below 1 cm, and d is the epicentral distance in kilometers.

Comparison of XMAG and FMAG estimates from equations (1) and (2) for 77 southern Alaskan shocks in the depth range 0 to 150 km and in the magnitude range 1.5 to 3.5 reveals a systematic linear decrease of FMAG relative to XMAG with increasing focal depth. However, no systematic dependence of T on D has been found. The following equation, including a linear depth-dependence term but not a distance term, is therefore used for Alaska:

$$FMAG = -1.15 + 2.00 \log_{10} T + 0.007 Z$$
 (3)

where Z is the focal depth in kilometers.

The coda duration magnitudes calculated from the network data are systematically less than the magnitudes reported in the Earthquake Data File (EDF) of NOAA (Lahr and Stephens, 1983). Based on a preliminary analysis (John Lahr, unpublished data), the empirical relationship between bodywave magnitude m<sub>b</sub> and duration magnitude, M<sub>D</sub> is:

$$m_b = 1.4 M_D - 0.39 \tag{4}$$

The magnitude preferentially assigned to each earthquake in this catalog is the mean of the XMAG (equation 1) estimates obtained for USGS stations. When no XMAG can be determined, the mean of the FMAG (equation 3) estimates for USGS stations is reported. For some earthquakes no XMAG or FMAG estimates are available and a magnitude calculated by another organization is used when available.

## ANALYSIS OF HYPOCENTRAL QUALITY

Two types of errors enter into the determination of hypocenters: systematic errors limiting the accuracy and random errors limiting the precision. Systematic errors result mainly from incorrect modeling of the seismic velocity structure in the earth and from incorrect phase identification. Random errors arise primarily from timing errors; their effect on the solution for each earthquake can be estimated through the use of standard statistical techniques.

The HYPOELLIPSE computer program determines hypocenters by minimizing differences between observed and computed traveltimes through an iterative least-squares process. For each earthquake, HYPOELLIPSE calculates the lengths and orientations of the principal axes of the joint confidence ellipsoid. The one-standard-deviation confidence ellipsoid describes the region of space within which one is 68 percent confident that the hypocenter lies, assuming that the only source of error is random reading errors. The confidence ellipsoid is a function of the geometry of the stations recording a particular event, the velocity model assumed, and the standard error of the arrival times; it is a measure of the precision of the hypocentral solution (see descriptions of SEH and SEZ in Appendix A). Repeated readings of the same phases by four seismologists have established that the standard deviation is as small as 0.01 to 0.02 s for the most impulsive arrivals and as large as 0.10 to 0.20s for emergent arrivals. The confidence ellipsoids are computed for a standard deviation of 0.1 s and therefore likely overestimate the 68 percent confidence regions. The standard deviation of the residuals for an individual solution is not used to calculate the confidence ellipsoid because it contains information not only about random reading errors but also about the incompatibility of the velocity model to the data.

In a few extreme cases the value calculated for one of the ellipsoid axes becomes very large corresponding to a spatial direction with very great uncertainty. In these cases an upperbound length of 50 km is tabulated. In most hypocentral solutions, the epicentral precision (SEH) is better determined than the focal depth precision (SEZ) so that SEH is generally smaller than SEZ.

To fully evaluate the quality of a hypocenter one must consider both the size and orientation of the confidence ellipsoid and the root-mean-square (RMS) residual (see description of RMS in Appendix A). In addition to reflecting random errors, the RMS residual can be large due to the misfit of the velocity model to the actual velocities within the earth, misinterpretation of phases, and systematic timing errors. In areas where the velocity structure is accurately known, a large RMS residual would probably indicate errors in the phase data. If the assumed velocity model does not represent the true seismic velocity structure within the earth, the RMS residuals could be large and reflect the incompatibility; alternatively, the RMS residuals could be small and not indicate the actual error in a mislocated hypocenter.

Other parameters provided by HYPOELLIPSE that are helpful in evaluating the quality of a hypocentral solution are: 1) GAP, the largest azimuthal separation between stations measured in degrees at the epicenter. If GAP exceeds 180°, the earthquake lies outside the network of stations used to locate the shock, and the solution is generally less reliable than that for an event occurring inside the network. 2) D1, the epicentral distance in kilometers of the closest station used in the solution. Solutions where the calculated depth is greater than D1 generally have smaller SEZ values (better depth precision) than events that have calculated depths less than the epicentral distance to the closest station. 3) NP and NS, the number of P- and S-arrivals, respectively, used in the solution. The accuracy of the solutions generally improves with an increase in the number of P- and S-arrivals. The RMS residual may actually increase, however, if distant stations are included in locating an event, because the differences between the observed and calculated traveltimes commonly increase with increasing epicentral distance due to the errors in the assumed velocity model.

#### HYPOCENTER PRECISION

The precision of the hypocenters, or the relative location accuracy of neighboring events, is represented by the confidence ellipsoids. The precision of epicenters, expressed in terms of the maximum semi-axis of the projected one-standard-deviation confidence ellipsoid (SEH), averages 4.4, 1.7, and 1.9 km, respectively, in the eastern (east of longitude 145°W.), central (between longitudes 145° and 150°W.) and western (west of longitude 150°W.) parts of the network. Similarly, the precision of focal depth (SEZ) averages about 4.7, 3.0 and 3.3 km, respectively. The variation in the precision of hypocenter determination across the network is strongly influenced by differences in the station density in the different regions. Hypocenter biases equal to and larger than the dimensions of the confidence ellipsoids are not unlikely as a consequence of the over-simplified velocity models assumed in the preparation of this catalog.

#### FOCAL DEPTHS

Previous studies (e.g., Francis and others, 1978; Lilwall and Francis, 1978; Uhrhammer, 1980; and McLaren and Frohlich, 1985) have shown that the accuracy of focal depths for shocks occurring in the vicinity of a seismic network is primarily a function of the geometry of the network, the number of P- and S-phase arrivals read, and the adequacy of the assumed velocity model. Depths are generally more accurate for earthquakes where the distance from the epicenter to the closest station (D1) is less than the calculated focal depth used for events located within the network or on its periphery. The accuracy of focal depths usually increases as the number of S-phase arrivals increases; however, systematic S-phase timing errors (due to mistaken identification of a converted

phase as S) or "bad" S picks can degrade focal depth estimation accuracy by several kilometers even when azimuthal coverage is good (Gomberg and others, 1990). Focal depths for shallow (depth less than about 20-30 km) shocks within the southern Alaska network generally are not well constrained due to the relatively large distances between stations and to a lack of knowledge about the velocity structure. Calculated depths for the same event can vary by several kilometers depending on the number of P- and S-phase arrivals used in the location, the trial focal depth, the velocity model, and the P-phase traveltime corrections used to locate the earthquake. Ambiguity in the calculated depth occasionally arises in cases where the traveltimes to receiving stations are similar for upward-leaving rays from a deep source and for downward-leaving rays from a shallow source; this situation leads to double minima in the variation of RMS residuals with depths and is common for events outside the boundaries of the network.

#### COMPLETENESS OF CATALOG

The magnitude threshold at which this catalog is complete varies geographically as a function of the density of stations and the criteria for timing earthquakes (see section on Data Processing and Appendix C). This catalog contains a number of time periods for which the seismic data collected by the project have not yet been processed:

November and December 1971
January, November and December 1972
May, June and August 1974
November and December 1975
December 1976
January - September 1977

The most thorough checking has been completed for time periods that have published catalogs (Appendix B) and for data processed since 1981. No attempt was made to assess the completeness for the time periods before 1981 that do not have published catalogs.

#### SEISMICITY OF SOUTHERN ALASKA

Reviews of the seismicity of the Aleutian arc, continental Alaska, and western Canada have been written by Taber and others (1991), Page and others (1991), and Rogers and Horner (1991), respectively. Meyers (1976) and Espinosa (1984) reviewed teleseismically recorded earthquakes in Alaska. The purpose of this section is to discuss briefly the distribution of the seismicity within southern Alaska for 50,048 earthquakes and other events recorded and located by the USGS seismograph network in southern Alaska for October 1971 to May 1989, and to cite the published literature written by the authors. No attempts were made to remove quarry and mine blasts or glacial earthquakes from the data set. The gross features of the pattern of seismicity in southern Alaska are shown in Figures 5-9.

The oceanic Pacific plate is being subducted beneath the continental North American plate (Alaska) along the Aleutian megathrust, which crops out on the seafloor at the Aleutian trench

(Figure 1). The seismicity related to various tectonic elements can be divided into five distinct source zones as follows: 1) Alcutian megathrust earthquakes along the interface between the subducting Pacific plate and the overriding North American plate; 2) subsea earthquakes within the Pacific plate beneath or seaward of the trench; 3) Wadati-Benioff Zone (WBZ) earthquakes within the subducted part of the Pacific plate landward of the trench; 4) overriding plate earthquakes in the North American plate exclusive of those along the volcanic arc; and 5) volcanic axis earthquakes within the North American plate along the axis of active volcanoes.

#### Aleutian megathrust earthquakes

Historically, most of the seismic energy in southern Alaska is released in major earthquakes that rupture the shallow part of the dipping thrust between the subducting and overriding plates. In September 1899 a pair of moment magnitude (M<sub>w</sub>) 8.1 shocks ruptured the transitional segment between Yakutat Bay and Kayak Island (McCann and others, 1980). In 1964 the convergent plate boundary west of Kayak Island slipped approximately 15 m in a M<sub>w</sub> 9.2 earthquake (Plafker, 1969; Hastie and Savage, 1970), the second largest earthquake of this century worldwide. Beneath a certain depth, estimated to be about 40 km in the Aleutian arc (Davies and House, 1979) and about 20 km in the northern Prince William Sound region (Page and others, 1989), the interface slips aseismically.

The coastal region from about Kayak Island to Yakutat Bay had been identified by the early 1970's as a seismic gap due to the lack of significant shocks since the turn of the century (Tobin and Sykes, 1968; Kelleher, 1970; Sykes, 1971). In 1979, the M<sub>s</sub> 7.1 St. Elias earthquake (Figure 5) occurred at the eastern end of this gap and involved low-angle, north-northwest-directed thrusting (Hasegawa and others, 1980) at depths between 10-15 km (Stephens and others, 1980) in a zone that at least locally is no more than 2-3 km thick (Page and others, 1984). The Yakataga seismic gap, which extends westward from the western limit of the St. Elias aftershock zone to the eastern extent of the 1964 rupture, is considered a likely site for a great (M<sub>s</sub> 7.8 or greater) thrust earthquake within the next few decades (McCann and others, 1980; Lahr and others, 1980; Jacob, 1984; Stephens and others, 1992).

Seismicity of the gap region has been monitored nearly continuously since 1974 when the network was extended across this region (see Appendix C). East of longitude 145°W, the apparent high rate of shallow activity is due at least in part to a lower magnitude threshold used in selecting events for processing. During this period the spatial distribution of seismicity landward of the continental margin has remained relatively stable and is characterized by broad concentrations of shallow seismicity (depths less than 30 km) beneath Icy Bay, Waxell Ridge, and the Copper River delta separated by areas of relative quiescence with most of the shocks occurring in a zone that extends about 90 km inland from the coast (Figure 7 and Figure 9, section D). The most active area has been the aftershock zone of the St. Elias earthquake, both prior to and since the mainshock. Hypocenters of better-recorded shocks from the diffuse patch of seismicity near Waxell Ridge at the center of the gap concentrate near a depth of 12 km (Stephens and others, 1992), which is comparable to the depth to the thrust plane defined by seismicity beneath the St. Elias area. The Copper River Delta seismicity on the western edge of the gap is discussed in the section on Wadati-

Benioff zone earthquakes. The largest earthquake located within the gap since 1971 is a  $M_x$  4.1 shock on December 1980 near the southwest corner of the Waxell Ridge concentration of seismicity.

#### Subsea earthquakes

There is very little historical evidence of seismicity in the Gulf of Alaska seaward of the continental margin. However, in November 1987 and March 1988, two M<sub>s</sub> 7.6 strike-slip earthquakes in the northern Gulf of Alaska ruptured a composite 250-km-long, north-striking zone in the Pacific plate (Figures 5 and 7) south of the Yakataga seismic gap (Lahr and others, 1988a). These events are thought to reflect shear stress in the Pacific plate seaward of the boundary between the locked Yakataga seismic gap and the recently slipped 1964 rupture.

## Wadati-Benioff zone earthquakes

Wadati-Benioff seismic zones are associated with both the northeast-to-north-trending Aleutian volcanic arc west of Cook Inlet and the east-southeast-trending Wrangell volcanic arc east of longitude 145°E. Below 30 km depth the distribution of earthquakes is dominated by activity within the Aleutian Benioff zone west and north of the Cook Inlet region (Figures 5 and 6, and Figure 9, sections C-E). This zone dips west-northwestward from 8-10° beneath the Kenai Peninsula and Anchorage, and then steepens to about 65° west of Cook Inlet, although the latter dip may be overestimated by abut 10-15° due to the systematic bias introduced by using a flat-layered velocity model. Beneath both the southern Kenai Peninsula and Anchorage the upper surface of this zone is at a depth of 30 to 35 km, and beneath the volcanic arc west of Cook Inlet it is at a depth of about 100 km. The average dip of the shallow (depth < 30 km) part of the zone decreases from about 7° beneath Prince William Sound (Page and others, 1991) to about 3-4° beneath the continental margin. On a regional scale using only events with well-constrained focal depths, the maximum thickness of the Aleutian WBZ is about 15 km. The zone appears to thicken slightly beneath the southern Kenai Peninsula and Anchorage. The maximum focal depths varies along the length of the zone from about 200 km near Iliamna (Figure 9, section E) to about 160 km beneath the northernmost Cook Inlet Volcano (Figure 9, section D) and to about 150 km beneath the Alaska Range (Figure 9, section C). Beneath continental Alaska, the Aleutian WBZ is made up of a single zone of earthquakes, in contrast to the two zones of seismicity beneath the eastern Aleutian arc near the Shumagin Islands (Taber and others, 1991). The geometry of the Aleutian WBZ changes along strike as documented in studies of both teleseismically and regionally recorded earthquakes (Van Wormer and others, 1974; Lahr, 1975; Davies, 1975; Agnew, 1980; Pulpan and Frohlich, 1985). In lower Cook Inlet near latitude 59°N, the strike of the deep WBZ rotates 15° counterclockwise (Pulpan and Frohlich, 1985), and about 350 km farther north the strike swings about 35° clockwise (Agnew, 1980). The deeper seismicity east of the Cook Inlet region appears to be bounded by a northwestsoutheast trending line, which passes about 50 km northeast of Valdez (Figure 6). Such a line approximately delineates the northeastern terminus of the Aleutian Benioff zone (Stephens and others, 1984a). The diffuse appearance of the Aleutian Benioff zone in Figure 9, section C, may be attributed in part to a lack of focal depth control for earthquakes north of the USGS network (north of 62°N).

The magnitude of the largest shocks recorded by the network within the southern Alaska segment of the Aleutian WBZ since 1971 have been in the magnitude 6 class (Figure 5). In July and September of 1983 the two M<sub>s</sub> 6.3 (m<sub>b</sub> 6.2) Columbia Bay shocks occurred in the shallow part of the WBZ beneath the north coast of Prince William Sound and involved normal faulting on steeply northwestward-dipping planes between depths of 22-35 km (Page and others, 1985, 1989). A M<sub>x</sub> 6.1 (m<sub>b</sub> 5.5) earthquake occurred beneath Mt. Spurr in November 1988 with a calculated depth of 138 km. The focal mechanism for this event using local network data exhibits down-dip tension and along strike compression which is typical for the deep WBZ zone activity in the Cook Inlet region (Lahr, 1975; Engle, K. Y., 1982; Pulpan and Frohlich, 1985).

The rate and character of seismicity between the shallow and deep parts of the Aleutian WBZ are significantly different, with the higher level of activity in the deep part of the zone beneath the western Kenai Peninsula, Cook Inlet, and the volcanic arc. Since 1971 the distribution of seismicity generally has not been uniform and has a significant component of spatial and temporal clustering. Clustering is more common in the shallow WBZ with mainshocks followed by energetic aftershock sequences such as the aftershock sequences following the two 1983 Columbia Bay earthquakes. In contrast, few, if any, aftershocks have been recorded following WBZ shocks in the deep part. Three persistent, diffuse concentration of events not associated with an aftershock sequence are located about 40 km northeast, 150 km southeast, and 75 km southwest of the Columbia Bay shocks beneath the Tazlina Glacier (latitude 61°N30'N, longitude 146°35'W), the Copper River Delta, and the northern end Knight Island, respectively. The Tazlina Glacier cluster is a north-northeast-trending swath of seismicity with depths between 25-45 km (Page and others, 1989). Preliminary relocations of well-recorded earthquakes in the persistent, low-magnitude seismicity beneath the Copper River Delta at the western edge of the gap occur at depths between 20-30 km and appear to occur below the megathrust based on one teleseismically recorded thrust event that occurred in 1970 with a depth of about 21 km (Stephens and others, 1992). The northern Knight Island source area was the focus of a special field recording effort in 1987 in which eight stations were operated for one month. Relocated hypocenters from that effort define a subhorizontal tabular zone, which is 6 km thick and whose upper surface is about 18 km deep (Lahr and others, 1988b). Spacial clustering in the deep WBZ is conspicuous (Figures 5 and 6) with concentrations of events deeper than 110 km beneath Mts. Iliamna and McKinley and northwest of Anchorage, while other parts are relatively quiet, such as in the vicinity of the two points 60.8°N, 152.5°W and 62.5°N, 150.2°W.

East of the Aleutian zone lies the weakly active Wrangell WBZ, which dips to the north-northeast beneath the Wrangell volcanoes (Stephens and others, 1984a; Page and others, 1989). This zone extends at least 150 km along strike and to depths of at least 100 km (Figure 6 and Figure 9, sections A and B). The Wrangell WBZ is about two orders of magnitude less active than the Aleutian zone. The largest earthquakes located in this zone by the network are a 4.5 m<sub>b</sub> (M<sub>x</sub> 3.7) event in September 1986 at a depth of 48 km and a M<sub>x</sub> 3.7 shock in February 1989 at a depth of 55 km. The deepest well constrained shock is a M<sub>D</sub> 2.4 shock in September 1986 at 97 km depth beneath Mt. Wrangell. The seismicity in the Aleutian and Wrangell zones appears to be continuous, at least in the depth range 20-45 km, and may define adjacent limbs of a buckle in the subducted plate (Page and others, 1989).

## Overriding plate earthquakes

Generally the distribution of shallow seismicity within the North American plate away from the volcanic axis since 1971 has been diffuse (Figure 7), and the majority of the events cannot be clearly associated with mapped fault traces. In a few cases, including the Talkeetna segment of the Castle Mountain fault, the Duke River fault near longitude 141°W, and the northernmost segment of the Fairweather fault, earthquakes are closely associated with mapped faults. The diffuse character of the seismicity north of latitude 62.5°N, south of latitude 59.5°N and east of longitude 138°W is at least partially attributed to these areas being outside the USGS seismograph network. The shallow seismicity does has three conspicuous concentrations: 1) a narrow zone parallel to the Duke River fault on the U.S-Canada border, 2) a diffuse north-northeast-trending band of seismicity near longitude 150°W extending between the Denali fault and Cook Inlet whose focal depths lie well above the Aleutian WBZ (Figure 9, section D), and 3) a band north of and parallel to the Castle Mountain fault east of longitude 151°W. To some degree, the apparent scatter of the seismicity near the Duke River fault reflects errors in the epicenters due to uncertainties in the velocity structure and the lack of nearby seismograph stations. Recent shocks located from both regional (Horner, 1983) and local (Power, 1988b) recordings reveal that the seismicity is mostly concentrated in the vicinity of the Duke River fault.

Since 1971 the largest shallow shocks in the overriding plate have been in the magnitude 5 class (Figure 5). The August 1984  $m_b$  5.7 Sutton earthquake ruptured a 10-km-long buried segment of the Castle Mountain fault and involved right-lateral slip on a steeply north-dipping plane (Lahr and others, 1986). This is the largest shock to be clearly associated with a mapped fault since the network was established. Two earthquakes, with magnitudes  $M_X$  5.3 ( $m_b$  5.3) and  $M_X$  5.0 ( $m_b$  4.9), occurred 3 days apart about 75 km northwest of Juneau and a few km east of the mapped trace of the Chatham Strait fault. A  $m_b$  5.7 ( $M_X$  5.6) shock occurred in November 1987 on the Duke River fault. The  $M_X$  5.6 ( $m_b$  4.7) shock of March 1989 was located 150 km west of Juneau near the entrance to Cross Sound is about 10 km west of the mapped trace of the Fairweather fault.

Earthquakes with reported depths of 30 km and deeper (Figure 6) east of longitude 145°W (excluding shocks beneath or more than 100 km southwest of the Wrangell Volcanoes), west and north of the Aleutian WBZ generally have poorly constrained focal depths and are probably in the overriding plate.

## Volcanic axis earthquakes

Small, shallow earthquakes are abundant along both the Aleutian (Kienle and others, 1983; Stephens and others, 1984b) and Wrangell (Page and others, 1989) volcanic axes. The shocks recorded by the network west of Cook Inlet form a diffuse band, approximately 30 km wide, punctuated by pronounced clusters (Figure 7 and Figure 9, sections D and E). The volcances are marked by dense clusters of earthquakes shallower than 5 km, whereas elsewhere along the Aleutian axis shocks fall in the depth range 5 to 20 km (Stephens and others, 1984b).

Both the 1976 and 1986 eruptions of Augustine Volcano in lower Cook Inlet were preceded and accompanied by large numbers of small earthquakes (Lalla and Kienle, 1986; Reeder and Lahr, 1987; Power, 1988). Although the earthquakes associated with these eruptions were too small to be located by the USGS network (and do not appear in the epicenter or cross section plots), the USGS stations were used to monitor activity during the 1986 eruption after all of the stations on the island had failed, possibly due to mudslides (Reeder and Lahr, 1987).

#### AVAILABILITY OF DATA

The summary and phase data for the 50,048 earthquakes and other events located by the USGS southern Alaska seismograph network and also the location program and control files used to locate all of the events are available on CD-ROM (Fogleman and others, 1993). Appendix B lists previously published catalogs available from the USGS, ESIC, Open-File Report Sales, Box 25286, Federal Center, Denver, CO 80225 (telephone: 303-236-7476). Information about the availability of this and other preliminary data can be obtained by contacting the principal investigators.

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#### APPENDIX A

## Magnitude 4 and Larger Earthquakes for October 1971 through May 1989

Earthquakes from southern Alaska are listed in chronological order. The following data are given for each event:

- Origin time in Universal Time (UT): date, hour (HR), minute (MN), and second (SEC).
   To convert to Alaska Standard Time (AST) or Alaska Daylight Time (ADT) subtract 9 or 8 hours, respectively.
- 2. North latitude (lat n) and west longitude (lon w) in degrees (deg) and minutes (min).
- 3. DEPTH, depth of focus in kilometers. A "\*" after the depth indicates the depth was fixed.
- 4. MAG, magnitude of the earthquake. The magnitude preferentially assigned to each earthquake in this catalog is the mean of the XMAG (M<sub>X</sub>) estimates obtained for USGS stations. When no XMAG can be determined, the mean of the FMAG (M<sub>D</sub>) estimates for USGS stations is reported. For some earthquakes no XMAG or FMAG estimates are available and a magnitude calculated by another organization is used when available. A letter following the magnitude indicates the type as follows:
  - X Amplitude magnitude (XMAG), USGS.
  - F Coda duration magnitude (FMAG), USGS.
  - B Body-wave magnitude (m<sub>b</sub>), USGS National Earthquake Information Service (NEIS).
  - S Surface-wave magnitude (M<sub>s</sub>), NEIS.
  - O Estimated magnitude, USGS.
- 5. NP, number of P arrivals used in locating earthquake.
- 6. NS, number of S arrivals used in locating earthquake.
- 7. GAP, largest azimuthal separarion in degrees between stations.
- 8. D1, epicentral distance in kilometers to the station closest to the epicenter.
- 9. RMS, root-mean-square traveltime residual in seconds:

$$RMS = \left[\frac{\sum_{i=1}^{N} W_{i}R_{i}^{2}}{\sum_{i=1}^{N} W_{i}}\right]^{\frac{1}{2}}$$

where R<sub>i</sub> is the observed minus computed arrival time of the i<sup>th</sup> arrival and W<sub>i</sub> is the corresponding weight of the P<sub>i</sub> S<sub>i</sub> or S-P reading used in the solution.

10. SEH, standard error in horizontal direction with least control in kilometers. SEH = MAXH/1.87, where MAXH is the largest horizontal deviation in kilometers of the one-standard-deviation confidence ellipsoid (see Figure 10). In previous catalogs MAXH was referred to as ERH. Values of SEH that exceed 50 km are tabulated as 50.0 km.

- 11. SEZ, standard error of depth inkilometers. SEZ = MAXZ/1.87 where MAXZ is the largest vertical deviation in kilometers of the one-standard-deviation confidence ellipsoid (see Figure 10). In previous catalogs MAXZ was referred to as ERZ. Values of SEZ that exceed 50 km are tabulated as 50.0 km.
- 12. Q, quality of the hypocenter. This index is a measure of the precision of the hypocenter (see section Analysis of Hypocentral Quality) and is calculated from SEH and SEZ as follows:

Q	Larger of SEH and SEZ (km)
Α	≤ 1.34
В	≤ 2.67
C	≤ 5.35
D	> 5.35

- 13. AZ1, DIP1, and SE1 are the azimuth in degrees (clockwise from north), dip in degrees, and length in kilometers of the most nearly horizontal of the three principal semi-axes of the one-standard-deviation error ellipsoid.
- 14. AZ2, DIP2, and SE2 are defined as above, but correspond to the principal semi-axis of intermediate dip.
- 15. AZ3, DIP3, and SE3 are defined as above, but correspond to the most nearly vertical principal semi-axis. Values that exceed 99.00 km are tabulated as 99.0 km.

		RELIMINARY DETERM	INATION OF	HYPOCE	MTERR	TM 80	CTHERN	ATARYS	- OCT	ORER	1 9 7 3	TO N	AY 1969		
origin time			mag np ne		d1	578-6	den						dip2 me2	A=3 (	fee frib
date hr mm sec		deg min km		deg	Rus	800	km	ka					deg ka		
71/10/29 13 16 36.8			4.67 15 2	-		0.99	0.9	1.4 8	30				17 1.5		
72/03/76 19 38 37.6			4.3P 15 0			0.29	1.3	1.5 8	7		1.0		38 1.6		
72/03/29 21 0 46.1			5.1B 15 2			0.39	1.4	1.1 3	_				35 2,9		
72/03/30 16 8 49.5			4.1x 13 0			0.57	0.7	1.5 B			0.6		19 1.0		
72/04/02 13 4 19.6			4.6F 16 0			0.33	0.0	1.1 A	94	â		3	5 1.1		
, a, - w		101 1017 1071				V.13	٧.0		,	•	1.3	•	J 4.1	~~~	** ***
72/04/07 3 16 24.8	59 58 0	152 24 5 98 8	4.6F 15 1	174	49	9.48	0.9	1.2 A	20	1		110	25 1.4	248	65 2.4
72/04/20 17 27 15.4			4.3F 16 2			0.39	0.8	0. B A	10			110	42 1.4		
72/05/07 9 15 14.2			4.1F 16 1		3		0.7	1.2 A	45		8.7		8 1.3		
72/06/10 22 50 38.2			4.0F 15 2		_										
72/06/14 0 52 35.3			5.0F 17 1			0.37	1.7	1.6 B		18			42 1.5		
72/00/14 0 32 33.3	00 17.0	132 39,0 100.3	3.00 17 1	/9	10	0.54	0.9	1.1 A	226	7	0.8	320	28 1.4	123	61 2.3
72/07/18 9 35 15.2	E0 32 2	184 6 6 748 A	4.67 11 2	237	41	0.25	3.4	3.1 c	7			276	41 7.6	34	49 4.1
72/00/17 13 48 3.5			4.17 9 1			1.97	1.3	2.4 B						_	
72/10/01 10 8 51,3			5.1r 15 0			0.29	1.7		15			108	24 1.3		
72/10/01 10 8 31.3								3.1 C		_		354	8 3.1	206	
			4.4x 5 0			1.01		50.0 D	271		11.4	1	0 99.0	٥	
72/10/21 19 52 6.0	63 2.4	130 53.3 129.4	4.6F 16 0	202	123	0.25	2.5	4.2 C	98	5	1.8	6	21 3.9	201	68 8.4
73/03/05 8 30 49.6	62 42 4	148 98 A 107 3	4.08 15 1	284	214	0.46	4.1	6. 6 D	270	٥		180	14 0.7	٥	76 12.6
73/05/18 18 32 57.2			4.18 15 0			0.18			78						
73/05/26 23 4 38.0			4.3F 15 0			0.32	1.7	4.9 C				175			
73/08/22 18 14 36,6								3.4 C				6		216	
73/08/31 2 30 57.0			4.3F 12 1		105		7.1	22.4 D	93			164	9 11.5		
13/06/31 2 30 31.0	•4 3.•	147 0.4 11.3	4.2F 15 0	62	19	0.33	0.5	G A	24	12	9.7	289	22 0.5	141	65 1.6
73/09/06 10 59 37.1	60 58 8	146 44 2 14 9	4.4F 15 0	65	13	0.38	0.5	0.4 A	251	13	0.6	152	36 0.9	358	51 0.7
73/09/10 6 53 36.0			4.0F 15 0			0.37		47.1 0	88			357	18 14.7		72 92.8
73/09/28 14 2 25.1			4.2F 16 0			0.21	0,9	2.9 C	54	_		144	5 1.6		83 5.5
73/10/27 7 5 46.9			4.0F 12 0				50.0					337	0 99.0		90 99.0
73/11/11 16 44 56.7			4.47 15 0		40	0.18	0.0	1.8 3	57			148			03 3.6
/3/11/11 10 44 30./	39 31.4	143 31,3 141.4	4.49 15 0	73	40	0.14	0.0	1.0 2	21	•	1.0	144	3 1.3	203	03 3.6
73/12/13 7 56 43.1	41 70 4	151 21 0 250 D	4.3F 13 0	240	73	0.66	11.8	23.7 D	76	2	2.4	169	25 5.0	331	64 49.3
73/12/31 17 4 51.6			4.0F 10 1			0.11	6.2	4.0 D		13 1			25 1.8		
74/01/07 8 27 5.2			4.5F 16 0			0.27	0.7	2.1 B	50			141			82 4.0
74/01/22 10 43 6.0			4.4F 16 0			0.19	0.8	3.1 c	276		1.2	6	2 1.4		
74/01/22 11 13 47.3			4.0F 17 1			0.28	0.7	1.6 B	24	-		293	9 1.0	155	76 3.0
10,01,11 11 15 11.5	01 35.0	2,0 24,12	1.00 11 1			V.2V	٠,,	1.4 2	**	•	4.2	2 93	, , ,	133	,. 3.0
74/01/24 18 43 27.4	61 31.4	147 26.2 20.0	4.27 15 0	65	35	0.36	0.5	1.3 A	288	G	0.5	198	14 0.7	18	76 2.5
74/02/02 15 55 29.0	61 29.9	147 25.6 19.8	4.3F 15 0	62		0.33	0.5	1.4 B	287			197	13 0.7	17	
74/02/04 14 6 52.2			4.0F 16 0			0.18	0.7	2.3 B				359	5 1.3		
74/02/05 2 25 23.0			4.47 17 0		95	0.33	0.	2.4	53		1.0		11 1.3	153	79 4.6
74/02/10 22 5 47.4			4.0F 16 1		71	0.33	0,6	1.8 8	33		0.9	301	7 1.1	131	83 3.4
								-,		~	-,-				J- 2/4
74/02/11 0 7 39.2	59 39,7	152 16.8 59.1	4.3x 15 2	143	45	0.12	0.5	2.0 2	264	8	0.9	356	17 1.3	150	71 3.9
74/02/15 6 6 29.5	63 6.4	150 40.8 118.8	4.3F 18 1	78	167	0.18	1.0	7.9 D	62	5	3.0	331	7 1.2	187	81 15.0
74/03/06 15 15 47.1			4.1F 15 1	76	22	0.17	1.1	3.5 C			1.8	66	13 1.4		
74/03/10 10 0 15.1	63 1.8	150 32.6 113.5	4.27 19 1	91	170	0.18	1.9	9.4 D	62	5	3,1	332	7 1.2	187	81 17.9
74/04/06 3 56 25.1			6.6F 17 0					27.4 D			4.4	41	34 56.0		
		···										_			
74/04/06 5 12 23.0			4.0F 16 0		91	0.39	9.8	0.8 D	279		1.4	182	42 24.4	18	47 3,0
74/09/08 5 27 32.6			4.1x 8 0	195	12	1,62	4.6	10.4 D	104	14	3,3	10	18 2.1	230	67 21.2
74/09/24 15 39 22.8	59 42,0	153 21,2 124,4	4.4X 14 2	83	57	0.27	1.0	2.0 B	47	2	1.1	130	17 1.6	310	73 3.4
74/10/09 7 38 10.0			6.0F 10 0	294	132	1.11	11.6	16,0 D	29	20	2.6	131	28 3,9	269	54 36.0
74/10/10 19 2 27.7			4.5x 15 3	264	418	0,74	27.6	48,1 D			5,1	6	26 31,4		64 99.0
														-	
74/11/02 5 6 24.2			4.0F 15 0			1.43		50.0 D		-		333	0 99.0	0	90 99,8
74/11/06 9 23 10.4			4.2F 15 1	-		0.32	1.1	3.0 C	39			129		290	80 5.7
74/11/14 4 48 51,8			4.47 15 0			0,32		15,5 D			0.9	54	6 0.8		
74/11/22 18 4 2.2			4.60 17 1		30		0.8				1,5	59			86 4,9
74/12/01 15 56 34.6	62 10.2	150 24.7 38.7	4.08 17 0	192	64	0.27	1.1	3.3 C	5#	1	0.5	328	7 1.9	156	83 6,1

		REINTHARY D	etermi nati	OF OF	HYPOCKE	TERS	IN 800	THERM	ALASKA -	OCTO	BER	1971	TO NO	Y 1989		
origin time	lat n		depth mag	np n				sah						dip2 se2		
date hama sea			lon		deg			ka	km					deg ku		
74/12/13 11 47 54.9					1 113		0.31	1.4	3.2 C					22 0.9		
74/12/17 18 40 25.5 74/12/29 18 25 2.4					2 341 0 102		0.19	0.5						25 99.0 11 0.8		
74/12/30 3 33 18.5					9 174		0.36	0.5	1.1 A					13 1.4		76 2.1
75/01/01 3 55 13.4					0 175		0.35	0.8	1.4 B					12 1.4		77 2.6
12/02/02 3 23 13:0	02 32.3	247 55.2	20.0 4.7			40	V.33	٠, ۵	2.4	414	•	٠,٠	143	** 1.4	4.	// 4.0
75/01/13 0 31 56.2	61 17.1	150 30.8	50.3 4.0	г 16	0 94	22	0.37	0.4	1.9 A	38	1	0.8	178	6 0.8	799	84 1.9
75/02/10 14 3 32.0					0 156		0,25	1.4	2.7 C			2.6		6 1.3	-	
75/03/06 9 2 15.7					1 160		0.30	4.0	7.2 D		16		36			
75/03/20 13 31 17,3					2 276			4.4		70				28 5,3		
75/04/30 4 29 11.4	60 6.0	153 14.8 1	37.4 4.2	r 15	1 148	26	0,26	1.2	2.6 B	316	3	2.2	46	5 1.2	195	84 4.5
75/05/1# 15 43 0.4					0 216			1.4	.2,€ 3			2,6		4 0.9		86 4.4
75/05/20 16 29 49.8					2 240			3.7	7.0 p					18 5,7		
75/05/21 6 34 55.3					0 123				50.0 D		-	0.7	1			90 99,8
75/05/25 19 4 31.9									50,0 D				331	0 13.9		90 99,8
75/05/31 23 35 20.2	37 39.1	156 11.5 1	44.4 4.3	1.3	1 284	61	0.16	9,6	4.3 D	331	2	2,1	239	40 23.9	13	30 2,7
73/06/01 13 10 57.3	50 30 6	162 52 4 1	15 1 4 2	. 16	2 135	80	0.36	1.1	2.4 B	45		1.0	136	14 1.9	205	75 4.7
75/06/24 7 15 5.5					4 169		1.25	1.3	3,1 C			2.3		4 1.3		
75/06/24 12 15 30.4					1 275			1.7		83				42 10,1		
75/06/29 10 45 28.8					0 160				48.3 D			1.6	84	23 1.8		
75/07/29 22 1 55,3				r 16	1 83		0.25	1.0	2.3 B	-		1.8	63			85 4.7
•																
75/08/21 22 19 22.5				F 16			9,37	0.5	1,2 A			0,5		6 0.7		
75/08/22 15 50 16.5				r 15			0,14	0.9	2,7 C				342			84 5.2
76/01/26 3 41 36.3					1 134		0.18	1.0					154			80 4.2
76/02/05 9 36 37.4	-				1 134		0,36	1.0	0.8 A					33 2,0		
76/02/19 7 49 31.8	62 20.5	151 21.0	87.3 4.0	7 15	1 273	41	0.60	3.0	3,0 C	72	•	1.5	170	44 2,6	334	45 7.6
76/03/14 3 31 48.1	60 0 5	162 20 6 1		- 14	0 124	25	0.31	1.0	2.2 B	50		٠.	140	14 1.6	200	76 4.2
78/04/10 19 36 33.0					1 124		1.65	0.8	2.4 B	43		1.2		1 1.4		
76/04/18 10 32 48.6				г 15			0.22	1.3	2.3 B	62	_	-		14 2,3		
76/05/03 17 47 30.8					1 199		0.08	2.8	5.6 D			1.5		23 1.8		65 11.6
76/06/17 2 44 56.4				r 11			0.26	6.3	8.5 D					25 9.9		
76/08/22 2 1 46.5	59 59.8	153 6.5 1	35.1 5.4	<b>7</b> 17	0 70		0.24	0.9	2.1 B	71		1,1		5 1.6		
76/08/30 8 17 52.2					1 148		0,10		-		_	1,4		13 2,2		
76/09/04 23 23 47.1					0 159			1.9	4.0 C			1.4	17			72 7.5
76/10/18 0 36 32.5					0 127			1,8	5,5 D			1.5		12 2.6		
76/10/22 18 35 31.7	56 18.0	153 9.6	10.0* 4.0	r 13	0 304	166	0.45	50.0	50.0 D	279	0	35.9	0	0 99.0	0	90 99,0
75/10/04 17 18 52 8	42 21 4	140 8 7 1	140 47	- 16	0 149	120		2.6	4,8 C	21		1 4	112	18 4.2	986	72 9.5
76/10/24 17 19 52.9 76/11/03 16 40 45.2					1 127			2.8	4,9 C			5.3		9 2.1		
76/11/06 0 8 18.9					0 125		0.24	1.1	2,3 B		_	2.0		13 1.1		
77/10/11 16 50 51.9				7 16			0.42		10,3 D					12 3.3		
77/10/19 2 16 2.9				r 21			0.12	1.0	4.0 C			1.8				81 7.5
77/10/24 7 \$ 47.6	56 17.3	160 39.0 1	76.2 4.5	F 15	3 309	472	0,48	21,1	23,6 D	145	1	7.4	54	26 34.0		
77/11/06 19 11 1.5					2 139		0.54	0,6			•	0.7	75			79 3.7
77/11/27 15 5 7.7					0 246			2.0	2.6 3			1.4		32 3,0		
77/12/16 21 49 23.4				F 15			0.24	0.7	1.9 B			1.4	51			41 3.7
77/12/27 15 9 51.7	60 19,4	153 28.0 1	75,2 4.8	r 15	1 91	39	0.22	0.8	2.8 C	315	0	1.5	45	2 1.2	225	88 5.2
													226			
78/01/28 2 25 2.1				r 16	1 74		0.12	1.5	9,1 D	320	_	2.7	336	7 1.2		83 17,2
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78/03/20 3 59 5.6				r 16			0.17		3.7 C			1.5		4 1.1		
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,,	35 40.2										-		- •			<del>-</del>

80/11/22 16 27 29.3 59 15.5 154 32.7 145.6 4.0F 16 3 148 42 0.20 1.0 1.2 A 350 6 1.5 85 40 1.0 253 49 2.7

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81/01/20 6 20 57.9	57 28.0	155 55.3	98.6 4.	OF 15	5	292	207	0.54	6.0	8.1 B			2.0	19				57 17.5
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81/03/21 23 1 40.2	58 52,6	154 39,3 17	22.6 4.	4F 15	3	197	66	0,20	1.3	1.5	14	19	1.6	117	33	0.9	259	51 3.5
81/05/21 20 29 34.1	59 45.5	152 50.7 10	03.9 4.	OF 15	1	144	48	0.20	1.0	1.7 m	144		1.3	54	14	1.8	238	76 3.3
81/08/01 1 42 18.4				6F 16	1	115	20	0.27	0.7	1.8 b	110	2	0.9	20	5	1.4	222	85 3.5
#1/09/16 12 11 6.0				18 17				0.28	1.7	4,6 C			3.1	79		2.0	234	80 8.7
<b>81/11/16 23 49 50.2</b>	60 5.1	153 1,4 17	18.6 4.	2 <b>F</b> 18	2	110	16	0.29	0.8	1.1 2	136	15	1.0	232	20	1.2	12	65 2,3
81/11/17 11 28 43.0				OF 16				0.31	0.6	1.6 B			0.6					79 3.1
82/02/26 7 16 59.2				57 17		70		0.35	0,0	1,8 \$				211			354	
82/02/27 12 18 9.1				OF 16				0.27	0.6	1.6 3				338				83 3.0
82/05/05 19 49 55.2				2F 18	_	47		0.26	0.4	1.0 A		_		161	6			,-
82/07/06 17 33 10.0	39 6.9	132 29.9	10.0 4.	1F 16	Q	123	23	0.24	0.7	1.7 8	252	1	V	162	2	1.4	,	86 3.3
82/07/14 12 15 51.7	£0 22 1	152 15 6 15	25 G A	7F 17	•	68	33	0.39	1.0	2.9 C	21	,	1 1	111		1 7	284	82 5.6
82/08/02 2 34 19.7				1F 17		124			2.7	2.0 C	63		1.5			5.0		40 2.9
82/08/10 16 25 42.1				OF 16	_	71	10	0.38	0.7	1.0 B	65			156			-	61 3.4
82/09/06 7 48 56.8				1F 15			124	0.72		36.3 D				350			168	
82/11/06 10 43 18.4				2x 11					50.0	50.0 D				356		99.0	0	
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83/01/01 11 18 8.8	61 16.9	146 58.3	25.2 4.	OF 16	1	75	38	0.25	0.4	3.1 C	293	1	0.4	23	1	0.7	158	89 5.8
83/03/15 23 53 22.1				4F 15	0	157	79	0.28	1,6	2.8 C	150	1	1,5	59	12	2.8	245	78 5.4
83/03/30 18 6 16,5	61 25.0	140 9.1	0.0 4.	1F 16	1	203	61	1.40	2,5	7.6 p	113	3	1.3	23	7	4.4	226	82 14.4
83/04/19 19 12 49.6	63 17,4	149 50.1 13	18.4 4.	7F 19	2	95	170	0.26	0.7	4.0 C	252	2	1.3	342	6	0.9	144	84 7.5
83/04/20 10 18 33.5	58 46.2	156 9.0 1	97.0 4.	6 <b>F</b> 15	2	250	243	0.36	5,3	6.9 D	129	5	2,5	222	25	8.9	28	64 13.6
83/04/23 6 13 48.2				7F 15				0.77	1.2	2.8 C				230		2.2	16	
83/06/16 21 19 25.5				OF 16	-			0.31	0,8	2.4 B						1.2		75 4.7
83/06/28 3 25 17.7				3F 15				0.26	0.5	0.5 A			0.5					48 1.1
83/07/12 15 10 3.6				38 18 2 <b>F</b> 15		56		0.34	0,3 3.8	2.5 B 8.3 D		_	5.1	283	_	0.4	76	88 4,6 66 17.1
83/07/21 17 35 17.0	28 6.1	152 28.5	<b>94.6</b> 4.	ZF 15	2	210	47	0.20	3.4	8.4 D	210	v	J. I		44	1.4	100	00 17.1
83/08/06 16 14 0.8	60 22 T	150 87 E 1	22 5 4	7r 16	1	140	23	0.31	1.2	3.0 C	331	0	1 1	241		2.1	61	82 5.7
83/08/19 4 59 33.6				2F 16		170		0.26	1.9	2.9 C		_	0.9		_		258	
83/09/07 19 22 3,6				38 14		70		0.32	0.4	0.9 A				106			316	
83/09/07 22 22 10.9				1X 14				0.33	0.4	1.5 B			0.5	16		0,8	151	
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83/09/26 23 52 54.4	57 8.3	156 5,1 1	22.4 4.	3F 16	- 6		226	0,72	7.7		59		-	125				-
63/10/06 11 10 12.6				7 <b>r 18</b>				0.34	0.7	1.8 8	93		0.9	_		1.2		77 3.4
84/01/14 11, 64 27.7	59 44,6	153 30.3 1	29.3 (.	17 18	3	39	57	0,25	1.0	2.5 8	127	1	1.7	217	19	1.2	34	71 4.9
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84/04/18 19 31 30.8				37 22		101 95		0.26	0.7	1.8 S			0.7	76 225		1.1	34	
84/04/20 4 24 49.3				1F 16		205	95	-	1.2	3.4 C		_		261		2.0	73	
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84/08/14 1 2 9.2				78 19		81	19		0.4	0.9 A	58			326				76 1.6
84/08/14 1 54 37.7				28 16		90		0,50	0.5	0.8 A	78			347			186	
84/09/20 4 17 24.4				1r 18				0,31	0.3	0.7 A	91			102				
85/01/09 19 28 21.1				7x 16			32	0.45	0.5	1.4 8	290	6	0.4	199	10	0.8	51	78 2.7

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85/03/03 13 38 50.5					90		0.35	0.7	1.0 B			1,4	64			211	81 3.	.5
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85/11/05 1 26 2.3					86													
85/11/16 7 11 13.1							0.41		1.6 B			0.9					75 3. 59 14.	
63/11/16 / 44 13.1	39 7.0	130 20.0	U 4.5X 9	,	414	133	4.19	14,3	14.3 D	313	•	4.6	32	31 3	1,0	222	33 14.	
85/12/08 7 32 23.6	50 51 6	150 16 0 35	7 4,3x 15	6	148	38	0.34	0.7	1.0 A	62	•	0.5	330	24	1.0	162	65 2.	
85/12/30 6 6 14.6							0.27	1.1	0.5 A			1.1					54 1.	
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86/05/23 23 18 43.7	58 47.4	153 15.6 75.	1 4,17 15	1	102	64	0.24	0.9	2,0 B	355	3	0.9	86	14	1.4	253	76 3.	. 9
86/07/28 14 31 14.6	60 35.7	150 26,8 40	0 4.4x 16	5	69	15	0,25	0.4	1.1 A	273	3	0.6	3	5	0.7	152	\$4 2.	. 0
86/09/18 20 56 7,7	61 45.7	149 43.6 41	2 4.9x 17	3	89	15	0.33	0.6	1,0 A	278	2	0.6	188	4	1.1	35	\$6 1.	. 9
86/09/29 17 13 33.9	58 49.4	156 7,0 217	8 4.1F 16	- 4	243	155	9.30	4,6	3.7 C	316	6	3,4	223	27	9.2	58	62 6.	. 2
86/10/09 1 21 7.8						49		0.7	1,7 =			0.6	-			171		
86/10/15 23 35 54.0					107		0.23	0.9	2.2 B			1.1	14			234		
86/10/26 22 19 45.3							0.13	5.1	9,2 D	60			329				82 17.	
86/10/27 19 39 36.7					45	-	0.18	0.3	0.9 A	-		0.6		-			84 1.	
86/11/04 6 14 18,9	61 24.1	151 55.4 96	2 4,6x 18	3	73	15	0.25	0.6	1,0 A	197	12	0,8	104	15	1.0	324	71 1.	. 9
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87/02/15 16 55 5.9					78		0.21	0.5	1.2 A	65	•		155				04 2.	
87/02/25 12 26 34.3					56		0,29	0.4		64	_		337				81 1.	
87/03/17 11 12 27.1					68		0.27	0,4	1.1 A							317		
87/04/18 2 1 39.7					83		0.28	0.6		234	_		144			333		
		322 22.0	A.	•							_			•				. –
87/04/18 2 13 39.7	61 21.5	150 39.1 53	7 4,1x 14	7	148	52	0.33	0,6	1,7 B	237	4	0,6	146	6	1,1	0	63 3.	. 2
87/05/11 11 30 11.3					86		0.15	1.6	2.8 C	24						294	66 5.	
87/06/10 1 44 42.0					115		0.29	0.6	2.3 B	177			267		1,0	65	85 4.	
97/06/13 13 9 23.4	60 44.3	147 36.1 26	3 4.5x 19	- 5	60	31	0.33	0.3	1.5 m	225	1	0.6	316	3	0.6	117	87 3.	. 6
87/06/20 7 12 38.8	58 30.0	156 15.6 198	8 4.0F 15	4	149	186	0.15	3.0	4,4 C	333	1	2.0	64	12			78 8,	. 4
87/07/25 1 11 50.9					82		0,34	1.2	2,7 C							205		
47/09/09 9 35 29.4				_	90		0.40	0.9	2,2 3				249		1.7		<b>81</b> 4.	
87/10/24 2 49 46,0					107		0.71	1.1	2.8 8						0.8		65 4.	
87/10/24 5 6 45.6					46		0.40	0.3	1.5				193					-
07/10/31 1 27 59.8	60 14.1	149 55.6 61	3 4.6x 21	10	105	30	0.42	0.5	1,1 A	257	0	0,6	347	3	1,0	167	87 2.	.0
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89/05/22 3 16 14.0 63 20.5 150 15.0 13.1 4.0x 17 4 102 41 0.23 0.4 0.5 A 187 2 0.6 277 12 0.7 88 78 1.0 89/05/22 4 18 31.1 64 55.8 156 18.5 10.0\* 4.2x 14 1 180 152 0.43 3.3 9.3 D 49 3 5.2 140 13 4.8 306 77 17.9 89/05/31 0 46 40.5 58 57.3 154 22.8 134.8 4.1x 15 12 119 26 0.23 1.7 2.1 B 168 3 1.8 76 34 2.6 262 56 4.4

PRELIMINARY DETERMINATION OF HYPOCKETERS IN SOUTHERN ALASEA - OCTOBER 1871 TO MAY 1889 origin time lat n long w depth mag mp no gap dl mas seh sex g aal dipl set as2 dip2 se2 as1 dip1 se1

88/09/30 4 50 3.9 61 7.4 151 30.1 85.1 5.0x 15 7 64 30 0.31 0.7 1.6 B 136 12 1.0 45 13 0.7 269 72 3.1

deg km sec km km deg deg km deg deg km deg deg km

date hr mn see deg min deg min km

### APPENDIX B

### List of Previously Published Catalogs

- Lahr, J. C., Page, R. A., and Thomas, J. A., 1974, Catalog of earthquakes in south central Alaska, April-June 1972, U.S. Geological Survey Open-File Report, 35 p.
- Fogleman, K. A., Stephens, Christopher, Lahr, J. C., Helton, S. M., and Allan, M. A., 1978, Catalog of earthquakes in southern Alaska, October-December 1977, U.S. Geological Survey Open-File Report 78-1097, 28 p.
- Stephens, C. D., Lahr, J. C., Fogleman, K. A., Allan, M. A., and Helton, S. M., 1979, Catalog of earthquakes in southern Alaska, January-March 1978, U.S. Geological Survey Open-File Report 79-718, 31 p.
- Stephens, C. D., Astrue, M. A., Pelton, J. R., Fogleman, K. A., Page, R. A., Lahr, J. C., Allan, M. A., and Helton, S. M., 1982, Catalog of earthquakes in southern Alaska, April-June 1978, U.S. Geological Survey Open-File Report 82-488, 36 p.
- Stephens, C. D., Lahr, J. C., Fogleman, K. A., Helton, S. M., Cancilla, R. S., Tam, Roy and Baldonado, K. A., 1980, Catalog of earthquakes in southern Alaska, October-December 1979, U.S. Geological Survey Open-File Report 80-2002, 53 p.
- Stephens, C. D., Fogleman, K. A., Lahr, J. C., Helton, S. M., Cancilla, R. S., Tam, Roy and Freiberg, J.A., 1980, Catalog of earthquakes in southern Alaska, January-March 1980, U.S. Geological Survey Open-File Report 80-1253, 55 p.
- Fogleman, K. A., Stephens, C. D., Lahr, J. C., Rogers, J. A., Helton, S. M., Cancilla, R. S., Tam, Roy, Freiberg, J. A., and Melnick, J. P., 1983, Catalog of earthquakes in southern Alaska, July-September 1980, U.S. Geological Survey Open-File Report 83-15, 54 p.
- Fogleman, K. A. Stephens, C. D., Lahr, J. C., and Rogers, J. A., 1986, Catalog of earthquakes in southern Alaska for 1984, U.S. Geological Survey Open-File Report 86-99, 106 p.
- Fogleman, K. A., Stephens, C. D., and Lahr, J. C., 1988, Catalog of earthquakes in southern Alaska for 1985, U. S. Geological Survey Open-File Report 88-31, 113 p.

#### APPENDIX C

## Summary of Timing Criteria

### Routine network processing.

October 1, 1971 - September 31, 1973

Time all shocks with:

A. Cook Inlet (west of longitude 150°W):

S-P  $\leq$  25 s and recorded at 6 or more stations (clearly recorded at 4 or more of these).

- B. Valdez (east of 150°W):
  - S-P ≤ 20 s and clearly recorded at 4 or more stations including 2 of WLM, ERN, VLZ and CVA.

OF

- S-P ≤ 10 s and clearly recorded at 3 stations including 2 of WLM, ERN, VLZ and CVA.
- C. Well recorded regional events that have clipped traces and do not come from the Aleutians.

October 1, 1973 - September 31, 1974

Time all earthquakes with:

A.  $F-P \ge 30$  s. F-P is signal duration (see section on magnitude).

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B.  $F-P \ge 15$  s and  $S-P \le 5$  s.

October 1, 1974 - January 31, 1975

Time all local events in network with F-P  $\geq$  20 s.

February 1, 1975 - September 31, 1977

Time all shocks with:

- A. WEST (west of 150°W): F-P  $\geq$  80 s.
- B. CENT (between 150°W and 145°W): F-P  $\geq$  20 s.
- C. EAST (east of 145°W): All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals.

October 1, 1977 - September 31, 1980

Time all shocks within the area 58-64°N and 134-156°W (eastern border moved to 134°W from 136°W) with:

- A. WEST:  $F-P \ge 80 \text{ s.}$
- B. CENT:  $F-P \ge 20 \text{ s}$ .
- C. EAST: All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals on SCAN and/or EAST films.

October 1, 1980 - June 30, 1981

Time all shocks within the area 58-64°N and 134-156°W with:

- A. WEST (West of 145°W): F-P ≥ 30 s (WEST and CENT regions combined)
- B. EAST: All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals on SCAN and/or EAST films.

## July 1, 1981 - September 30, 1981

Time all shocks within the area 58-64°N and 134-156°W with:

A. WEST:  $F-P \ge 30 \text{ s.}$ 

#### B. EAST:

1. All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals on SCAN film only.

or

If an event has a coda on the SCAN film F-P ≥ 10 s but does not have 4 P's or 3
P's and 1 S, check the EAST film for additional arrivals and time if have 4 P's
or 3 P's and 1 S on SCAN plus EAST.

or

3. If at any time less than five stations are operating east of 145°W use both SCAN plus EAST films to see if have 4 P-arrivals or 3 P-arrivals and 1 S-arrival.

## December 1, 1980 - May 31, 1989

No rereads on magnitude M < 1.0 events unless solution is unacceptable.

#### October 1, 1981 - March 31, 1984

Network borders for timing reduced to decrease the number of earthquakes processed per month. Northern border moved from 64°N to 63°N. Less emphasize placed on rereading events between latitude 58-59°N and longitude 134-138°W.

Time all shocks within the area 58-63°N and 134-156°W with:

A. WEST:  $F-P \ge 30 \text{ s.}$ 

#### B. EAST:

1. All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals on SCAN film only.

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2. If an event has a coda on the SCAN film F-P ≥ 10 s but does not have 4 P's or

 If an event has a coda on the SCAN film F-P ≥ 10 s but does not have 4 P's or 3 P's and 1 S, check the EAST film for additional arrivals and time if have 4 P's or 3 P's and 1 S on SCAN plus EAST.

OL

3. If at any time less than five stations are operating east of 145°W use both SCAN plus EAST films to see if have 4 P-arrivals or 3 P-arrivals and 1 S-arrival.

### April 1, 1984 - August 31, 1985

Time all shocks within the area 58-62.5°N (northern border moved from 63.0°N to 62.5°N) and 134-156°W with:

A. WEST:  $F-P \ge 30 \text{ s}$ .

### B. EAST:

1. All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals on SCAN film only.

or

 If an event has a coda on the SCAN film F-P ≥ 10 s but does not have 4 P's or 3 P's and 1 S, check the EAST film for additional arrivals and time if have 4 P's or 3 P's and 1 S on SCAN plus EAST.

or

- 3. If at any time less than five stations are operating east of 145°W use both SCAN plus EAST films to see if have 4 P-arrivals or 3 P-arrivals and 1 S-arrival.
- C. Time any earthquake with S-P  $\leq$  2 s on SCAN film station except for AGA whose S-P time must be  $\leq$  1.75 s.

### September 1, 1985 - May 31, 1989

Time all shocks within the area 58-62.5°N and 138-156°W (eastern borders moved from 134°W to 138°W due to major reduction of recordable stations in east) with:

A. WEST:  $F-P \ge 30 \text{ s.}$ 

#### B. EAST:

1. All events with 3 clearly recorded P-arrivals and 1 S-arrival or 4 P-arrivals on SCAN film only.

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 If an event has a coda on the SCAN film F-P ≥ 10 s but does not have 4 P's or 3 P's and 1 S, check the EAST film for additional arrivals and time if have 4 P's or 3 P's and 1 S on SCAN plus EAST.

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- 3. If at any time less than five stations are operating east of 145°W use both SCAN plus EAST films to see if have 4 P-arrivals or 3 P-arrivals and 1 S-arrival.
- C. Time any earthquake with S-P  $\leq$  2 s on SCAN film station except for AGA whose S-P time must be  $\leq$  1.75 s.

# Special Studies

1. Special study of microearthquakes around Anchorage,

Time all shocks with:

January 1, 1979 - December 31, 1979

A.  $S-P \le 13.5$  s at KNK

and

B. P-arrival at SPU or MSP or SKL before TOA and VZW.

January 1, 1980 - March 31, 1984

A.  $S-P \le 10$  s at PMS

and

B. at least one station other than PMS with a F-P coda above 1 cm pk-to-pk for at least 5 s.

April 1, 1984 - May 31, 1989

A. S-P  $\leq$  5 s at PMS or SSN

and

- B. At least one station other than PMS with a F-P coda above 1 cm pk-to-pk for at least 5 s.
- 2. Special criterion for reduction in timing of aftershocks of February 28, 1979 St. Elias earthquake.

April 1 - September 30, 1979

A. If CHX is operational and YAH or PIN is working, do not time shocks where:

1. CHX S-P < 6 s.

and

2. YAH or PIN F-P < 40 s (changed to 25 s for May - September).

and

- 3. GYO P-arrival ≤ 4 s before CHX P-arrival.
- B. If CHX is not working but GYO is, do not time events where:
  - 1. GYO S-P < 7 s.

and

2. YAH or PIN F-P < 25 s.

and

3. GYO P-arrival is before SSP P-arrival.

- C. If CHX and GYO are out, do not time events where:
  - 1. SSP and PIN P-arrivals are approximately the same time and
  - 2. SSP and PIN S-P < 10 s.

### May 2-19, 1982

Special criterion for large aftershock sequence in St. Elias aftershock zone. Time only aftershocks with a measureable coda seen on SCAN film station or if YAH amplitude ≥ 20 mm pk-to-pk. This coda crtiteria only applied to shocks where YAH, WRG, and PIN were first three stations seen on SCAN and their S-P intervals were about 7 to 9 s.

### April 1, 1984 - September 4, 1985

Do not locate events near Icy Bay with:

- A. S-P  $\leq$  4 s at AGA (or S-P  $\leq$  5 s at CHX if AGA out) and
- B. F-P at YAH and PIN < 10 s. (or F-P < 8 s at AGA if YAH and PIN are dead)

### 3. Bradley Lake array on southern Kenai Peninsula

November 27, 1980 - January 31, 1981

Time any event with S-P  $\leq$  10 s at any Bradely Lake station (BRLK, BRNE, BRNW, BRSE, and BRSW).

February 1, 1981 - September 31, 1981

Time any shock with S-P ≤ 12 s at any Bradley Lake station.

October 1, 1981 - February 28, 1983

Time any shock with S-P  $\leq$  12 s at any Bradley Lake station excluding events which arrive at ILM or RDT before SLV, or BRNE, or BRSW, or BRSE.

March 1, 1983 - February 28, 1985

Time any earthquake with S-P  $\leq$  6 s at one of the Bradley Lake stations or SLV or SWD.

March 1, 1985 - September 19, 1986

Time all events with S-P  $\leq 3.5$  s at BRLK.

## 4. Cook Inlet volcanoes

July 1, 1981 - May 31, 1989

Time any earthquake with S-P  $\leq$  5 s at SPU, RDT, or ILM on SCAN film.

## 5. Aftershocks of August 14, 1984, Sutton m, 5.7 event

September 11, 1984 - April 30, 1986

Time all earthquakes with GHO S-P  $\leq$  4 s and KNK F-P  $\geq$  8 s on SCAN film.

# 6. Knight Island in Prince William Sound

July 7, 1985 - August 31, 1985

Timed all earthquakes which:

A. Have O wt. P arrivals on KNI, LOU, GBY

and

B. KNI P-arrival is before GLI and MTU P-arrivals

and

C. GBY S-P  $\leq 8$  s.

and

D. A measurable code on KNI or LOU or GBY or pk-to-pk amplitude ≥ 10 mm on all three stations (KNI, LOU, GBY).

September 1, 1985 - December 31, 1987

Time all shocks with KNI S-P  $\leq$  7 s and F-P  $\geq$  8 s.

## 7. Montague Island in southern Prince William Sound

September 1, 1985 - April 30, 1986

Time all events with MTU S-P  $\leq 5$  s.

### 8. Joint USGS/GIUA data set

January 1, 1989 - May 31 1989

Starting in 1989, the Geophysical Institute of the University of Alaska (GIUA), Fairbanks began recording 13 USGS stations along with their own network. The GIUA locates all events triggered on by their automatic detection and recording system regardless of magnitude or location. Consequently, the GIUA locates shocks within the USGS routine processing borders which would normally not be processed by the USGS. Duplicate readings from USGS stations taken by GIUA staff for earthquakes also processed by the USGS are weighted out.

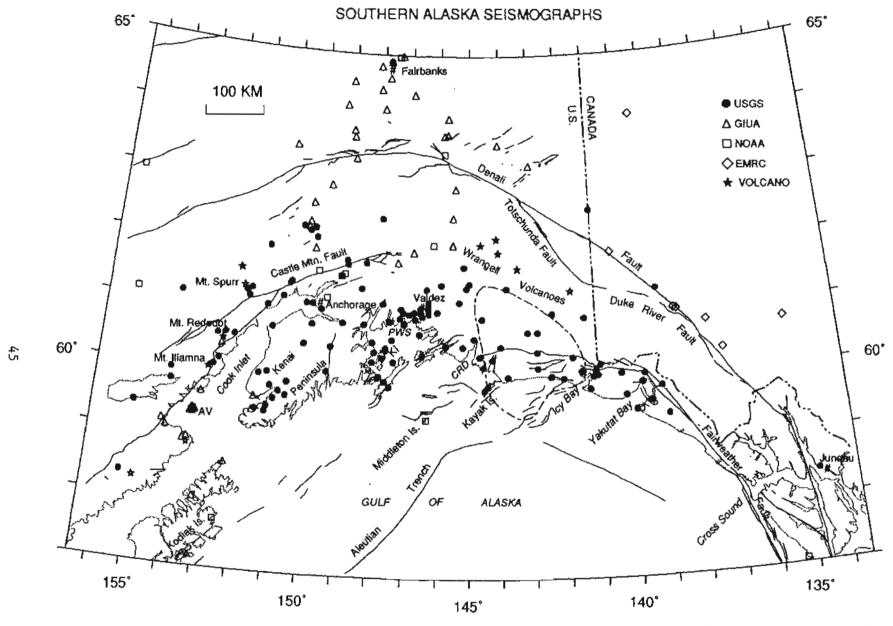
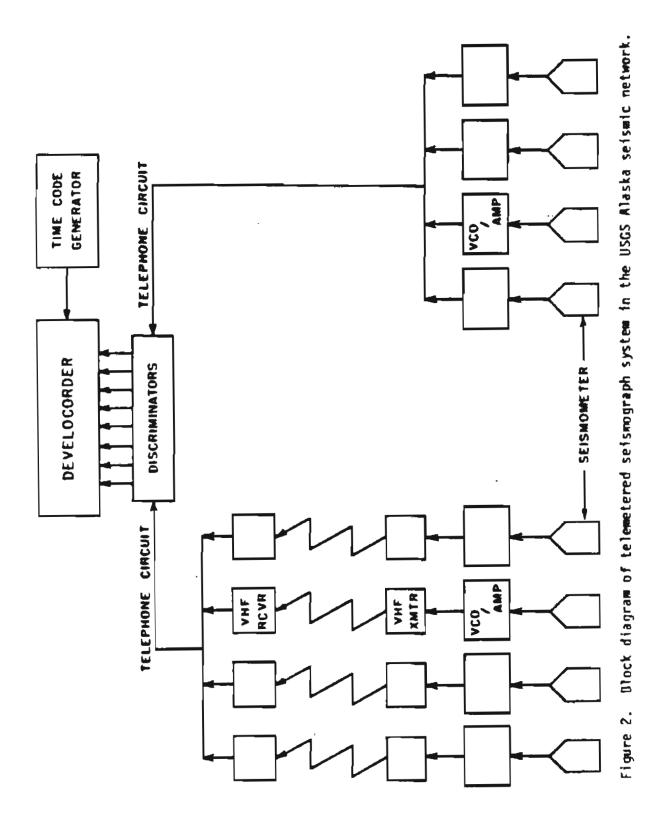


Figure 1. Map showing the locations of all USGS seismograph stations in southern Alaska along with stations operated by other institutions used in the prepartation of this catalog. Contour with alternating long and short dashes outlines inferred extent of Yakataga seismic gap. Neogene and younger faults (George Plafter, personal communication, 1988) are shown as solid lines. AV, Augustine volcano; CRD, Copper River Delta; KI, Knight Island; PWS, Prince William Sound. Quaternary volcanoes are indicated by stars.



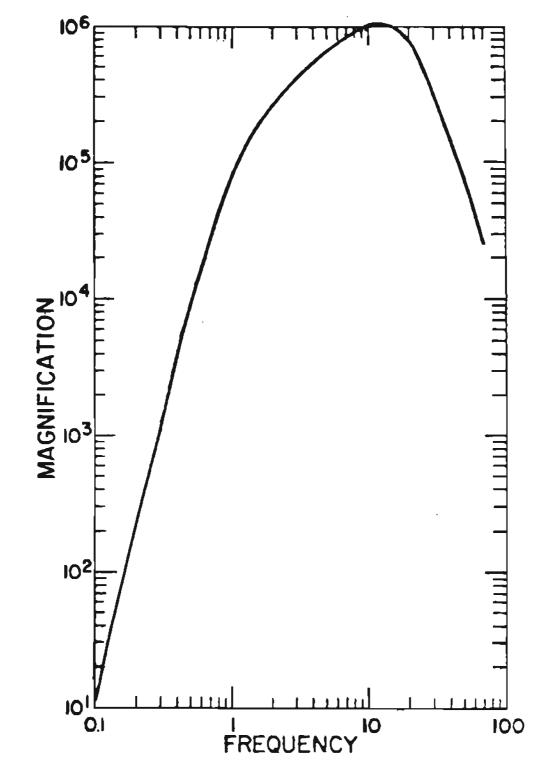


Figure 3. System response curve for typical USGS Alaska seismographs that incorporate the AlVCO unit.

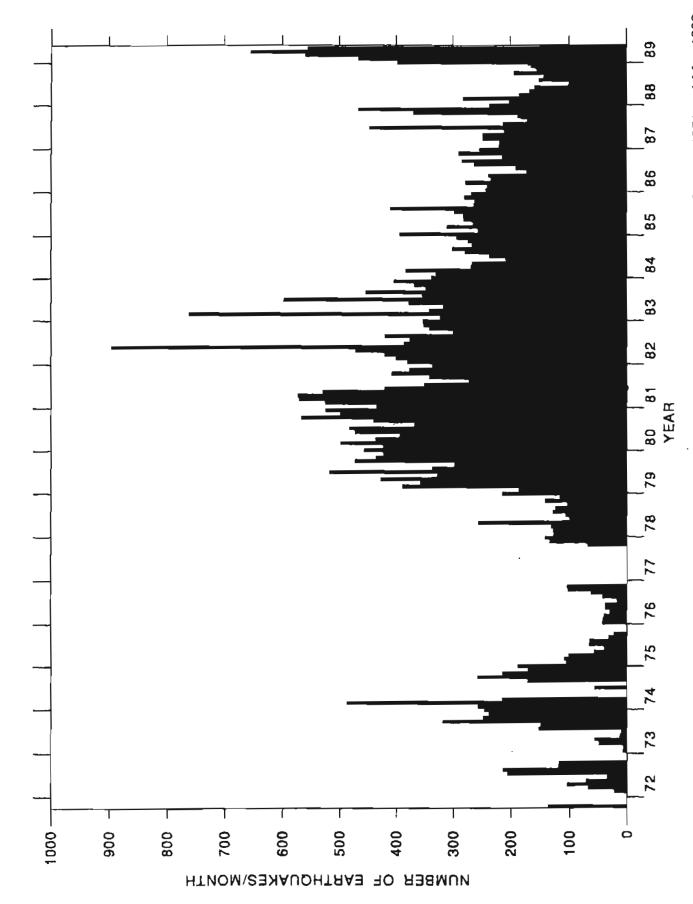


Figure 4. Histograph showing the number of earthquakes per month located by the network between October 1971 and May 1989.

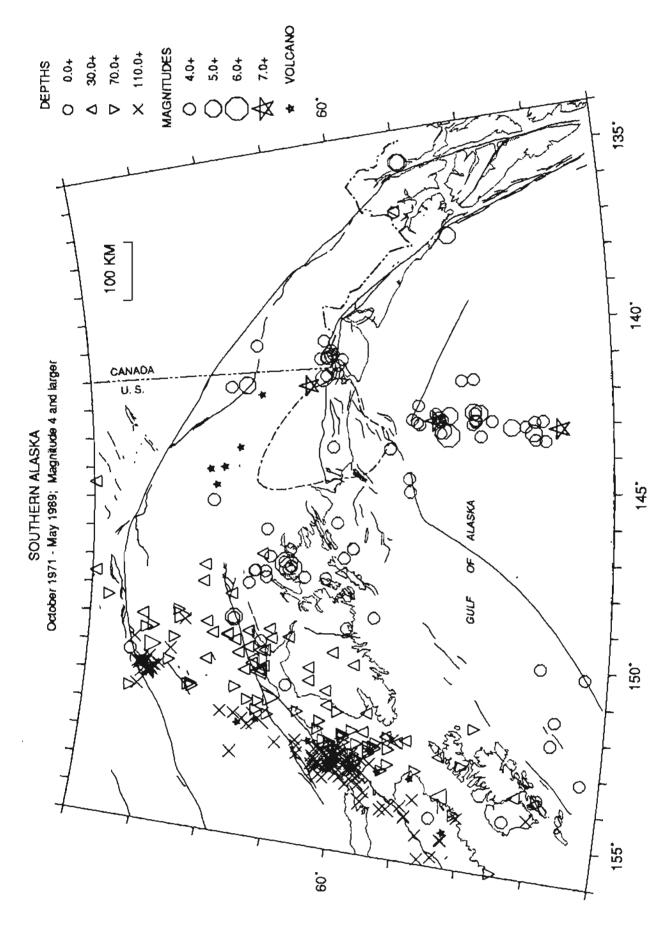


Figure 5. Epicenters of 307 earthquakes of magnitude of 4.0 or larger that occurred between October 1971 and May 1989. See Figure 1 for details about identification of map features.

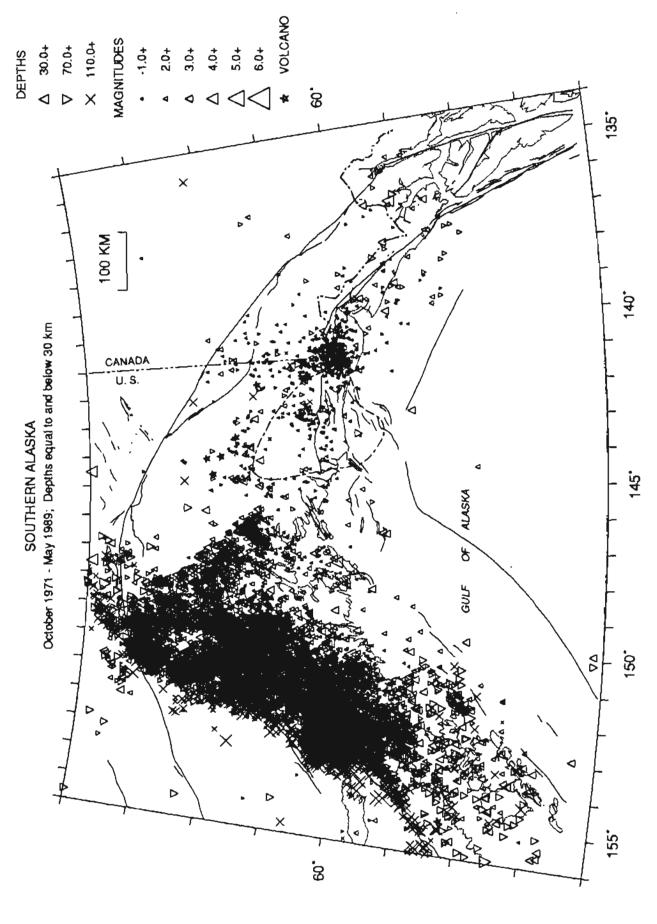


Figure 6. Epicenters of 14,013 earthquakes with depths equal to and below 30 km that occurred between October 1971 and May 1989.

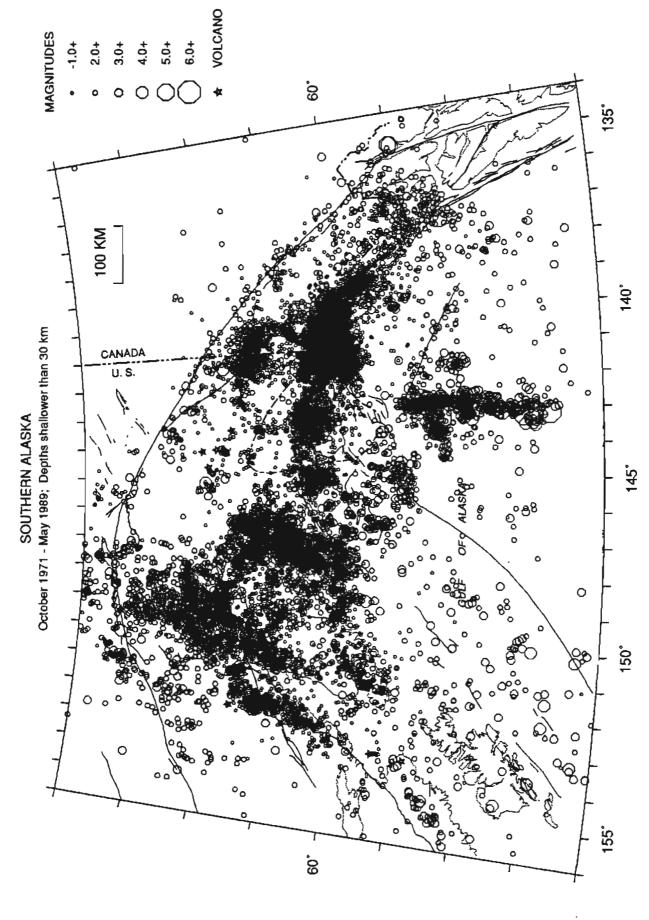


Figure 7. Epicenters of 35,225 earthquakes with depths shallower than 30 km that occurred between October 1971 and May 1989.

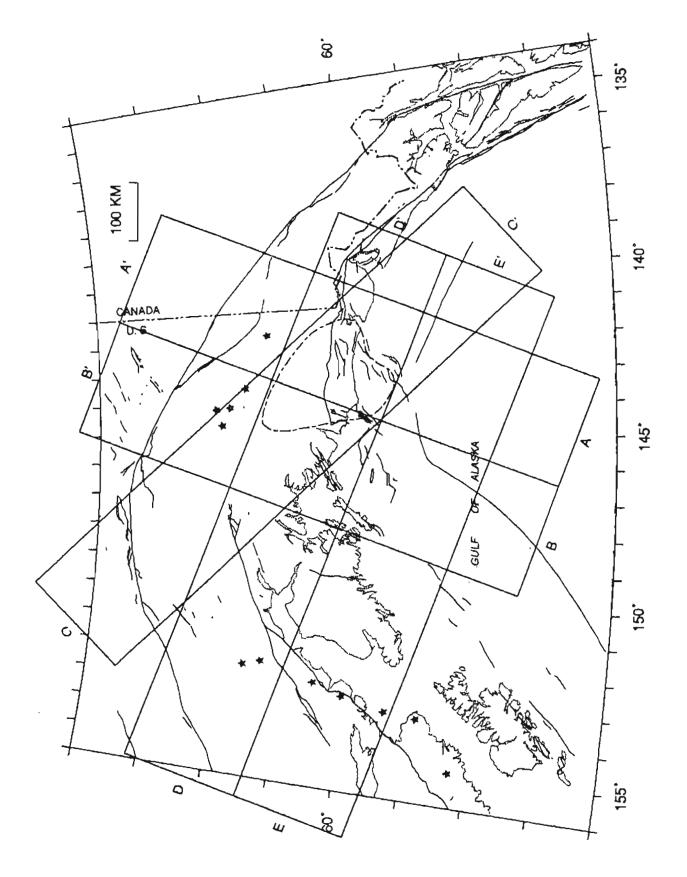


Figure 8. Reference map showing the areas represented in the cross sections in Figure 9.

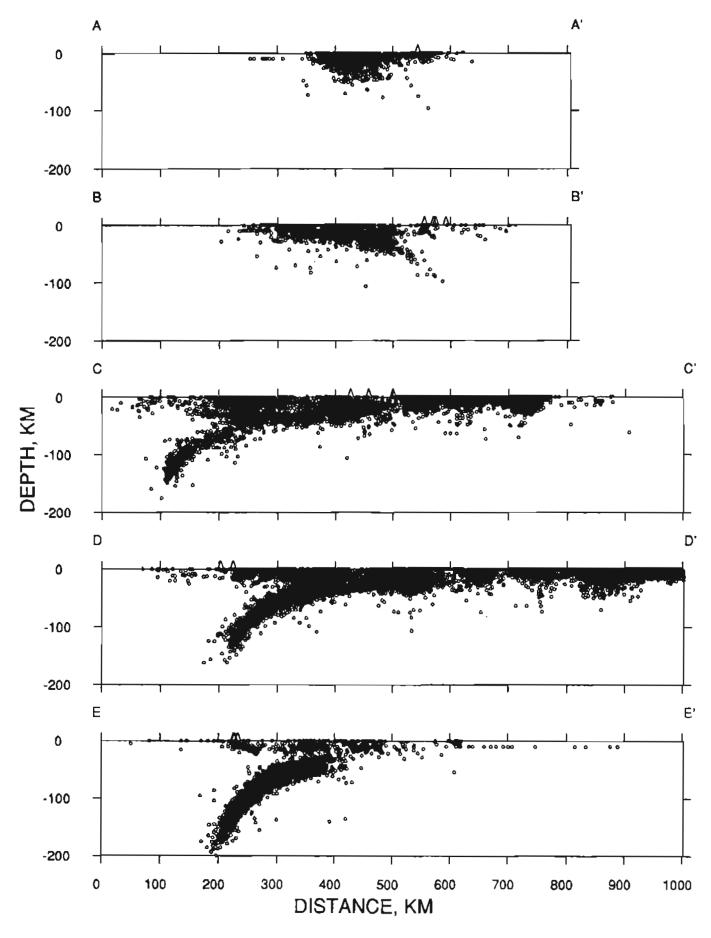


Figure 9. Sections of better-constrained hypocenters (SEH  $\leq$  5 km and SEZ  $\leq$  10 km) for areas indicated in Figure 8. Quaternary volcanoes plotted as triangles above zero depth. No vertical exaggeration.

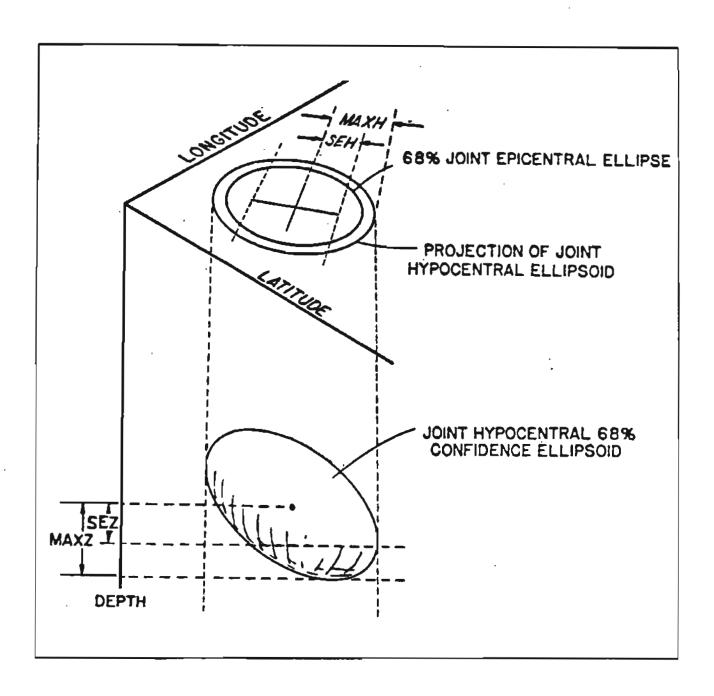


Figure 10. Relationship between the confidence ellipsoid and SEH, MAXH, SEZ, and MAXZ. The projected ellipse has the same orientation and eccentricity as the joint epicentral 68-percent confidence region, but is 1.23 times larger. The error ellipsoid is calculated assuming a constant standard deviation of 0.1 sec for the arrival time readings.